

7

- MIXING OF A MANTLE DOMAIN DURING CONVECTION

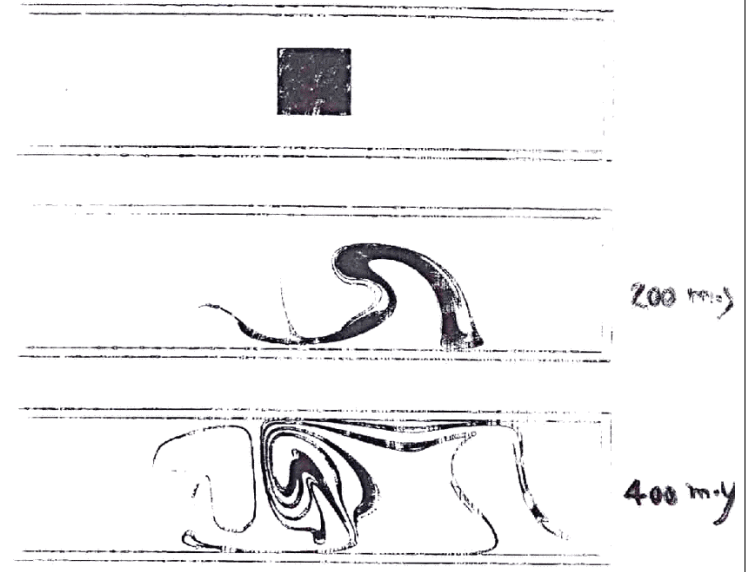
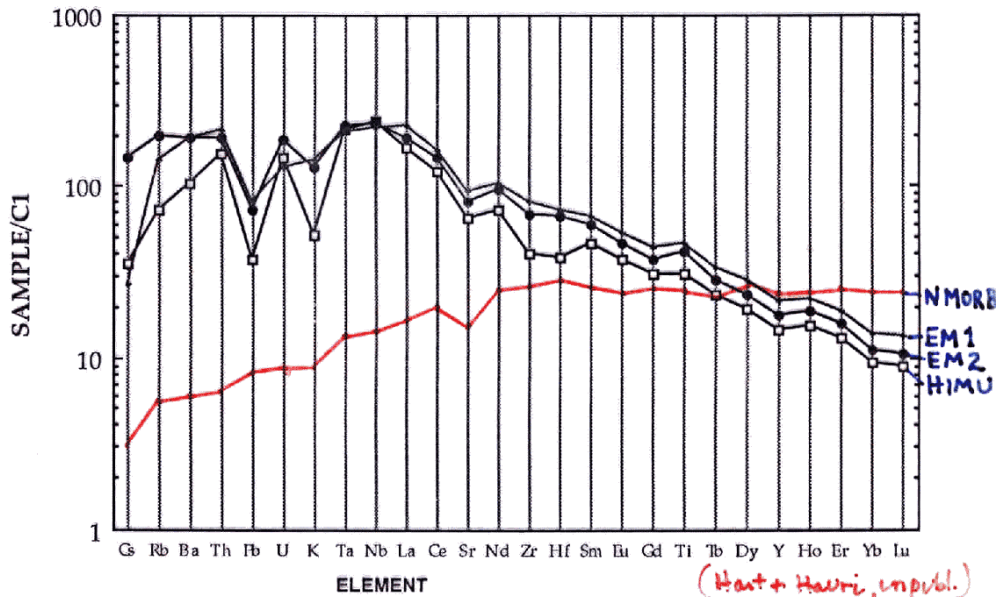


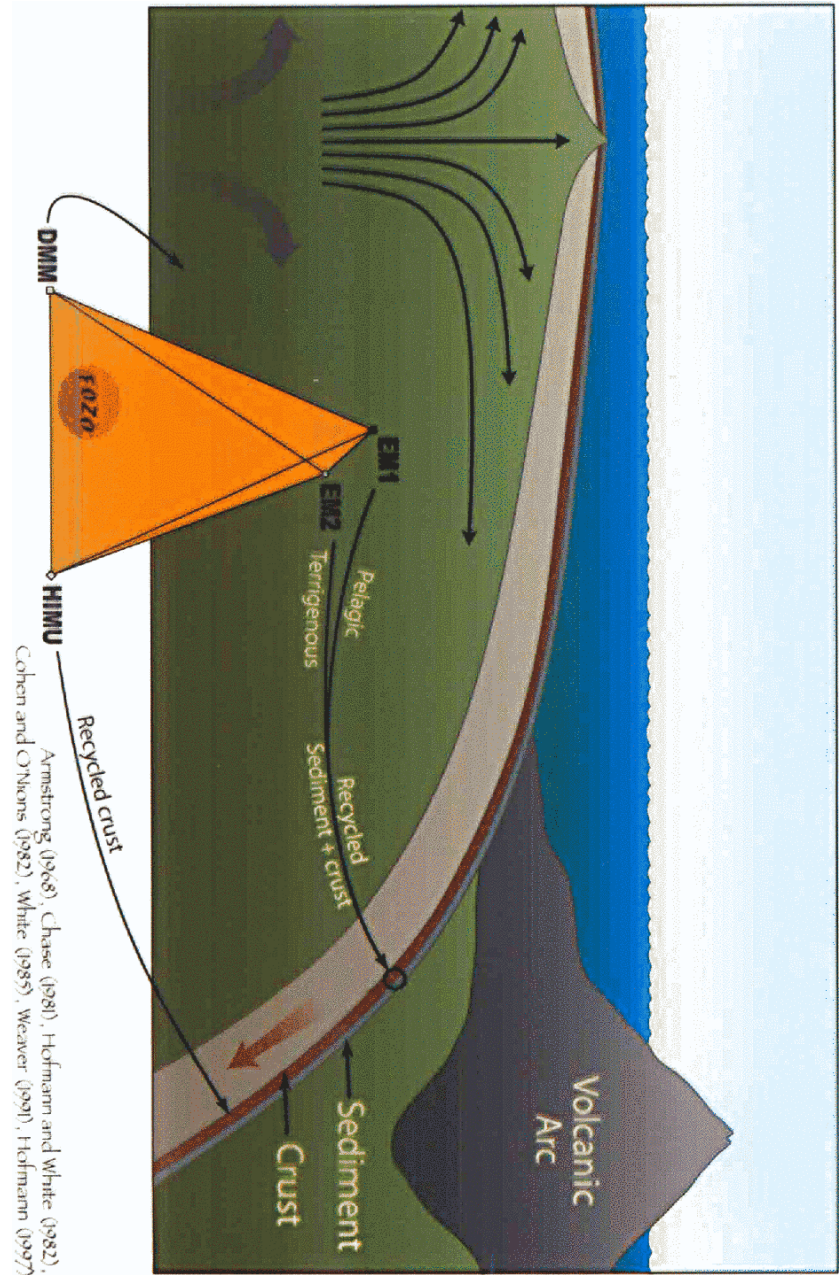
FIG. 5 - The evolution of a marked volume in the flow shown in Fig. 4. The time increment is 3×10^8 years for depth equal to 700 km. This calculation is by N. Hoffman, Cambridge University (personal communication).

