

Conference on *Friction, Fracture and Earthquake Physics*,
Kavli Institute for Theoretical Physics, Santa Barbara, 15-19 August 2005

Dynamics of rupture through branched and offset fault systems

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Alexei N. B. Poliakov (Royal Bank of Canada, London)

James R. Rice (Harvard)

Ares J. Rosakis (Caltech)

Carl E. Rousseau (URI)

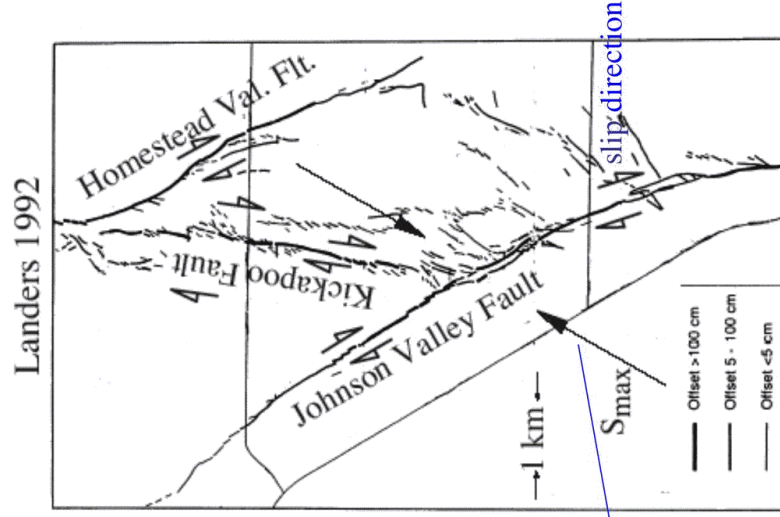
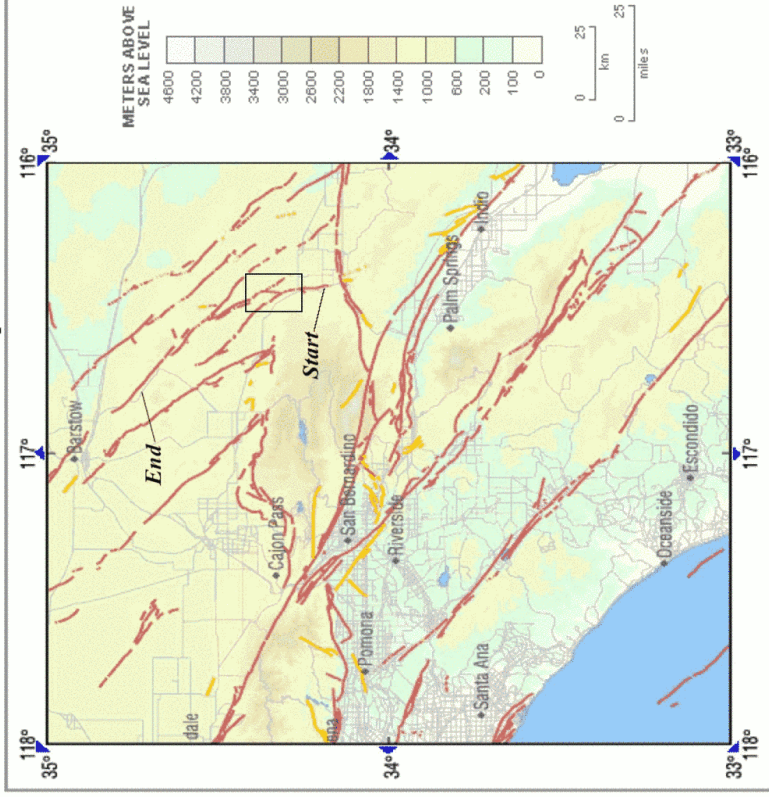
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Elizabeth L. Templeton (Harvard)

Questions

- How the earthquake rupture chooses its path through geometric complexities like bends, branches and stepovers? When and why a fault branch might be preferred? Would the rupture continue as well along the main fault?
- Could the directivity of a complex earthquake be inferred from a pattern of fault branches it ruptured?
- How do small fault branches interact with the main rupture propagation? Could they arrest rupture? How they contribute to seismic radiation?
- What is the off-fault damage pattern related to supershear (versus sub-Rayleigh) rupture propagation?
- Can laboratory experiments be used to constrain branching theory?

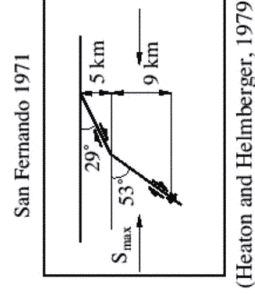
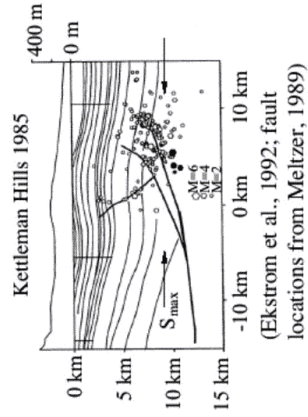
Faults in Southern California, Mohave Area (from USGS)
Start and End of 1992 Landers rupture shown



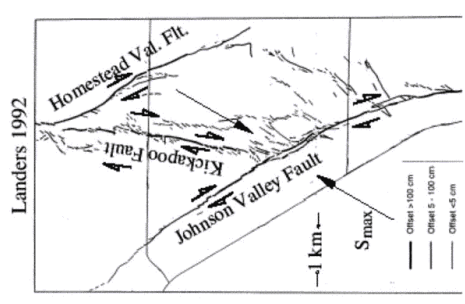
Portion of 1992
Landers earthquake
rupture path
(transition from
Johnson Valley to
Kickapoo fault,
then to Homestead
Valley fault).

Fault map and slip,
Sowers et al. (1994).

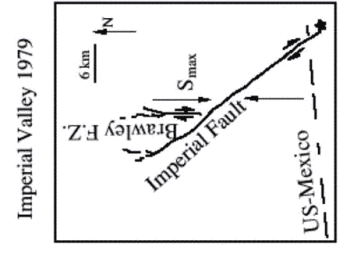
Pre-stress direction,
Hardebeck and
Hauksson (2001).



(Poliakov, Dmowska and Rice, *JGR*, 2002; Kame, Rice and Dmowska, *JGR*, 2003)



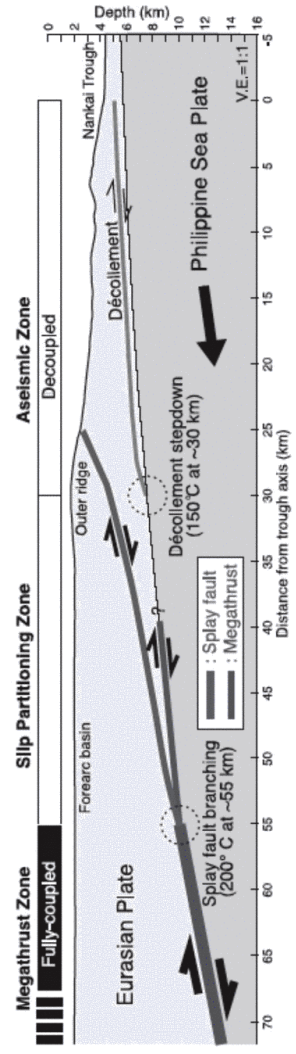
(fault map from Sowers et al., 1994; stress from Hardebeck and Hauksson, 2001)



(Archuleta, 1984; stress guessed from Hardebeck and Hauksson, 1999, to NW)

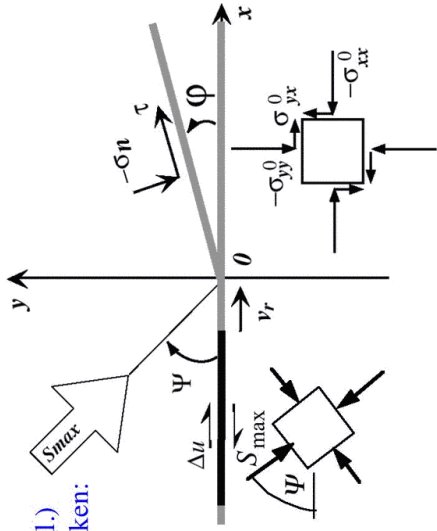
Nankai Trough, area of 1944 Mw 8.1 Tonankai earthquake (Park et al., *Sci.*, 2002; Nakanishi et al., *JGR*, 2002)

Splay fault branching. Relation to tsunami generation?