(Internally heated, $R_a = 1.4 \times 10^6$)

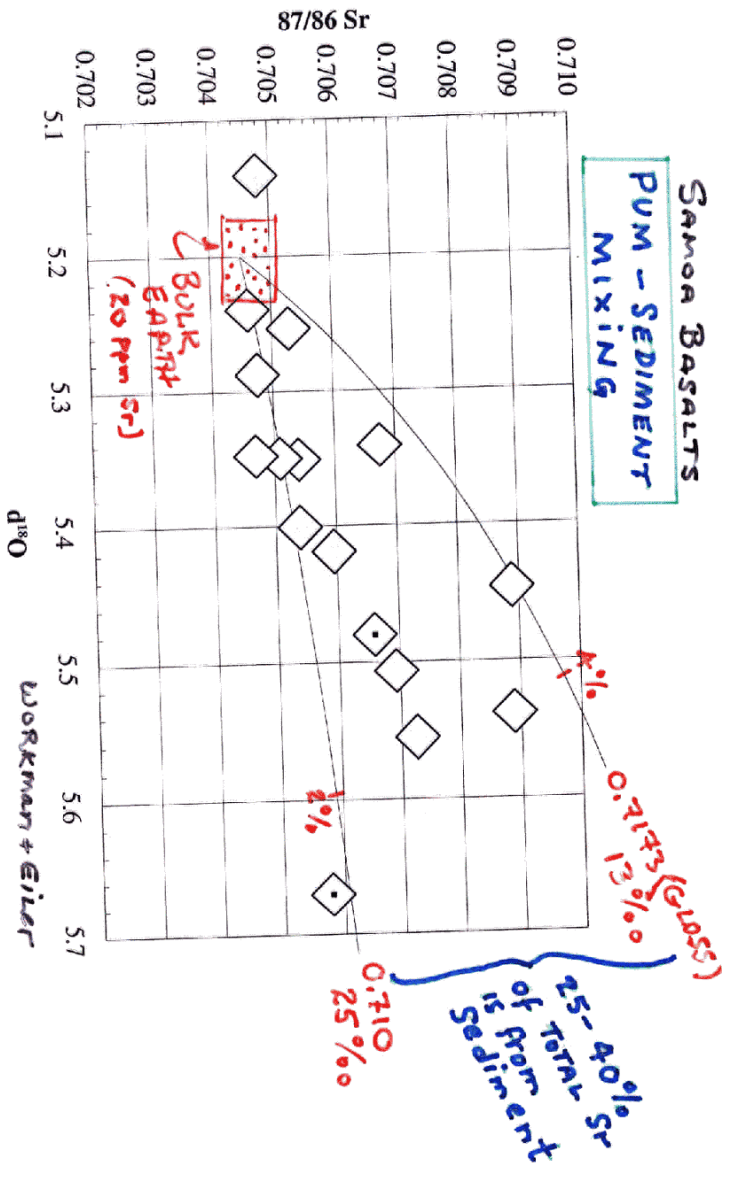
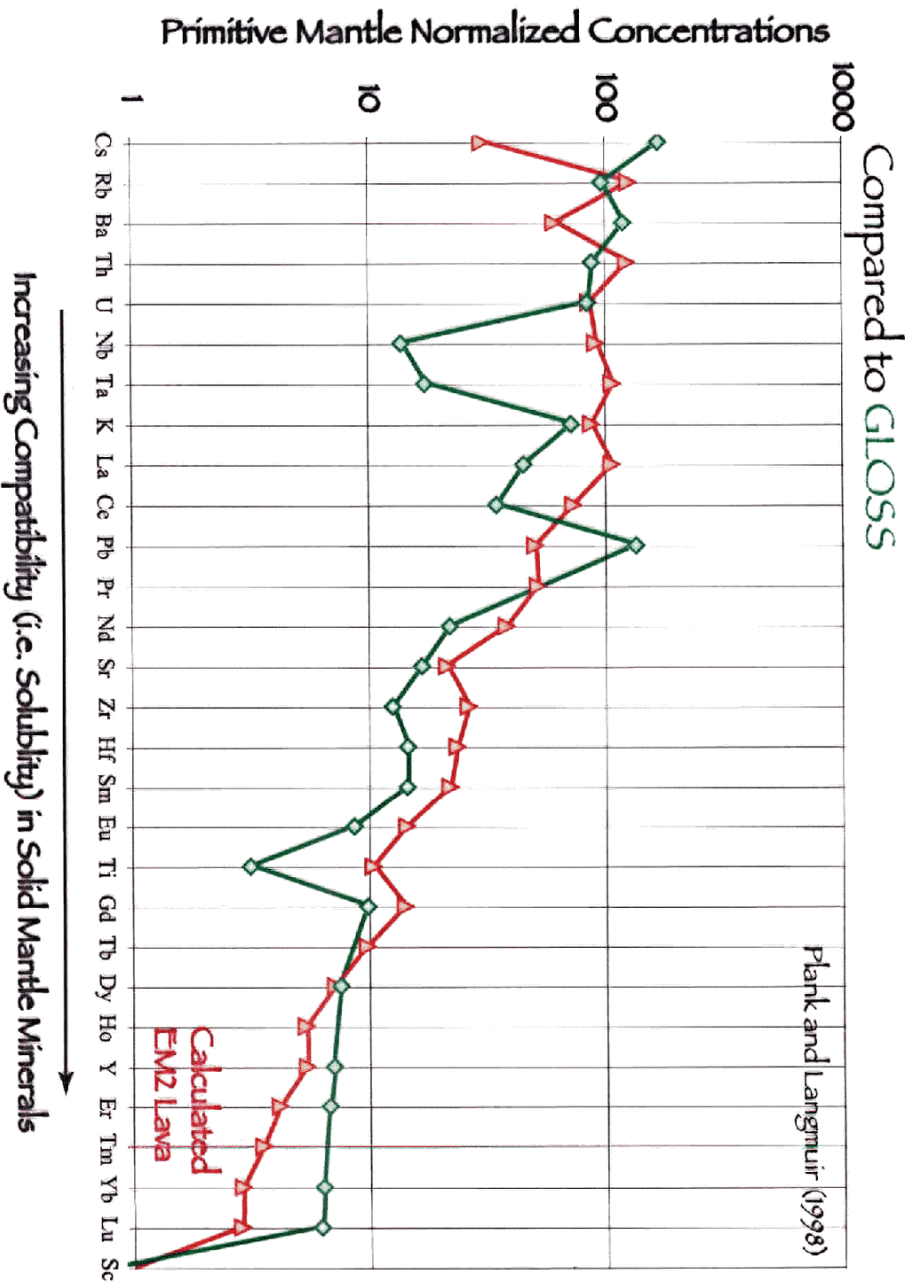
SPIDERGRAMS for MANTLE END-MEMBER BASALTS

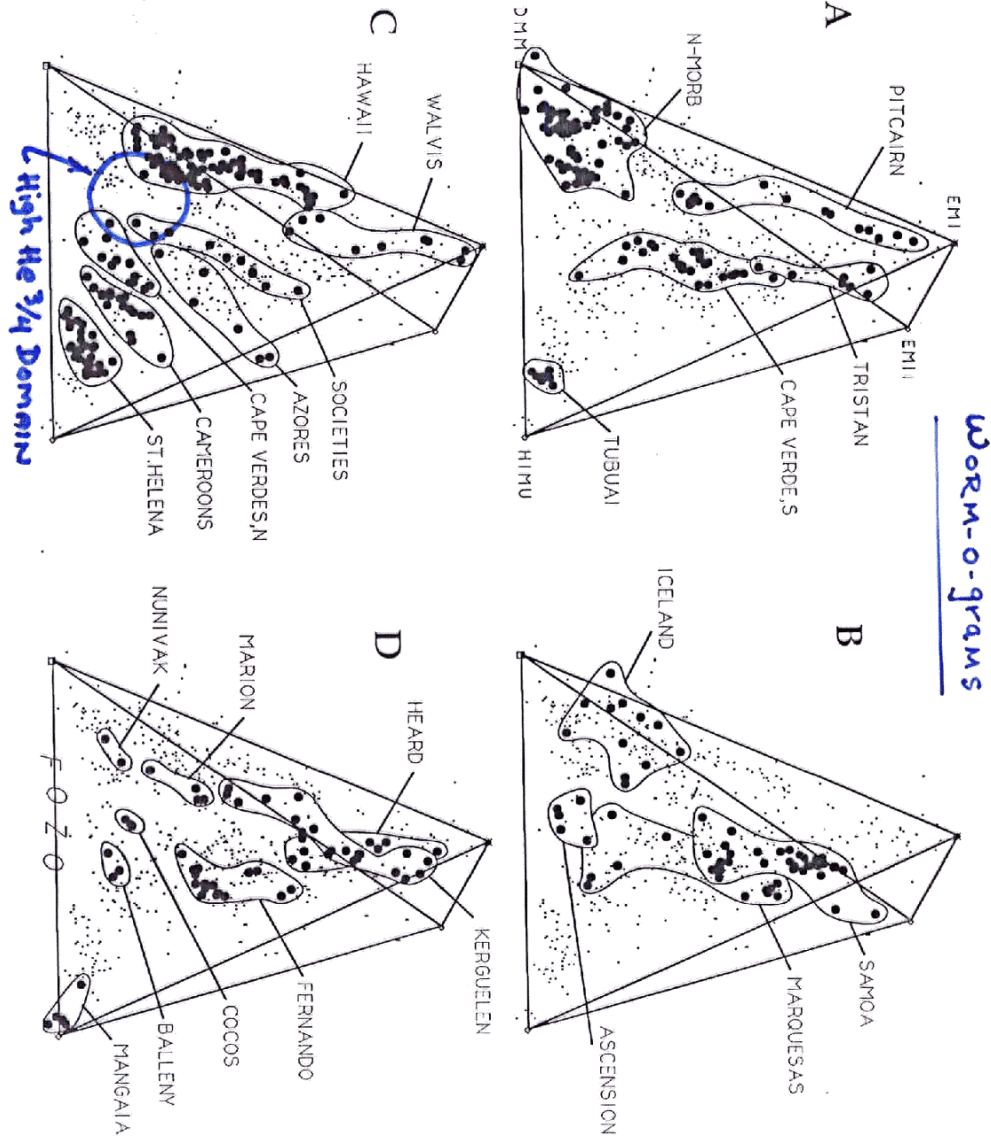


- EM1 ~ EM2 ~ HIMU ≠ N-MORB
- MINOR differences in OIB:
 - EM2 is "flat" from Cs → Th
 - HIMU has negative Zr-Hf anomaly
 - EM1 has no negative K anomaly