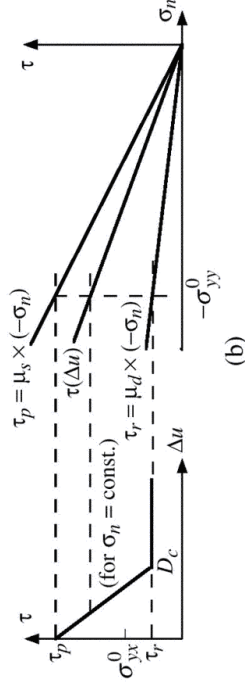


Theory and computational modeling (Poliakov et al., *JGR*, 2002; Kame et al., *JGR*, 2003)

Parameters argued (Poliakov et al.) to control whether branch path taken:
 φ -- branch angle
 Ψ -- direction of max. pre-stress
 v_r -- rupture speed at junction

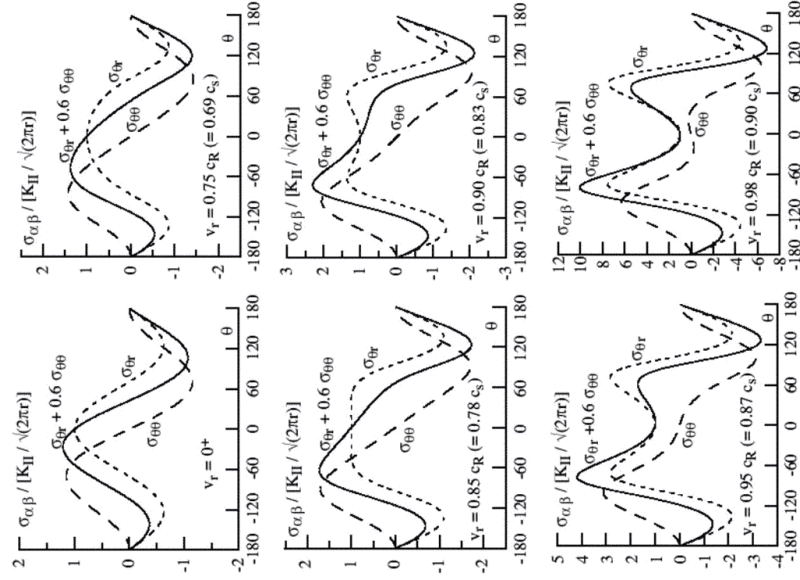
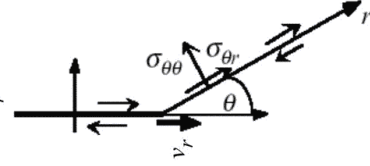


Slip-weakening model of frictional failure (Kame et al.);
 τ = shear stress,
 σ_n = normal stress,
 Δu = fault slip.



Perspective from singular elastic crack theory on the expected importance of rupture propagation speed v_r at the branch junction (Poliakov et al., *JGR*, 2002)

Plotted:
 Singular $1/\sqrt{r}$ stress terms, normalized by $K_{II} / \sqrt{(2\pi r)}$ ($=\sigma_{\theta r} |_{\theta=0}$), for various rupture speeds v_r .

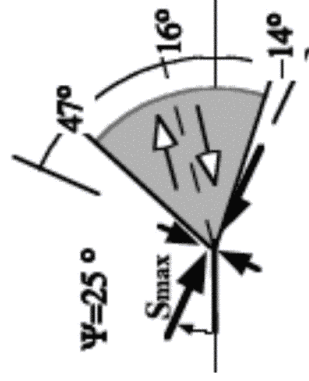
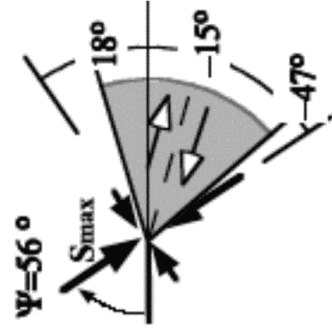


(Poliakov et al., *JGR*, 2002; Kame et al., *JGR*, 2003)

Once initiated in the high stress region, can a branched rupture become large?

Importance of direction Ψ of maximum principal compression in pre-stress field

steep pre-stress angle favors extensional side



shallow pre-stress angle favors compressional side

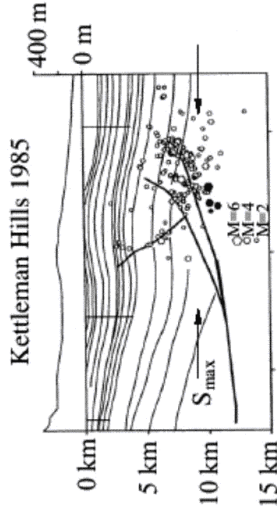
Dashed lines: Directions of maximum $\tau^0 / (-\sigma_n^0)$

Shaded regions: Sectors where $\tau^0 / (-\sigma_n^0) > \mu_d$ (= dynamic friction coef.)

(Poliakov, Dmowska and Rice, *JGR*, 2002)

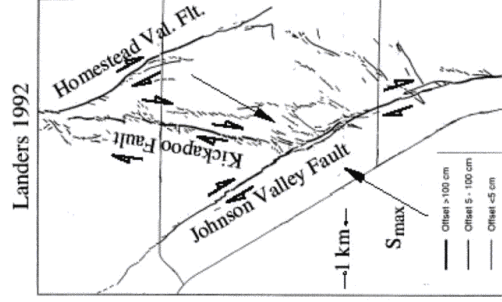
Correlation with natural examples

Depth cross-section view: Shallow S_{max} direction, $\Psi \approx 12-18^\circ$; secondary failures on compressional side:



(Ekstrom et al., 1992; fault locations from Meltzer, 1989)

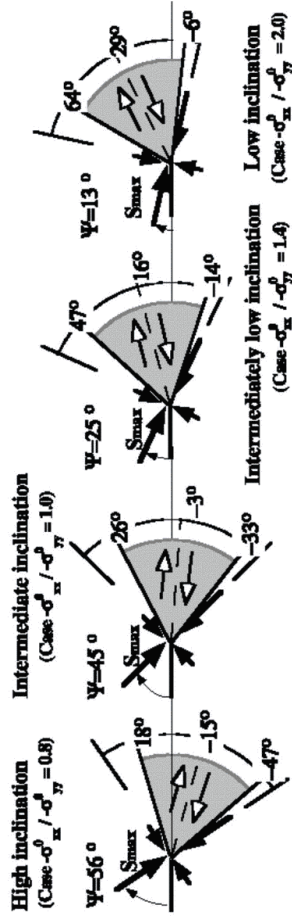
Map view: Steep S_{max} direction, $\Psi \approx 60^\circ$; secondary failures on extensional side:



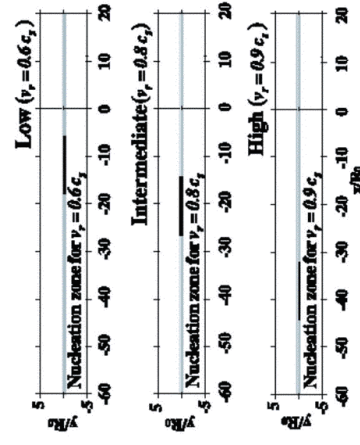
(fault map from Sowers et al., 1994; stress from Hardebeck and Hauksson, 2001)

(Kame et al., JGR, 2003)

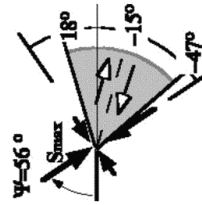
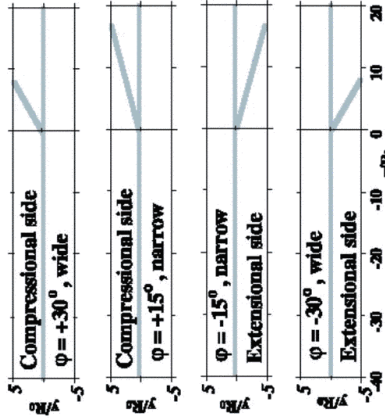
(a) Pre-stress state: Inclination of S_{max}



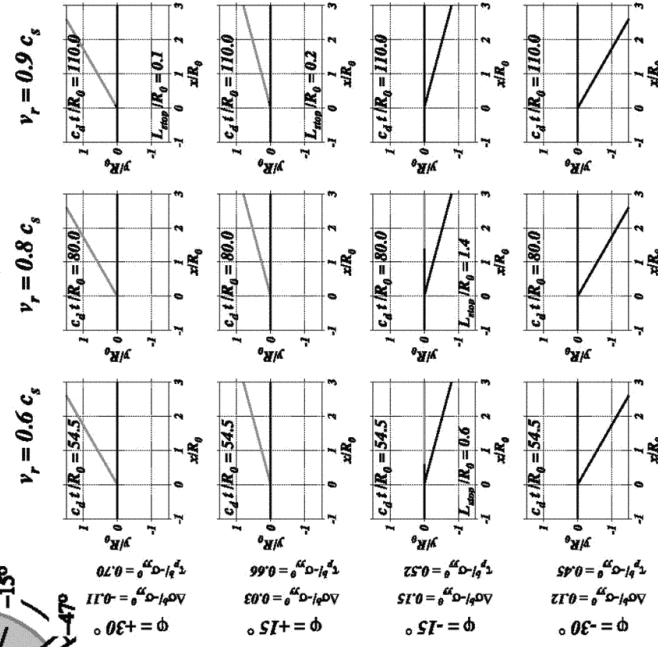
(b) Rupture velocity v_r when reaching the intersection



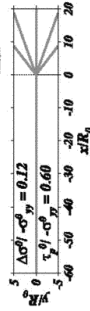
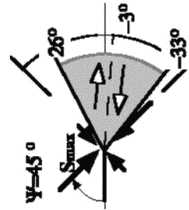
(c) Branching angle ϕ



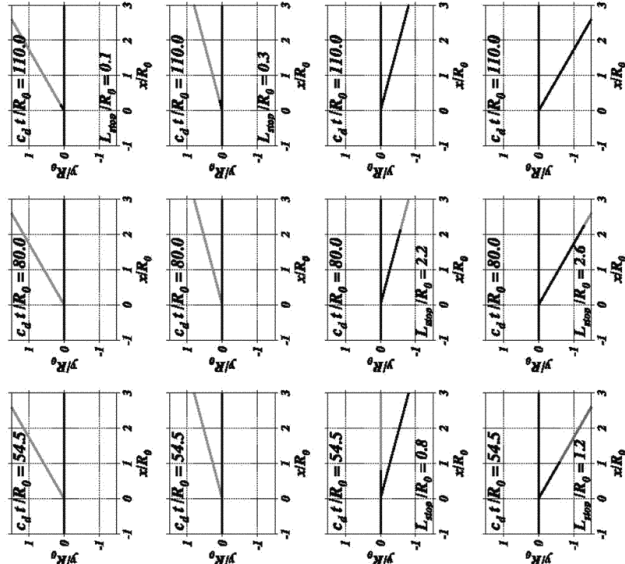
High Inclination of S_{max} , $\psi = 56^\circ$