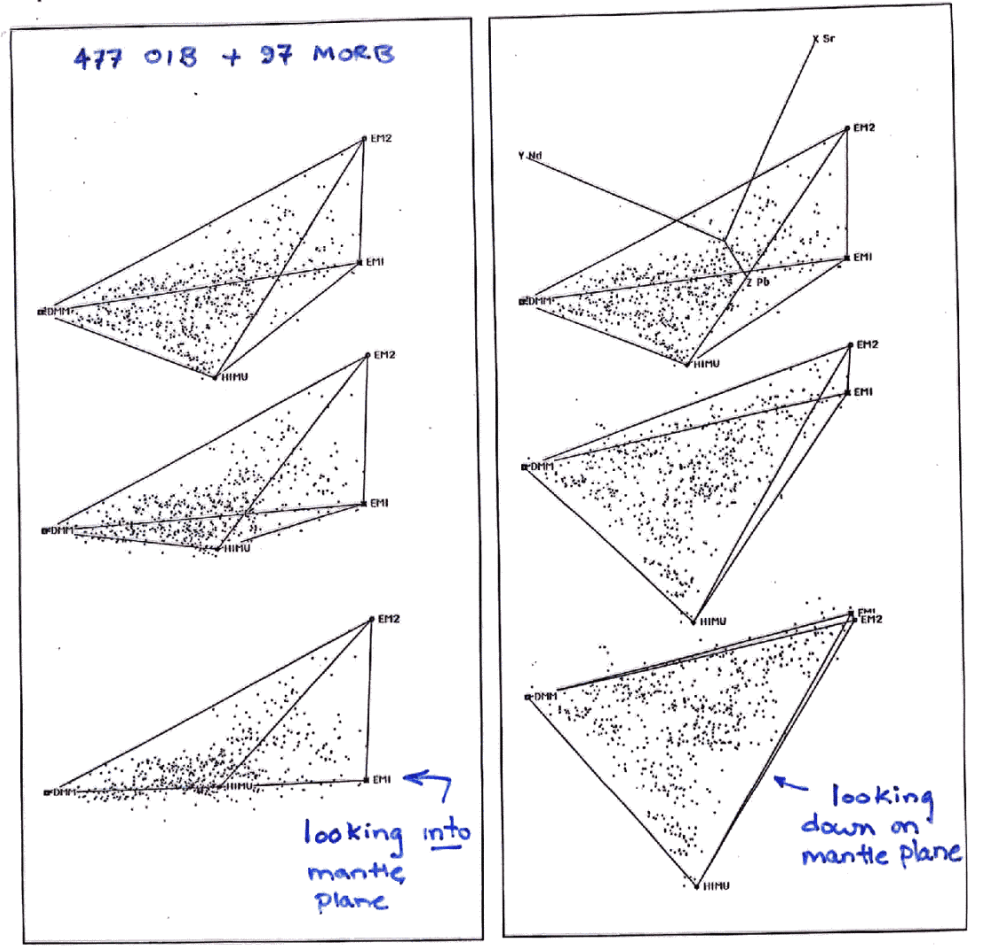


Standard Model for the Origin of Isotopic Endmembers





Worm-o-grams



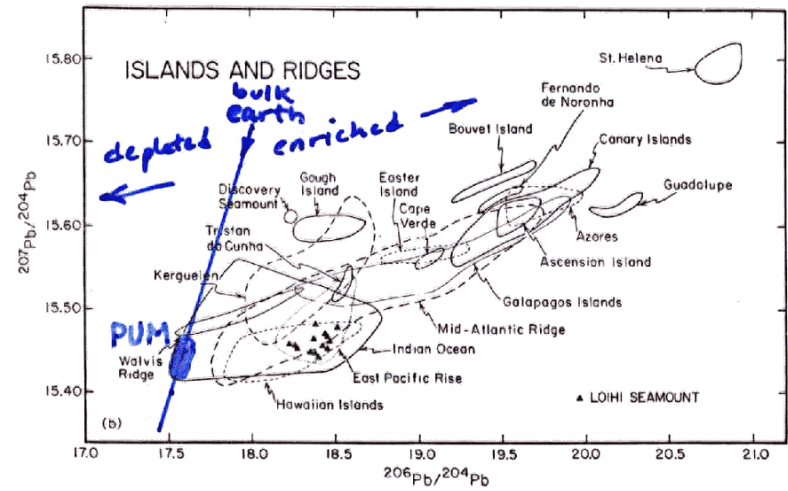
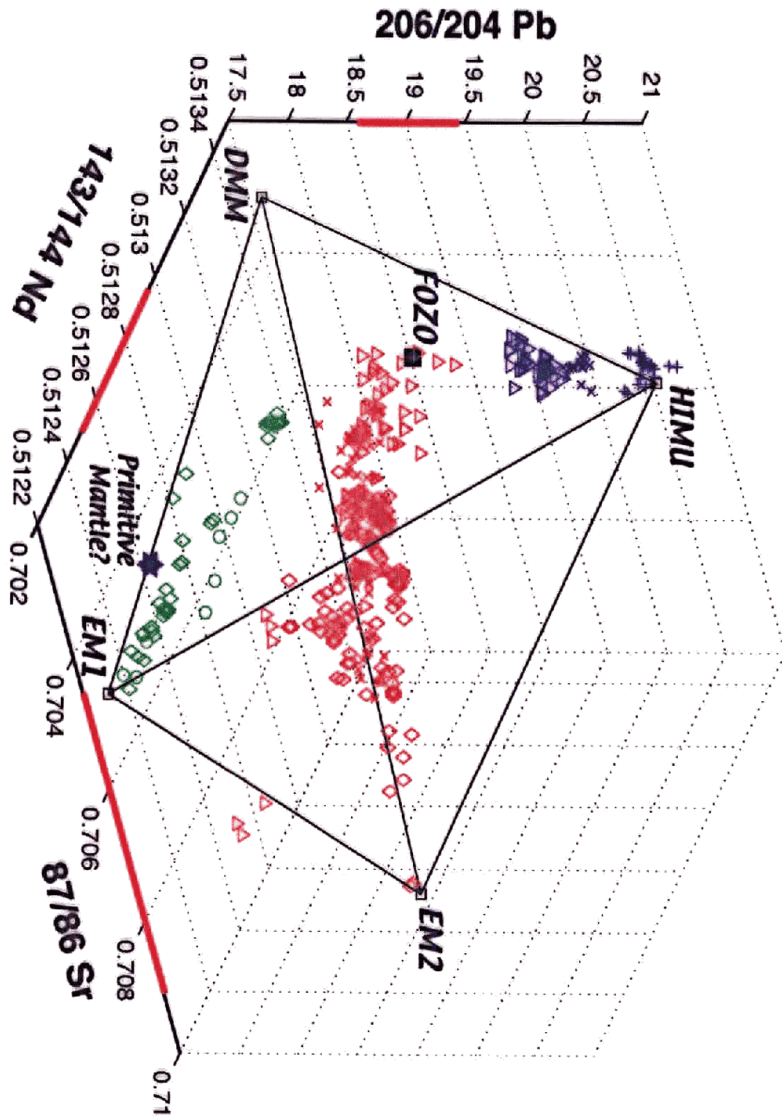
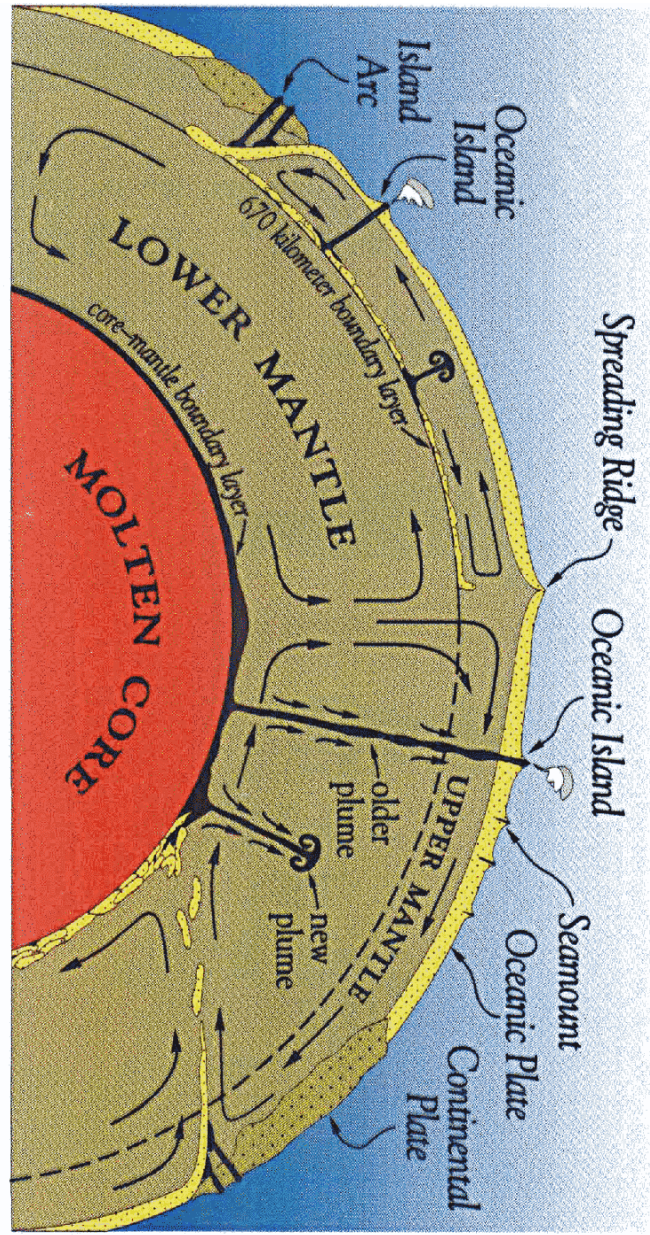
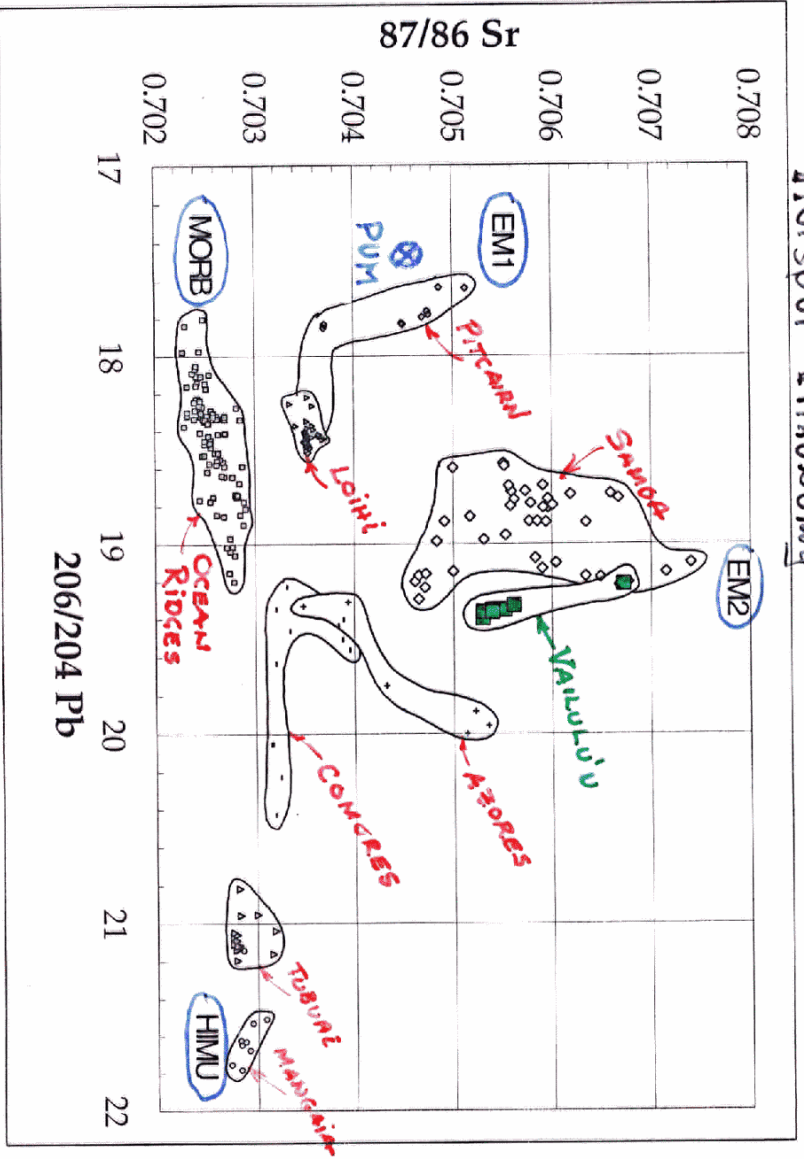


Fig. 5. (a) $^{207}\text{Pb}/^{204}\text{Pb}$ and $^{208}\text{Pb}/^{204}\text{Pb}$ versus $^{206}\text{Pb}/^{204}\text{Pb}$ for Loihi basalts compared to other Hawaiian volcanoes. (b) Loihi data are compared to published analyses for oceanic basalts. Data are from references 31,32,52; symbols as in Fig. 4. Analytical uncertainties are indicated in (a).



Horspor Taxonomy



WHAT NEXT?

- **NEED TO CONSIDER “INTERNAL” PRODUCTION OF CHEMICAL HETEROGENEITIES.**
 - melt/solid partitioning in the SLVZ (D’)
 - core/mantle reactions
 - inter-phase fractionations at the 410 and 670 km phase changes
- **NEED DYNAMICAL MODELING OF THE DESTRUCTION OF SMALL (1-10 KM) HETEROGENEITIES BY STIRRING AND MIXING.**

Mantle Heterogeneities:

- How do we know they are there?
- How are they made?
- What size are they?
- How do they survive (likely not this talk!)?
- Physical consequences –density and thermal impacts
- What next?

HELP NEEDED FROM OTHER DISCIPLINES!!

Seismologists: Do plumes come from the CMB?

Dynamicists: What will the 7 km ocean crust layer look like after a billion years of mixing/stirring/stretching in the mantle?

Seismologists: Does the oceanic asthenosphere have small degrees of melt everywhere?

Mineral physicists: What is the light element in the core, and can we also store helium, potassium, uranium and thorium there?

SENSITIVITY ANALYSIS
Chemical versus Physical Properties

	<u>Bulk Silicate Earth</u>	<u>Depleted MORB Mantle</u>	<u>EM2 Source (Harzburgitic Lithosphere)</u>
Whole Rock Mg #	89.33	89.77	90.60
Modal Olivine, %	55.2	??	76.9
Density, g/cc	3.374	3.367	3.355
Vp, km/sec	8.258	8.302	8.308
Vs, km/sec	4.807	4.814	4.826
Heat Generation, pW/kg	4.96	0.9	10.6

Delta, relative to BSE

Density	-0.19%	-0.54%
P-velocity	0.53%	0.61%
S-velocity	0.15%	0.40%
Heat Gen'n	0.18x	2.1x

- density, Vp and Vs from garnet-facies Mg# parameterization of C.-T. Lee, 2003.
- Heat generation from Workman et al., 2003.