Intermediate Inclination of S_{max} , $\Psi = 45^\circ$

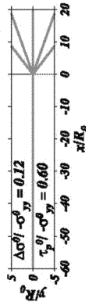
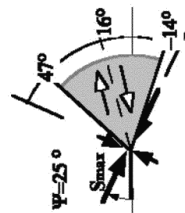


$v_r = 0.6 c_s$ $v_r = 0.8 c_s$ $v_r = 0.9 c_s$

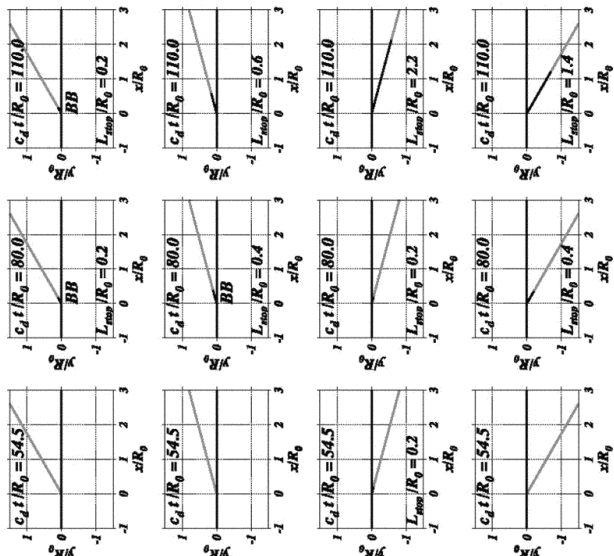


$\phi = +30^\circ$ $\Delta\phi^1 - \alpha_y^0 = -0.03$ $\xi_y^1 - \alpha_y^0 = 0.73$
 $\phi = +15^\circ$ $\Delta\phi^1 - \alpha_y^0 = 0.07$ $\xi_y^1 - \alpha_y^0 = 0.67$
 $\phi = -15^\circ$ $\Delta\phi^1 - \alpha_y^0 = 0.10$ $\xi_y^1 - \alpha_y^0 = 0.53$
 $\phi = +30^\circ$ $\Delta\phi^1 - \alpha_y^0 = 0.03$ $\xi_y^1 - \alpha_y^0 = 0.48$

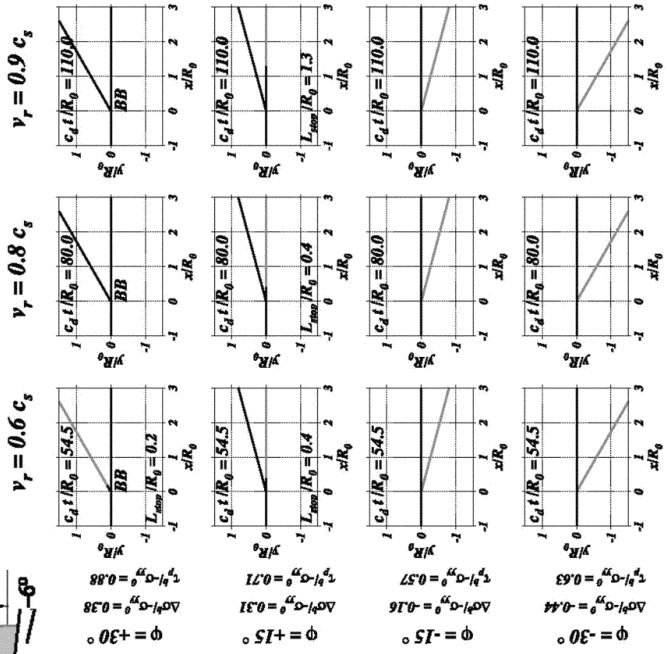
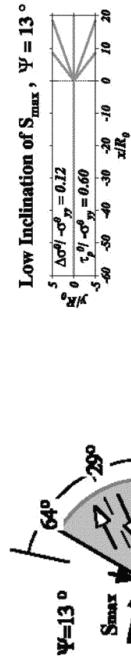
Intermediately Low Inclination of S_{max} , $\Psi = 25^\circ$



$v_r = 0.6 c_s$ $v_r = 0.8 c_s$ $v_r = 0.9 c_s$

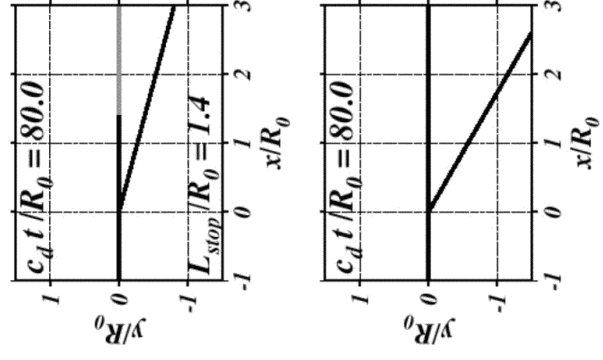


$\phi = +30^\circ$ $\Delta\phi^1 - \alpha_y^0 = 0.14$ $\xi_y^1 - \alpha_y^0 = 0.79$
 $\phi = +15^\circ$ $\Delta\phi^1 - \alpha_y^0 = 0.17$ $\xi_y^1 - \alpha_y^0 = 0.69$
 $\phi = -15^\circ$ $\Delta\phi^1 - \alpha_y^0 = 0.00$ $\xi_y^1 - \alpha_y^0 = 0.54$
 $\phi = +30^\circ$ $\Delta\phi^1 - \alpha_y^0 = -0.16$ $\xi_y^1 - \alpha_y^0 = 0.54$

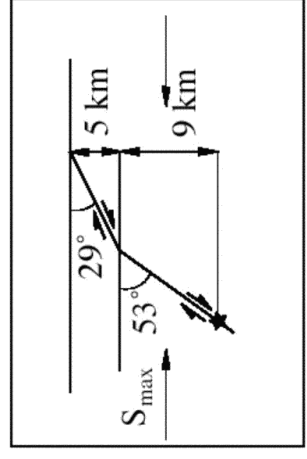


(Kame et al., JGR, 2003)

$\Psi = 56^\circ$, $v_r = 0.8 c_s$

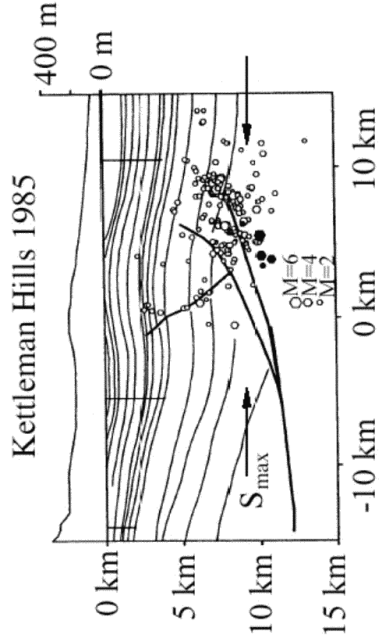


San Fernando 1971

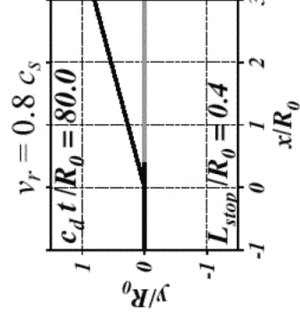
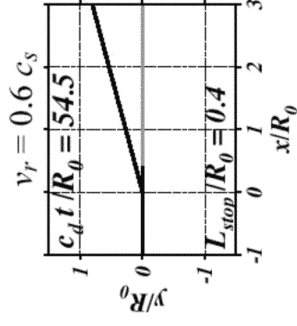


(Heaton and Helmberger, 1979)

(Kame et al., JGR, 2003)



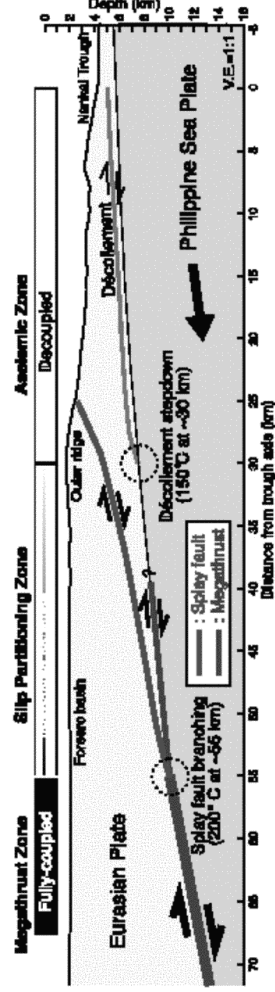
$\Psi = 13^\circ$



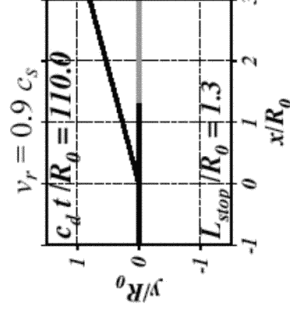
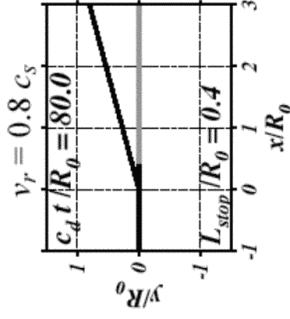
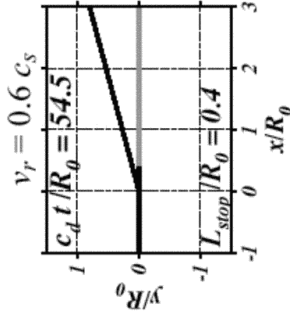
(Ekstrom et al., 1992; fault locations from Meltzer, 1989)

(Kame et al., JGR, 2003)

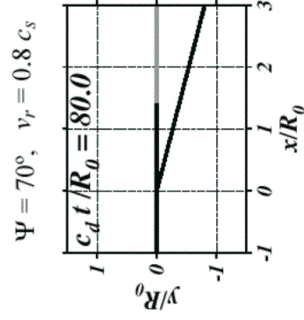
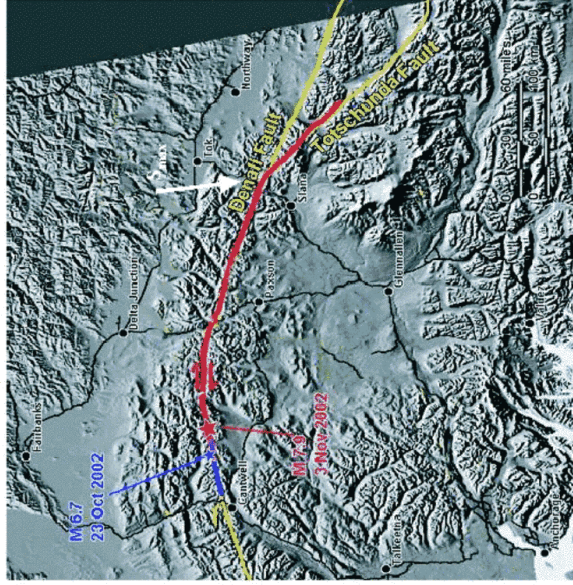
Nankai Trough, area of M 8.1 Tonankai 1944
(Park et al., Science, 2002; Nakanishi et al., JGR, 2002)



$\Psi = 13^\circ$



M 7.9 Denali, 3 November 2002 -- Branch from Denali Fault to Totschunda Fault
 (Bhat, Dmowska, Rice and Kame, Fall 2002 AGU abstract S72F-1364)
 (and *Bull. Seismol. Soc. Amer.*, 2004)



Stress direction from Ratchkovski and Hansen, 2002; Nakamura et al., 1980; Estabrook and Davies, 1991.

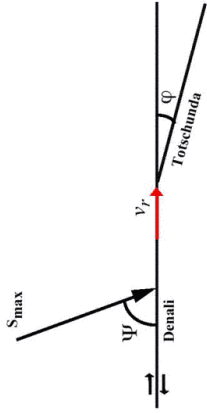
Fault map from Department of Natural Resources, Division of Geological and Geophysical Surveys, Alaska

Branching parameters for Denali-Totschunda

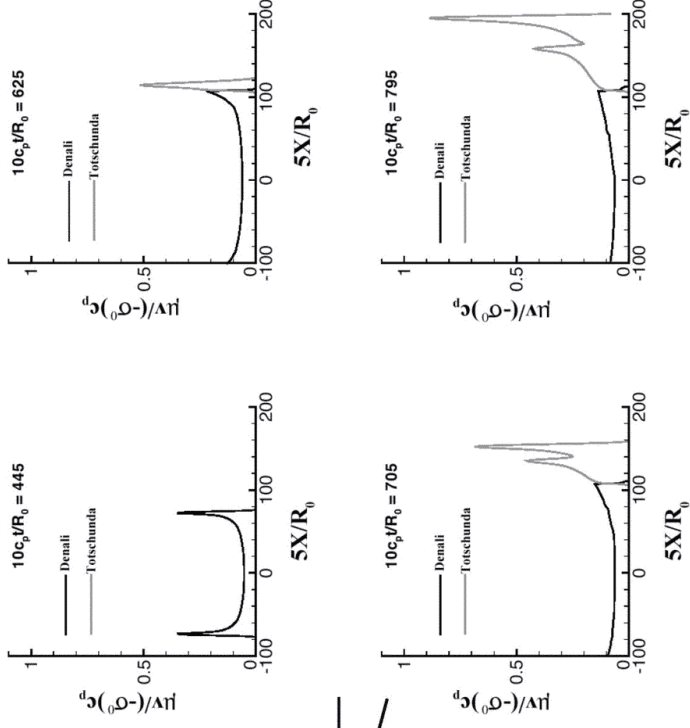
- Orientation of the principal maximum stress with the main fault, Ψ
 - Ratchkovski [2003]: $\Psi \approx 80^\circ$
 - Ratchkovski and Hansen [2002]: $\Psi \approx 73^\circ$
 - Estabrook et al. [1988] and Nakamura et al. [1980]: $\Psi \approx 75^\circ$
- Rupture velocity near the branching region, v_r : not well constrained
 - Kikuchi and Yamanaka [2002]: Average $v_r = 0.8c_s$
 - Ellsworth et al. [2004]: $v_r > c_s$ near PS10. $v_r = 0.8c_s$ beyond PS10
- Branching angle, φ
 - Savage and Lisowski [1991]: $\varphi \approx -15^\circ$ to the extensional side

(Bhat et al., 2002, 2004)

Nondimensional slip velocity v
vs. distance X along fault traces

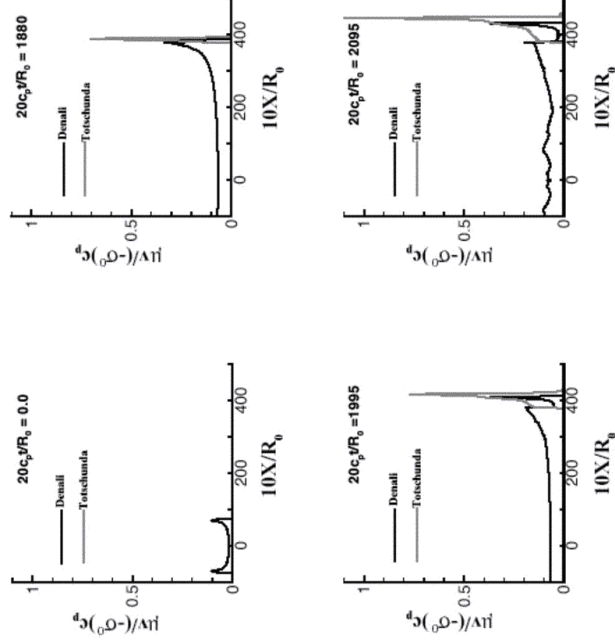


$\Psi=70^\circ, v_r = 0.8c_s, \phi = -15^\circ$



[For $v_r = 0.6 c_s$;
similar results]

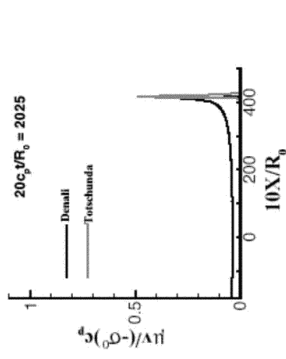
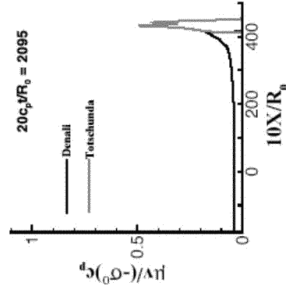
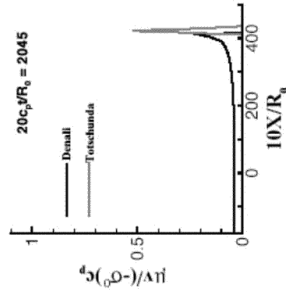
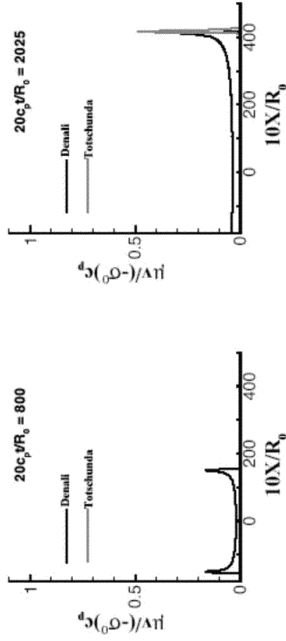
$\Psi=70^\circ, v_r = 0.9c_s, \phi = -15^\circ$



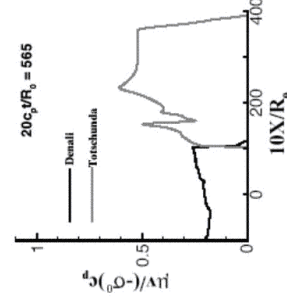
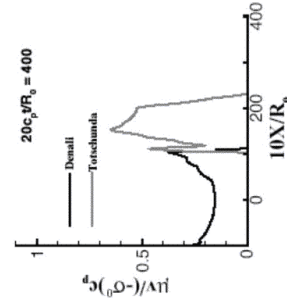
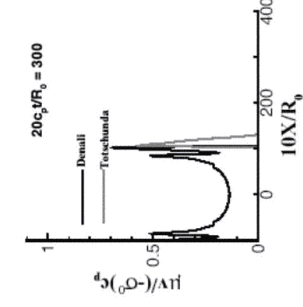
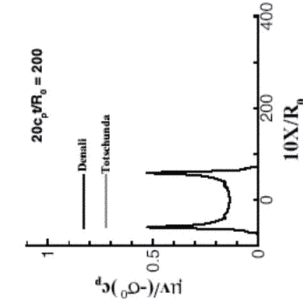
For $v_r = 0.87 c_s$:

both faults
rupture -- but
only Totschunda
ruptures if $\Psi = 80^\circ$
(next panel).