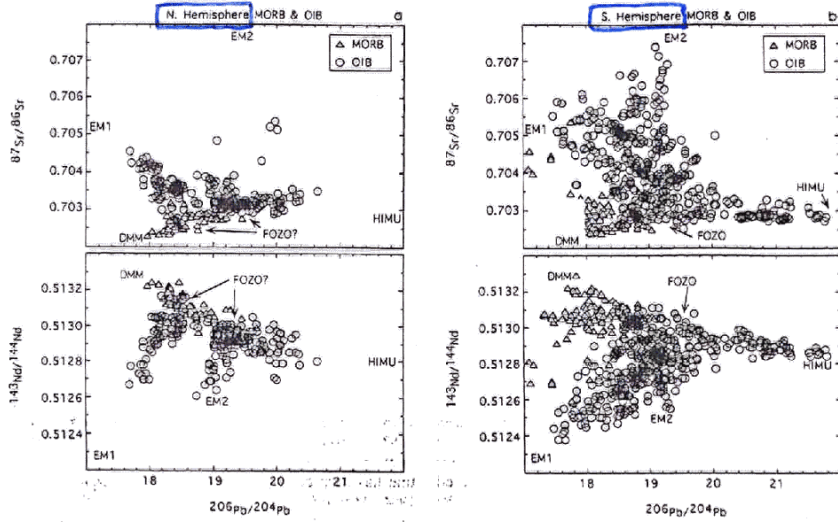
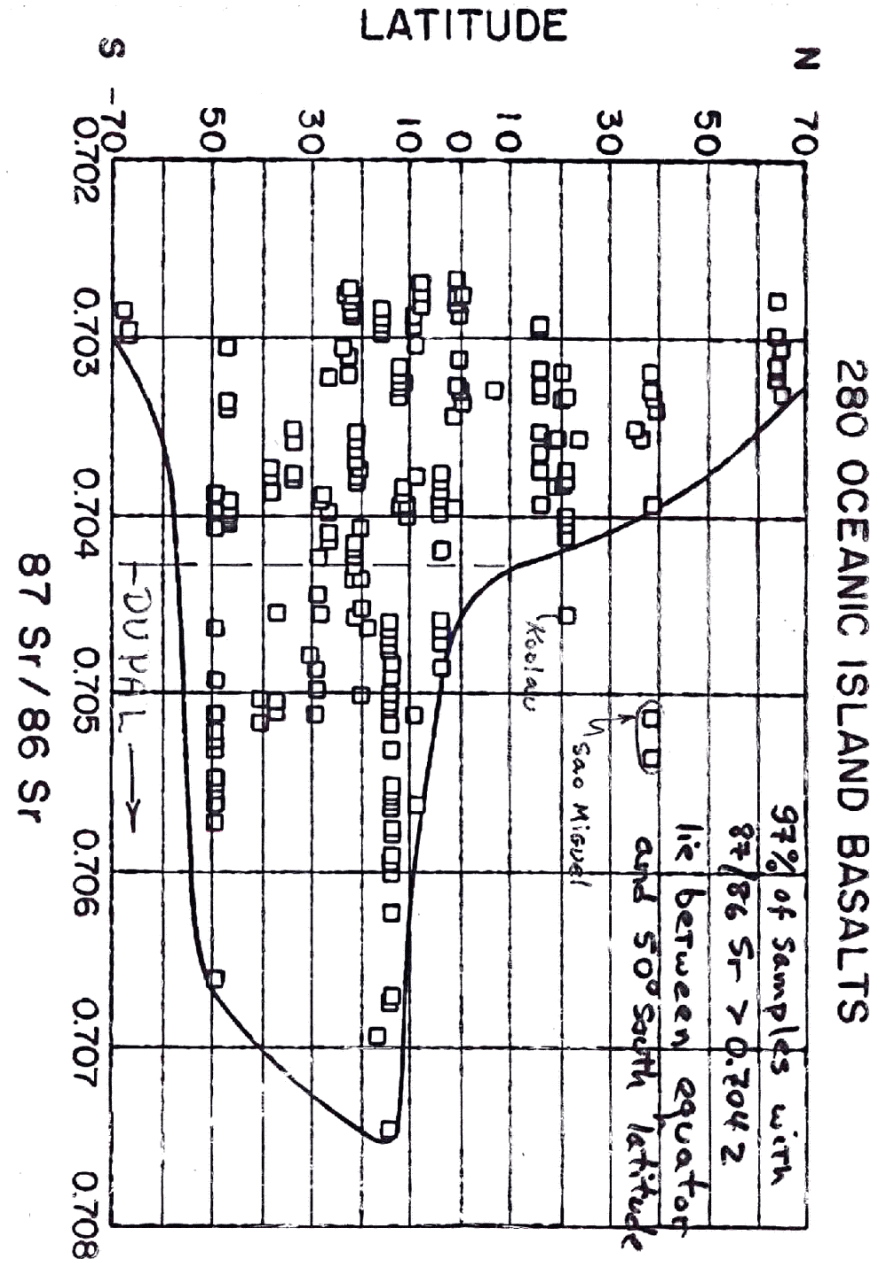
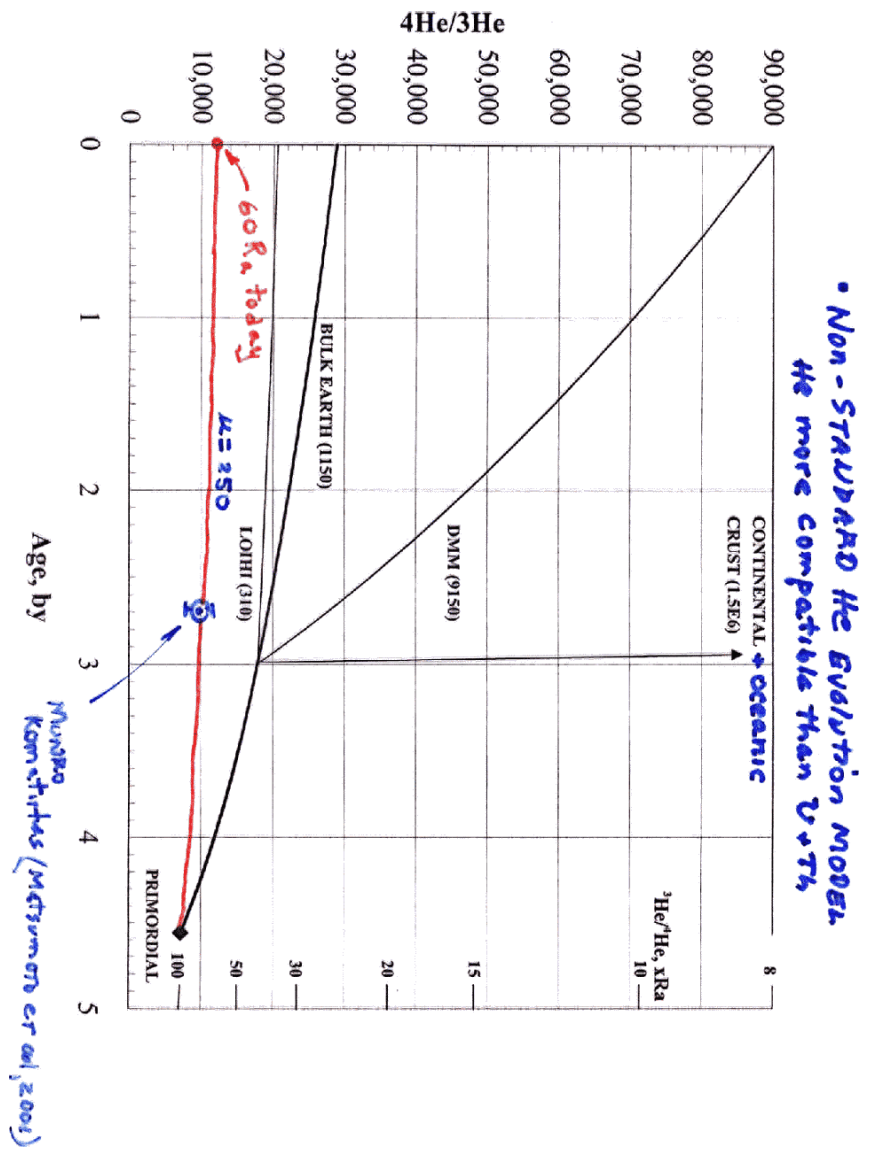