$\Psi=80^\circ, \nu_r = 0.87c_s, \phi = -15^\circ$

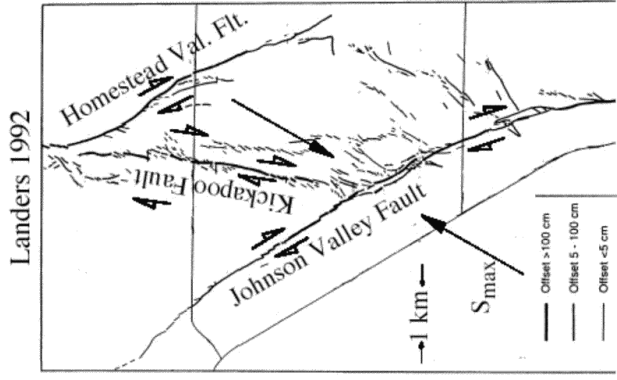


$\Psi=70^\circ, \nu_r = 1.4c_s, \phi = -15^\circ$

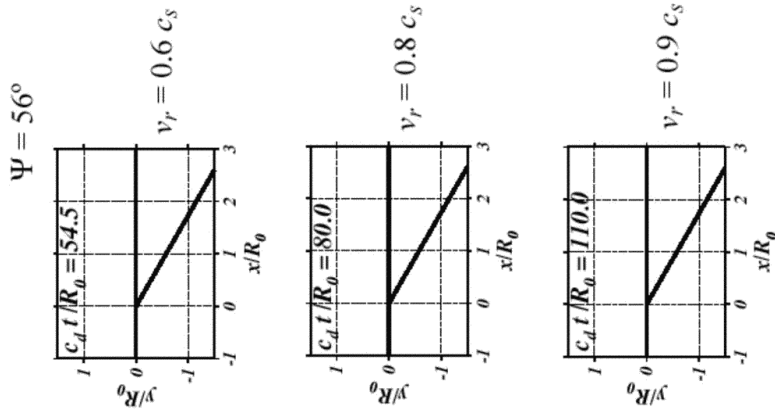


Interseismic
incident
rupture speed

(Kame et al., *JGR*, 2003)

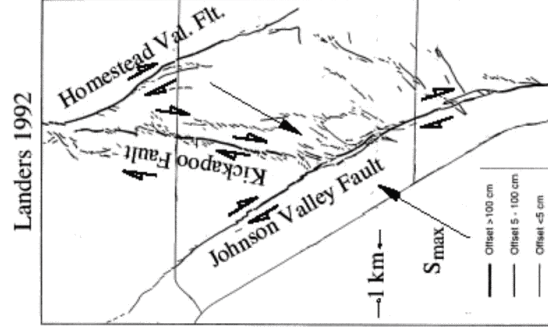
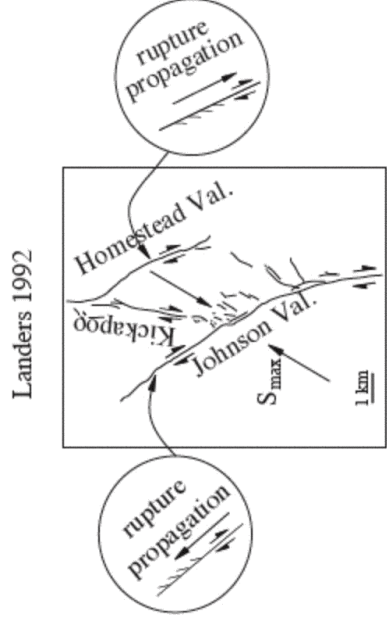


(fault map from Sowers et al., 1994; stress from Hardebeck and Hauksson, 2001)



Small branching features show the direction of propagation
[Poliakov, Dmowska & Rice, *JGR*, 2002]

Map view: Steep S_{max} direction, $\psi \approx 60^\circ$;
secondary failures on *extensional* side:



(fault map from Sowers et al., 1994; stress from Hardebeck and Hauksson, 2001)

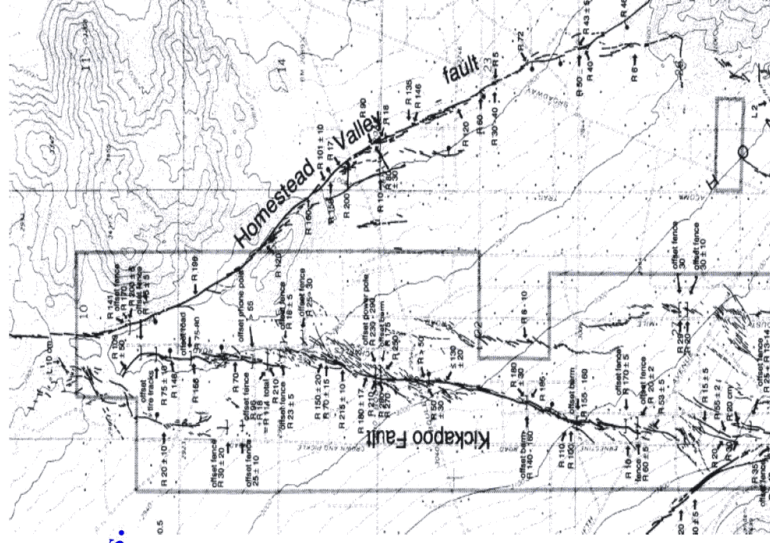
There are lots of *minor branch faults*.

How do they interact with rupture on the major fault strands?

Could they make rupture propagation very nonuniform? -- and be a source of high-frequency ground motion?

Could they arrest rupture on a major fault strand?

Examples given by Olives, Bhat, Rice & Dmowska [EOS 2004]



(Dmowska et al., EOS, 2002; Fliss et al., JGR. 2005)

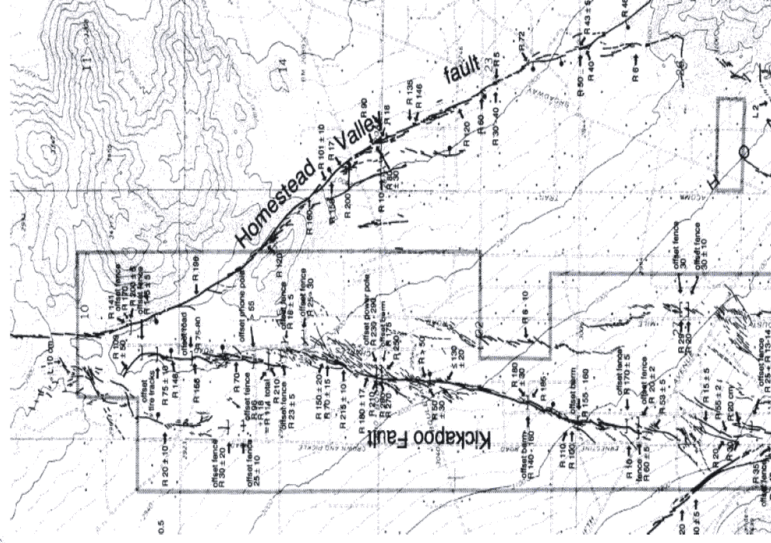
An example of *backward branching*:

Landers 1992 Earthquake

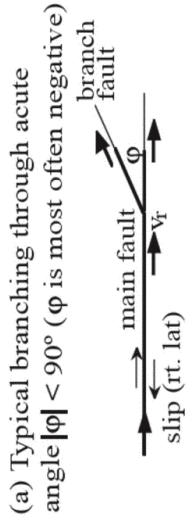
Rupture transition from Kikapoo Fault to southern part of Homestead Valley Fault (which ruptured much further to the north, off the map here):

How does *backward branching* happen?

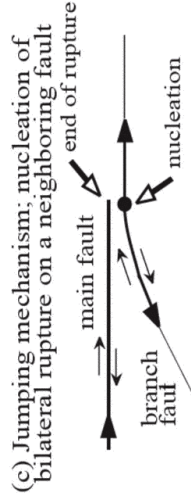
(Fault map: Sowers et al., 1994.)



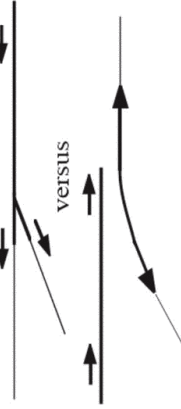
Backward Branching and Rupture Directivity



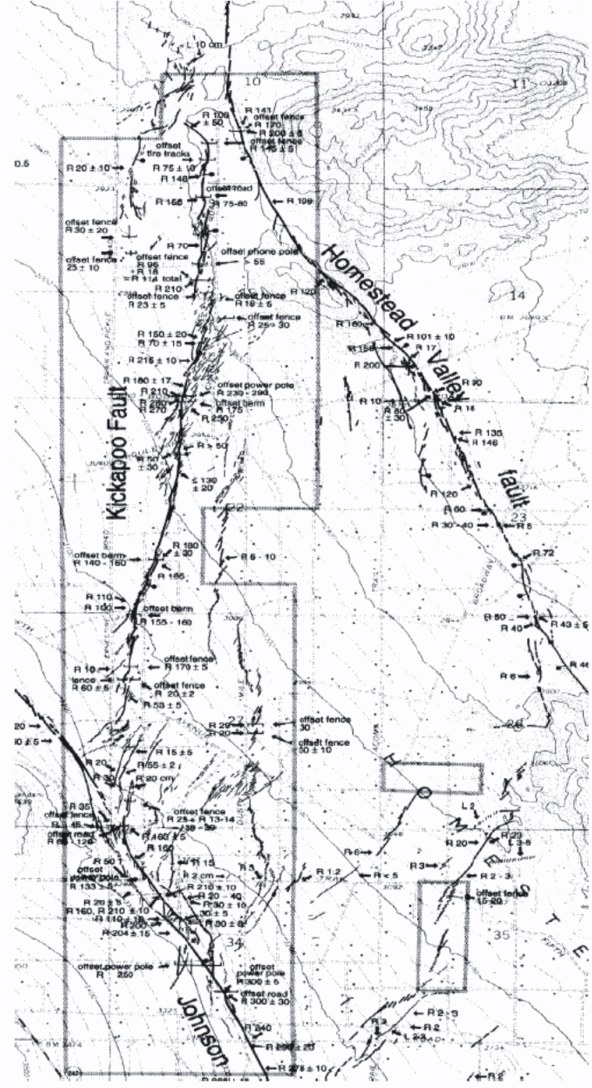
(b) Turn of rupture path through obtuse angle while continuing on main fault (never favored by stress field)



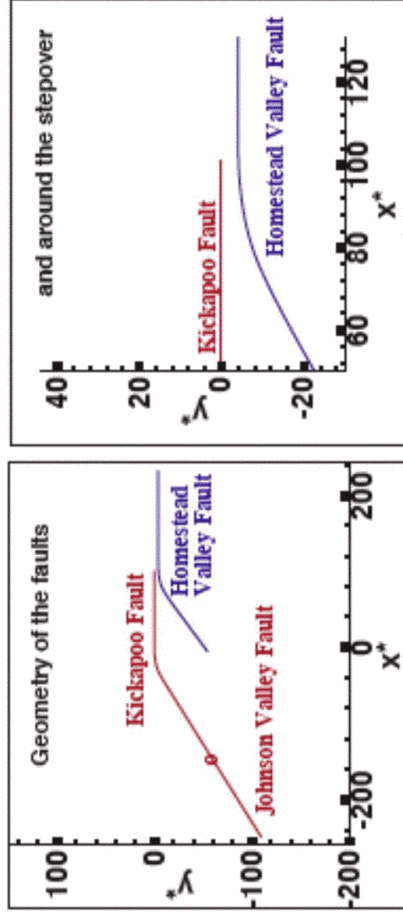
(d) Given the branch geometry, what was the direction of rupture propagation?



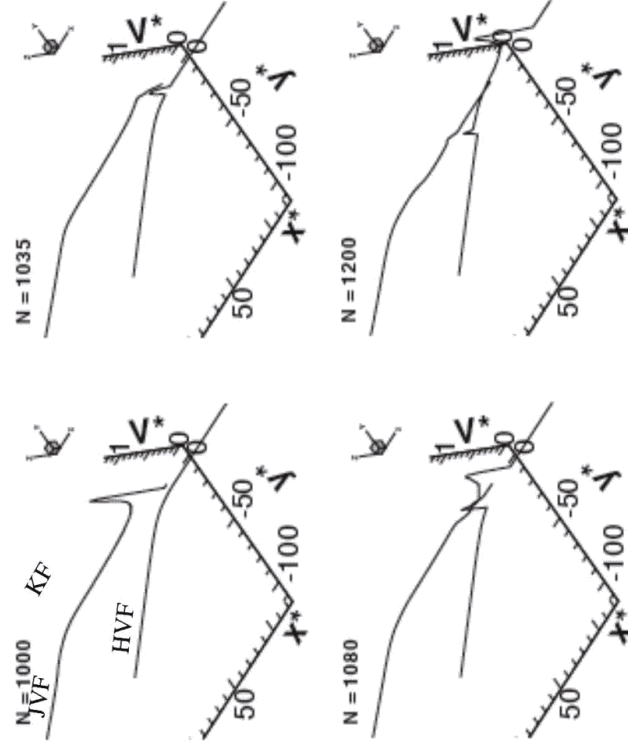
North \rightarrow



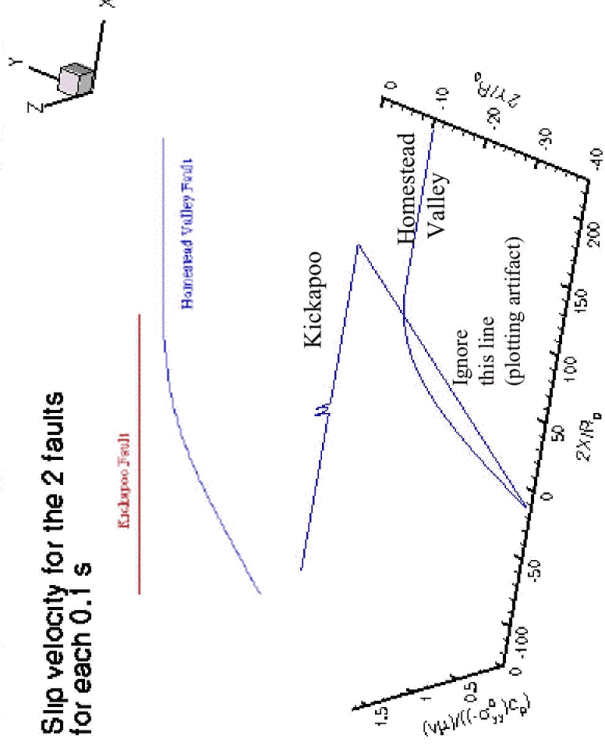
Fault geometry assumed for simulation



Fliss, Bhat, Dmowska & Rice [JGR, 2005]: simulation showing jump from the Kickapoo Fault (KF) to the Homestead Valley Fault (HVF), for a case with sub-Rayleigh v_r on the KF

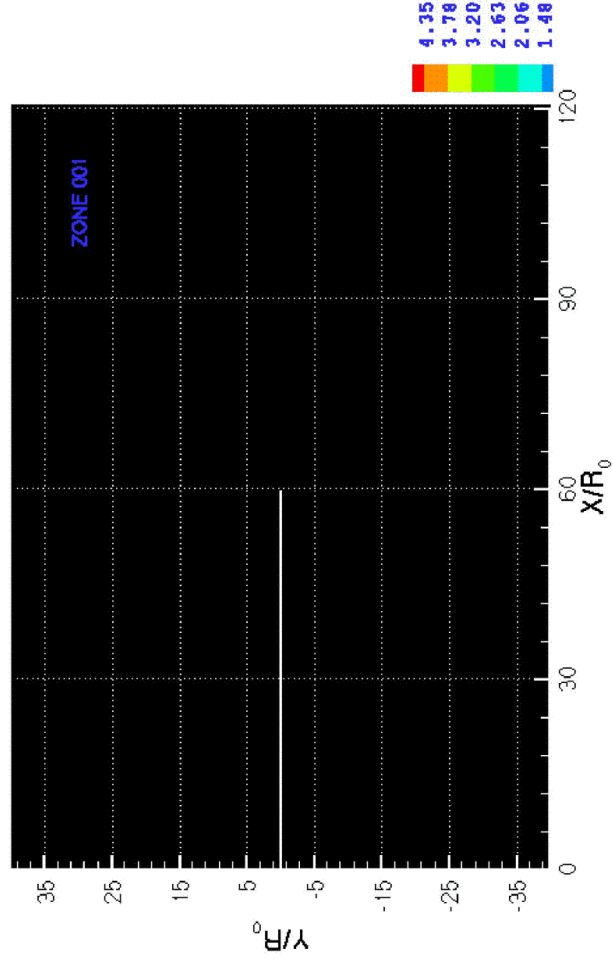


(Fliss et al. [EOS, 2003] version with supershear on KF; related sub-Rayleigh case [JGR, 2005])