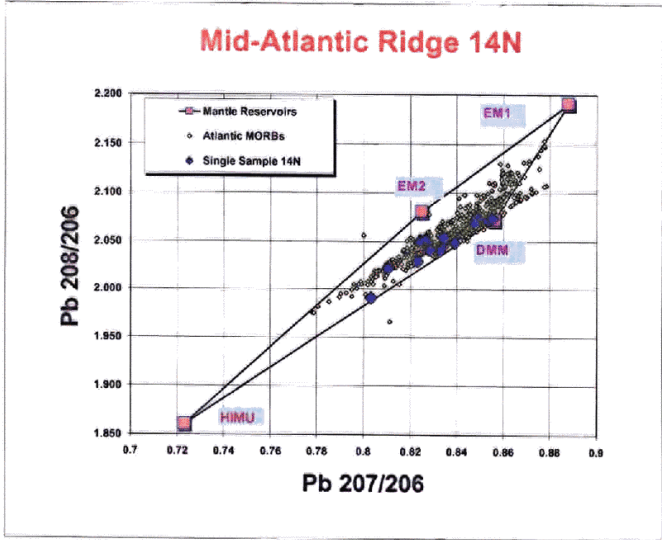
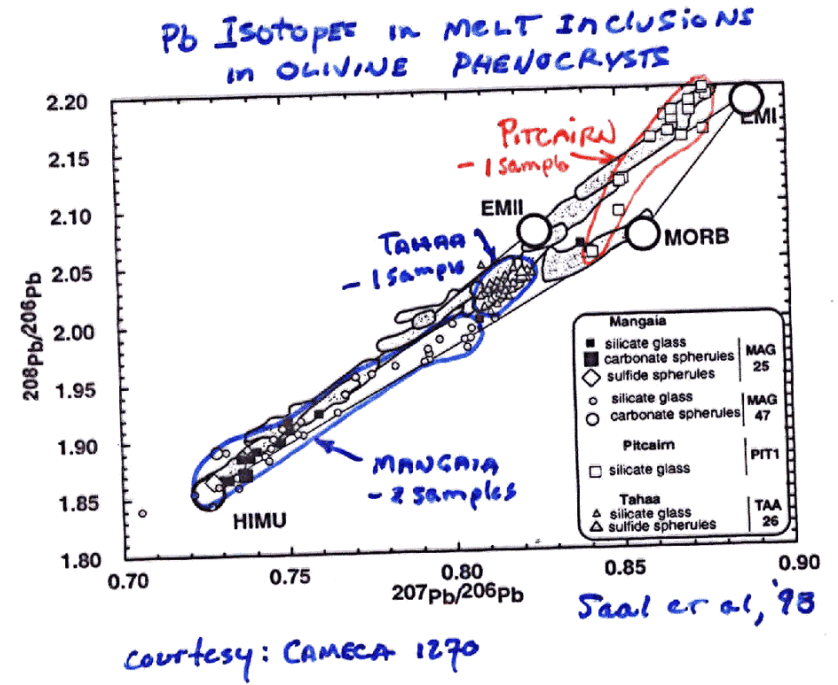
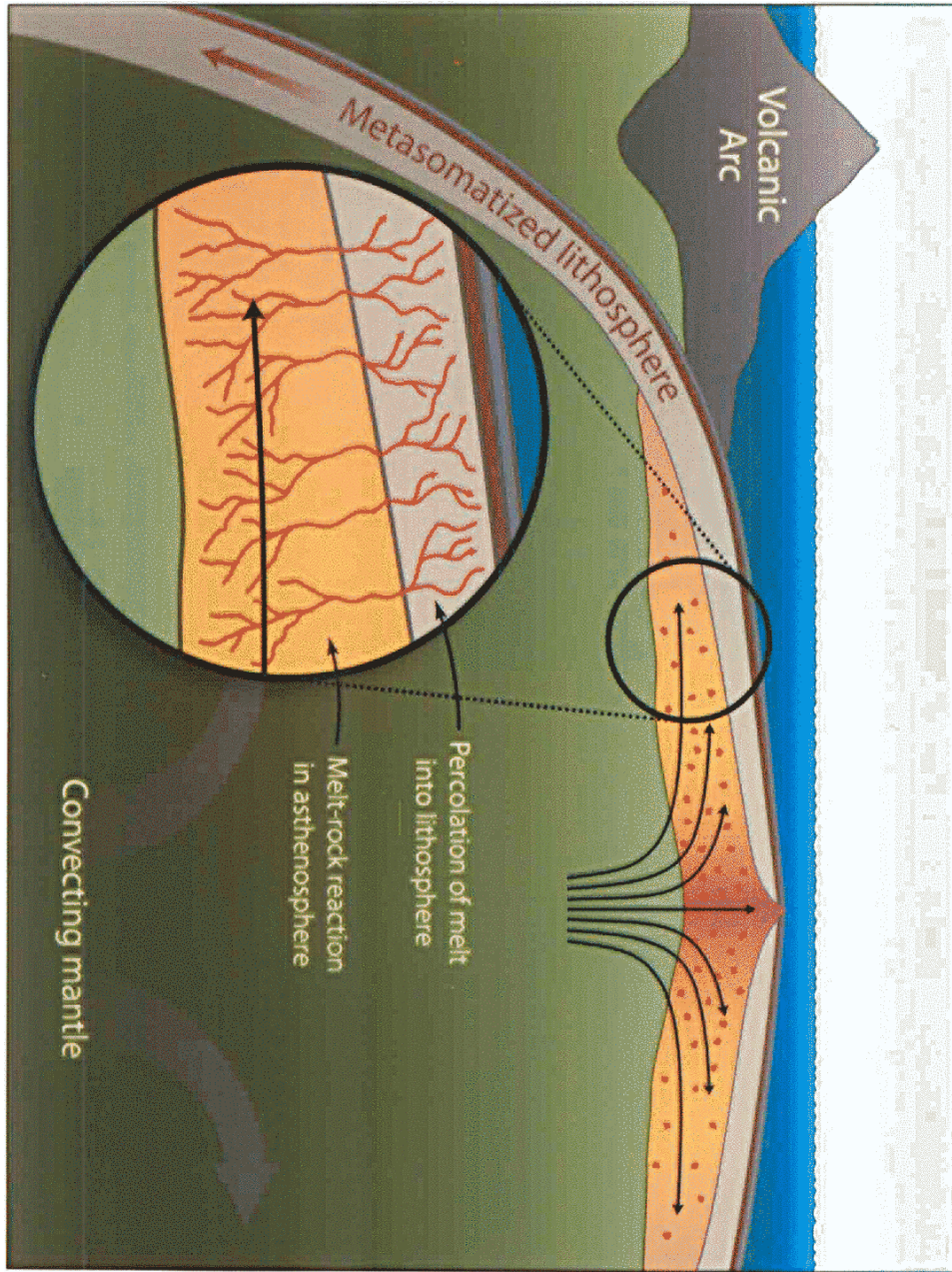