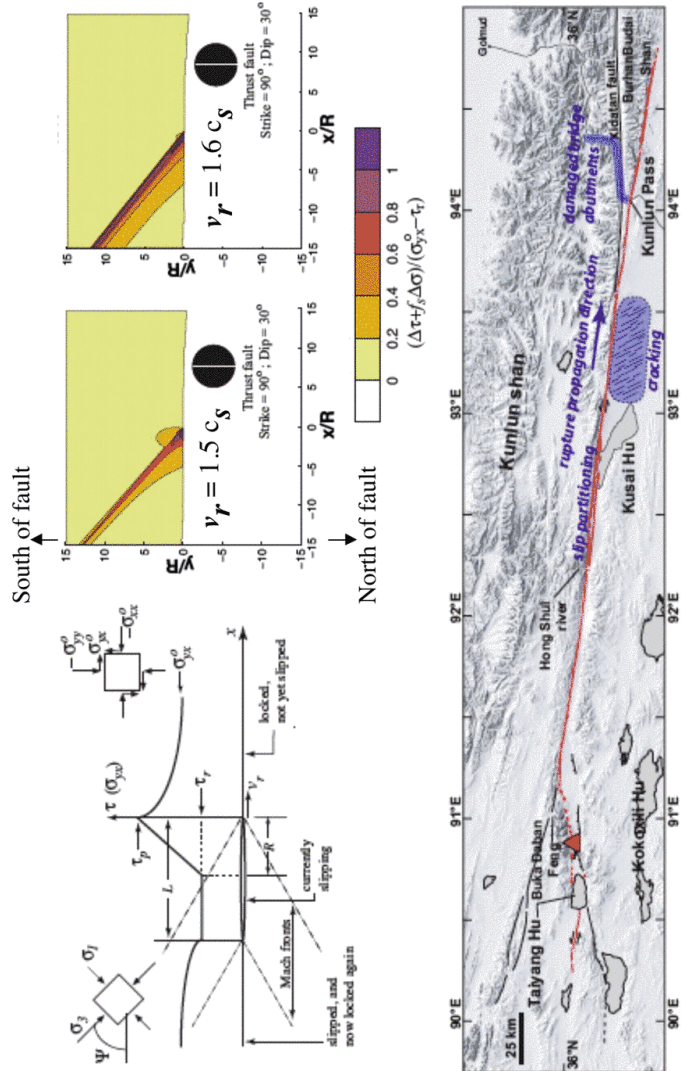


(Fliss, Bhat, et al., 2003)

Contour plot of $T_{\text{coll}} = \tau_{xy} / \mu_s (-\sigma_{yy})$



Off-fault stressing associated with supershear rupture, $c_s < v_r < c_p$, and damage in 2001 $M_w=8.1$ Kokoxili (Kunlun), Tibet, event [Bhat, Dmowska, King, Klinger & Rice, in prep. 2005]



Summary

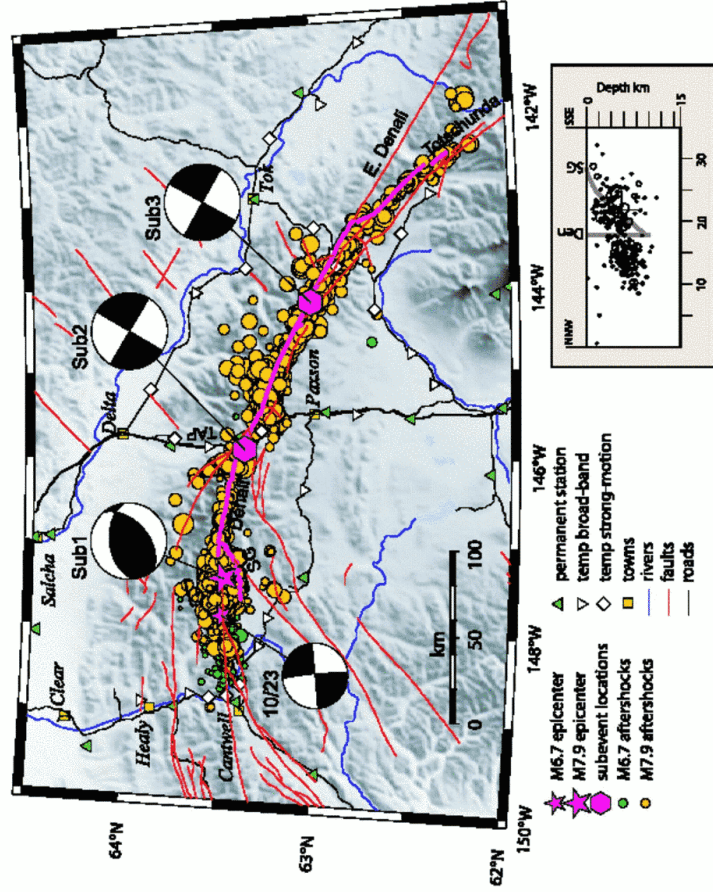
- Strongly driven rupture on a planar fault accelerates towards its limiting speed, c_{Limit} ($= c_R$ for mode II, c_s for mode III).
As $v_r \rightarrow c_{Limit}$, stresses *off* the main fault plane become much larger than *on* it. That nucleates failure along favorably oriented branches near the rupture front.
- Whether such failure, once nucleated, can continue to larger scales depends on the *pre-stress* state.
For mode II rupture, S_{max} at a *shallow* angle Ψ to the fault (e.g., $\Psi < 20^\circ$) favors rupture to the *compressional* side; S_{max} at *steeper* angle $\Psi > 45^\circ$) favors *extensional* side.
Concepts are generally supported by field observations and numerical simulations.
- Simulations also support more abundant branching at higher rupture velocities.
- Forward branching and off-fault damage in a given stress field is clearly related to rupture directivity.
- Backward branching most likely achieved as abrupt arrest on primary fault, followed by jump to a neighboring fault and bilateral propagation on it.
Such mechanism makes diagnosing directivity of a past earthquake difficult without detailed knowledge of the branching process.

Papers and download links:

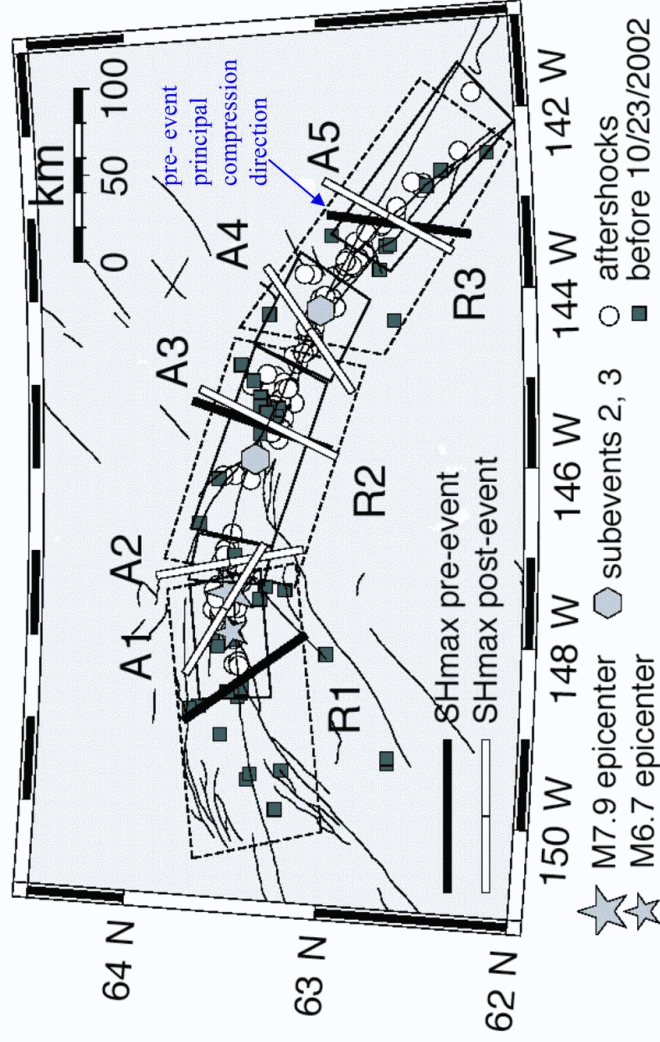
- A. N. B. Poliakov, R. Dmowska and J. R. Rice, 2002: Dynamic shear rupture interactions with fault bends and off-axis secondary faulting. *Journal of Geophysical Research*, **107** (B11), *cn*: 2295, *doi*:10.1029/2001JB000572, pp. ESE 6-1 to 6-18
http://esag.harvard.edu/dmowska/PoliakovDmowskaRice_JGR02.pdf
- N. Kame, J. R. Rice and R. Dmowska, 2003: Effects of pre-stress state and rupture velocity on dynamic fault branching. *Journal of Geophysical Research*, **108**(B5), *cn*: 2265, *doi*: 10.1029/2002JB002189, pp. ESE 13-1 to 13-21.
http://esag.harvard.edu/dmowska/KameRiceDmowska_JGR03.pdf
- H. S. Bhat, R. Dmowska, J. R. Rice and N. Kame, 2004: Dynamic slip transfer from the Denali to Totschunda Faults, Alaska: Testing theory for fault branching. *Bulletin of the Seismological Society of America*, **94**(6B), pp. S202-S213.
http://esag.harvard.edu/dmowska/BhatDmRiKa_Denali_BSSA04.pdf
- S. Fliss, H. S. Bhat, R. Dmowska and J. R. Rice, 2005: Fault branching and rupture directivity. *Journal of Geophysical Research*, **110**, B06312, *doi*:10.1029/2004JB003368, 22 pages.
http://esag.harvard.edu/dmowska/FlissBhatDmRi_JGR05.pdf
- H. S. Bhat, R. Dmowska, G. C. P. King, Y. Klinger and J. R. Rice, 2005: Off-fault damage patterns due to supershear ruptures with application to the 2001 Mw 8.1 Kokoxili (Kunlun) Tibet earthquake. *semi-final draft, to be submitted to Journal of Geophysical Research*.
http://esag.harvard.edu/dmowska/BhatDmKIKIRI_supershear_drft03Aug05.pdf

Additional materials, in case needed.

M_w = 7.9 Denali Earthquake, 3 Nov 2002 (Eberhart-Phillips *et al.*, *Science*, 2003)



Pre-stress directions (black lines) along path of $M_w = 7.9$ Denali Earthquake, 3 Nov 2002 (Ratchkovski, *Geophys. Res. Lett.*, 2003); shows $\Psi \approx 70^\circ$ to 80° near Totschunda branch



Boundary Integral Equation formulation, based on Kame and Yamashita (*GJI*, 1999), with Coulomb slip-weakening

$$u_k(x, t) = \int_{-\infty}^t \int_{\text{Fault}} C_{pqrs} \Delta v_r(x', t') v_s(x') \frac{\partial G_{kp}(x - x', t - t')}{\partial x'_q} ds' dt'$$

Heaviside Green's function

stress $\sigma_{ij}(x, t) = C_{ijkl} \partial u_k(x, t) / \partial x_l$

slip velocity

crack surface

stress evaluation point for the fracture criterion

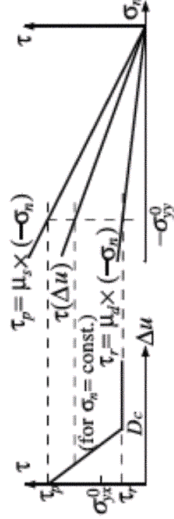
discrete elements for slip velocity

stress evaluation points for the B.C.

Stress can be evaluated anywhere.

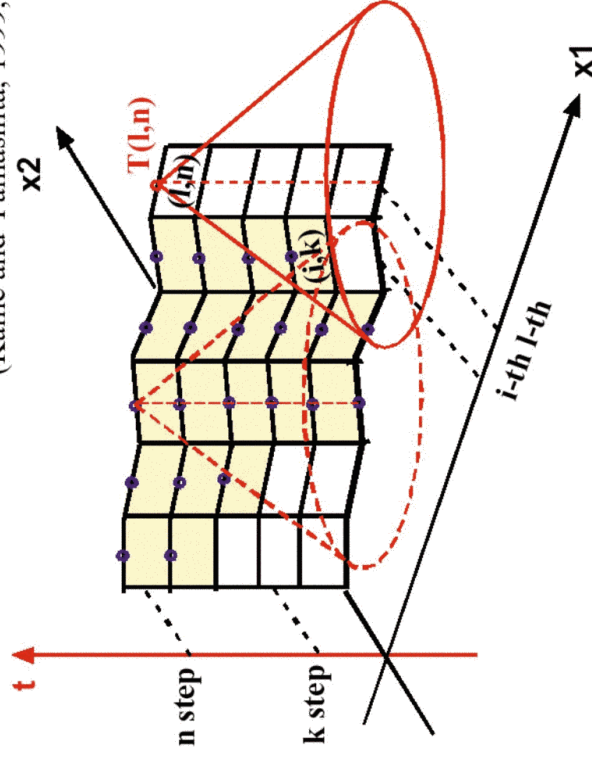
Slip-weakening model of frictional failure:

τ = shear stress,
 σ_n = normal stress,
 Δu = fault slip.



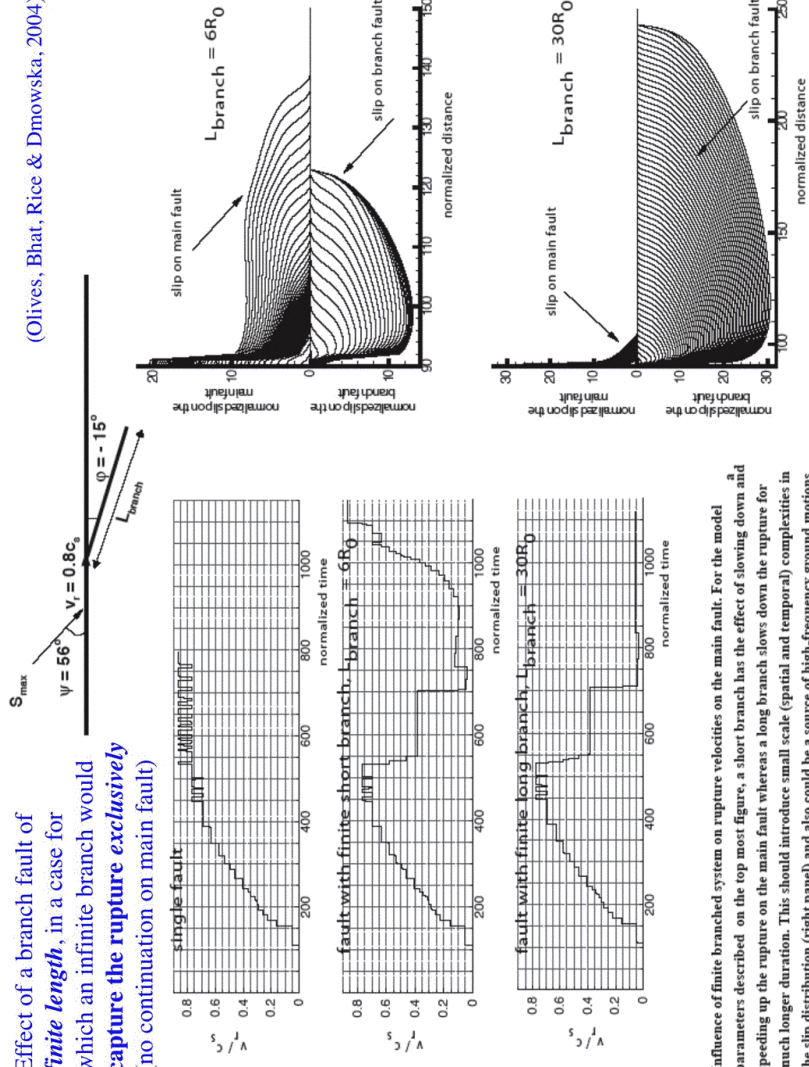
Boundary Integral Equation Method

(Kame and Yamashita, 1999, *GJI*)



Effect of a branch fault of *finite length*, in a case for which an infinite branch would capture the rupture *exclusively* (no continuation on main fault)

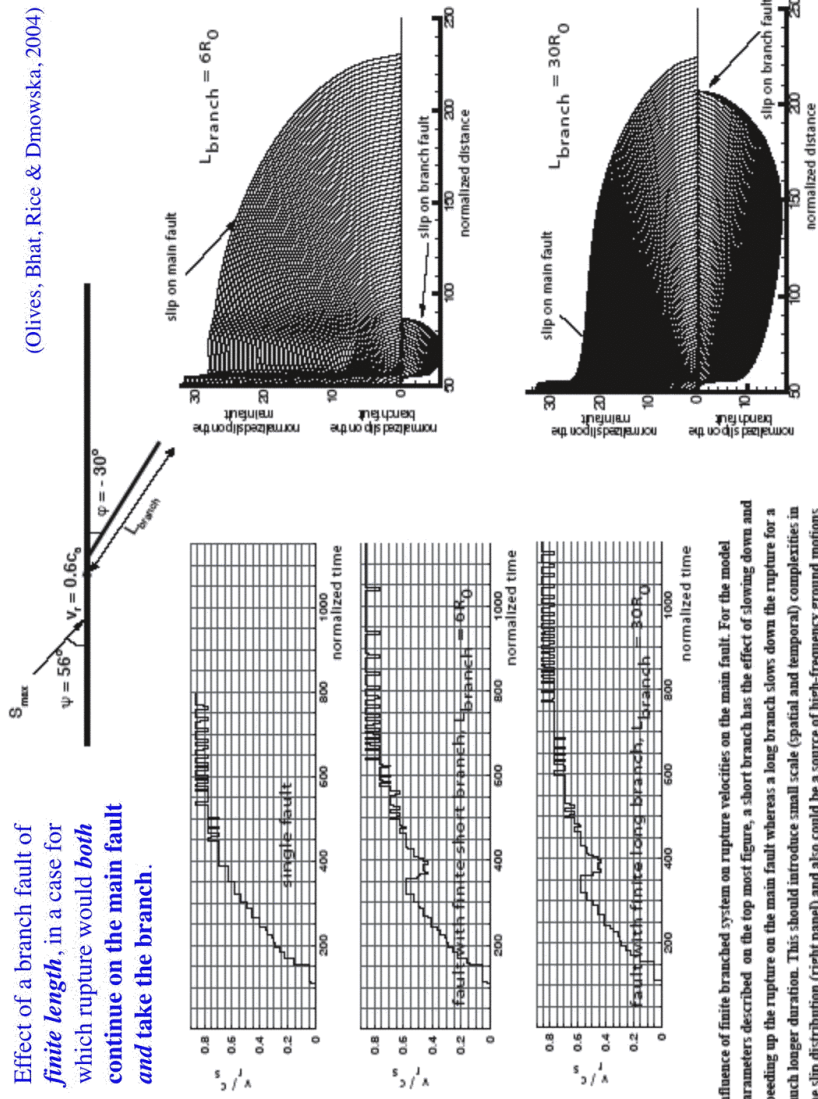
(Olives, Bhat, Rice & Dmowska, 2004)



Influence of finite branched system on rupture velocities on the main fault. For the model parameters described on the top most figure, a short branch has the effect of slowing down and speeding up the rupture on the main fault whereas a long branch slows down the rupture for much longer duration. This should introduce small scale (spatial and temporal) complexities in the slip distribution (right panel) and also could be a source of high-frequency ground motions [Olives et al. 2004].

Effect of a branch fault of *finite length*, in a case for which rupture would *both* continue on the main fault *and* take the branch.

(Olives, Bhat, Rice & Dmowska, 2004)



Influence of finite branched system on rupture velocities on the main fault. For the model parameters described on the top most figure, a short branch has the effect of slowing down and speeding up the rupture on the main fault whereas a long branch slows down the rupture for a much longer duration. This should introduce small scale (spatial and temporal) complexities in the slip distribution (right panel) and also could be a source of high-frequency ground motions (Olives et al. 2004).