Figure 11. The Sr-Nd-Pb isotopic variation in MORB and ocean island basalts. Carlson, 1994

- Note much larger "dispersion" of southern hemisphere basalts







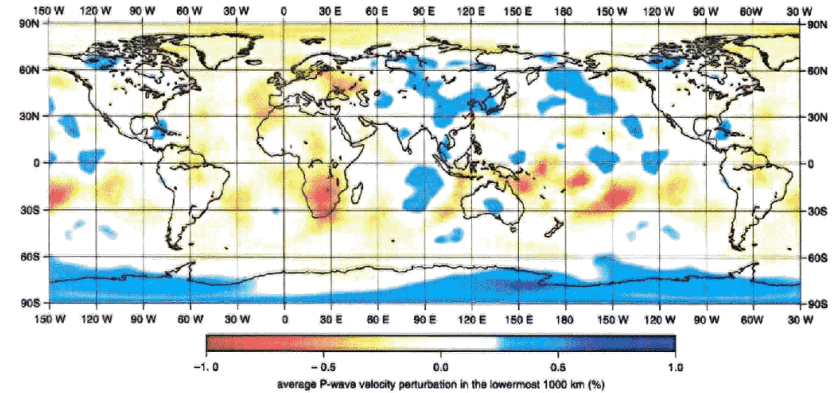
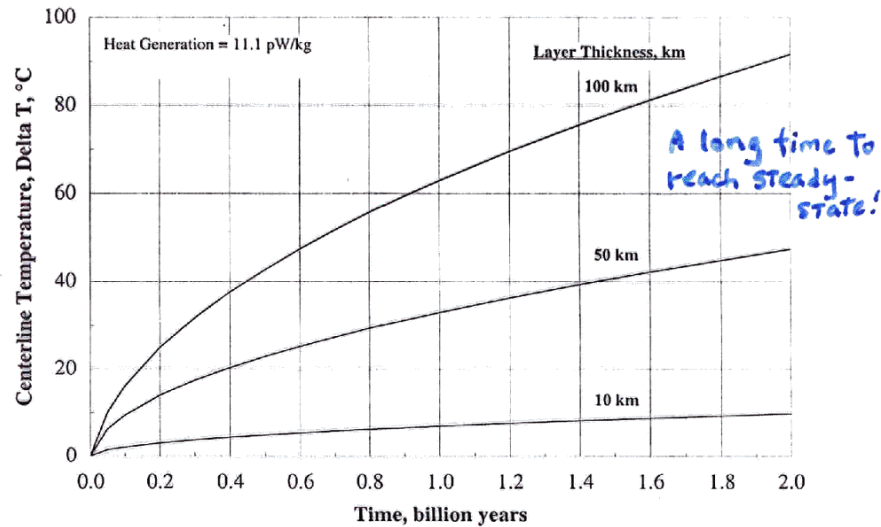
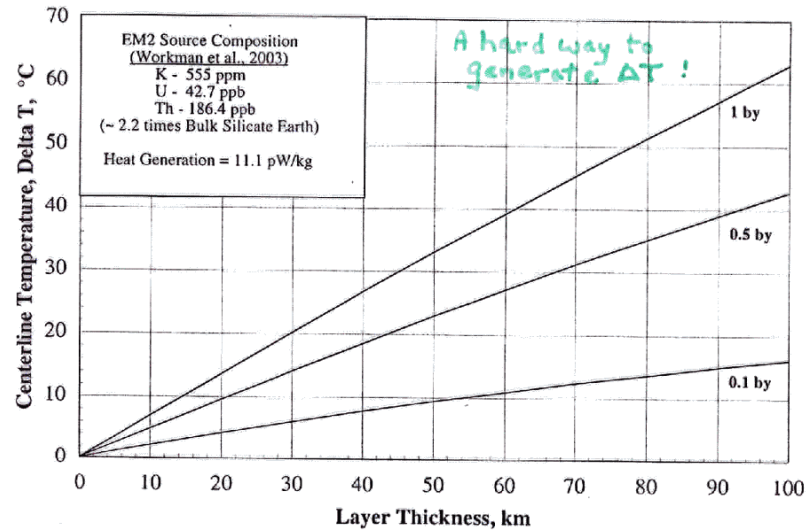
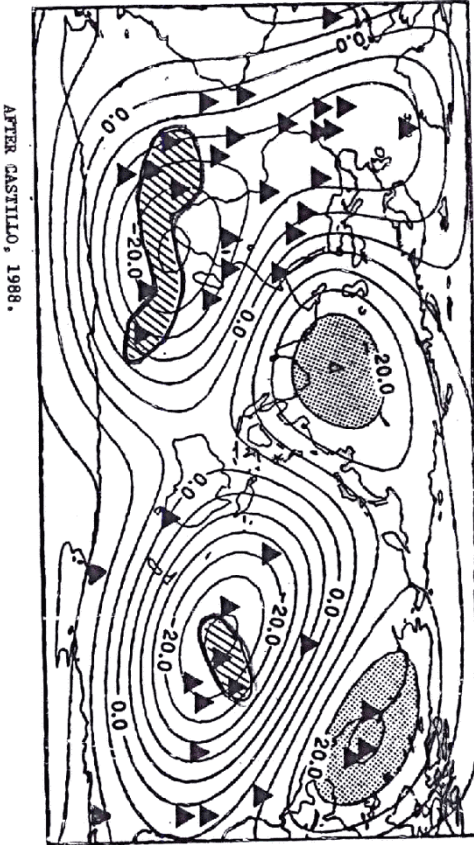


Fig. 1. Vertical average over the lowermost 1000 km of the mantle of the relative velocity perturbation $\delta v/v_p$. The averaging emphasizes features that are continuous with depth. Map has been wrapped around to have complete views of both the Atlantic and the Pacific oceans.

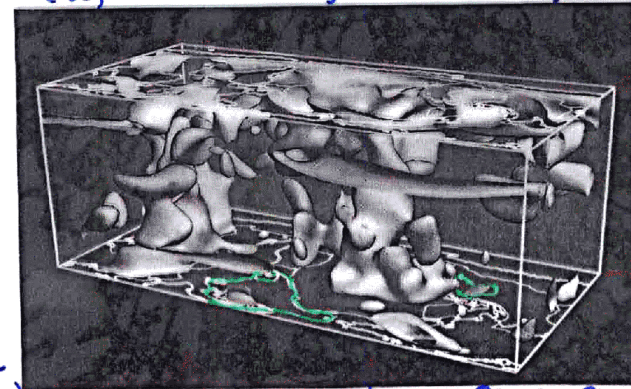
▲ - hot spots
Dupal "highs"

P-wave velocity deviations
averaged over lower mantle



THE HARVARD 'GRAND STRUCTURES' OF THE EARTH

(Su, Woodward + Dziewonski 1994)



Moho LEVEL

CMB

GREAT AFRICAN PLUME

PACIFIC PLUME GROUP

- Volumes with SHEAR VELOCITIES 0.6% Slower than mean
- VIEW is from the NW
- NORTH AMERICA + AUSTRALIA ARE outlined in green.



2500 km

3