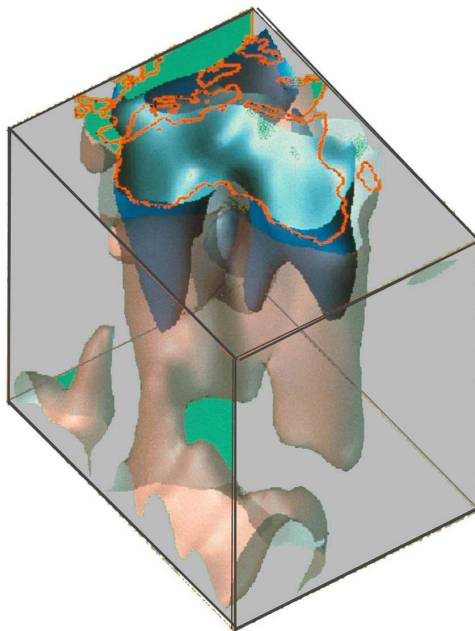


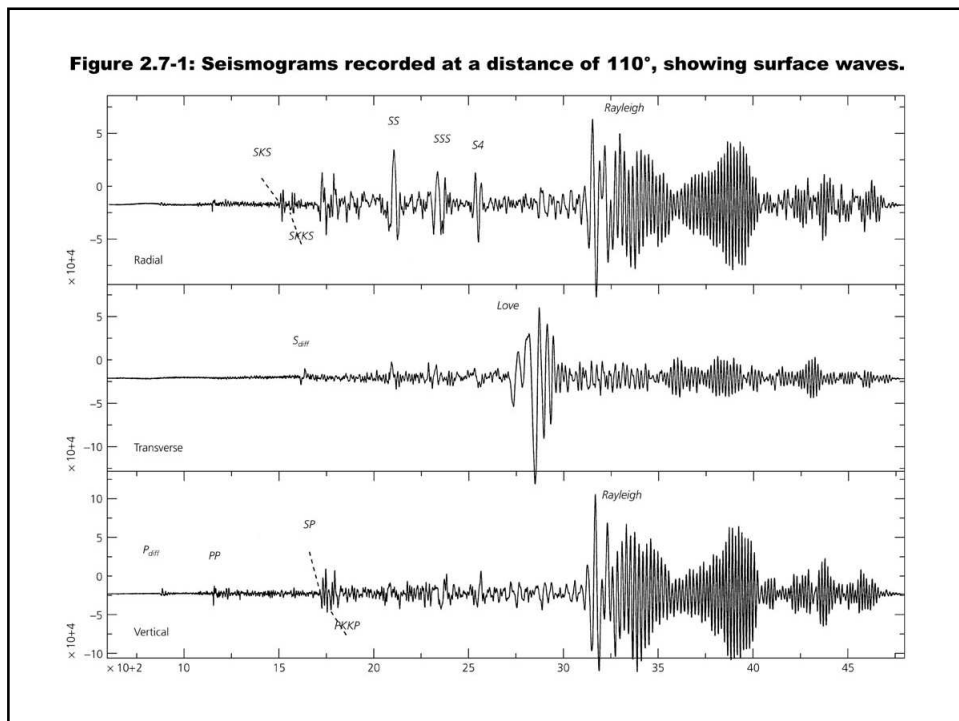
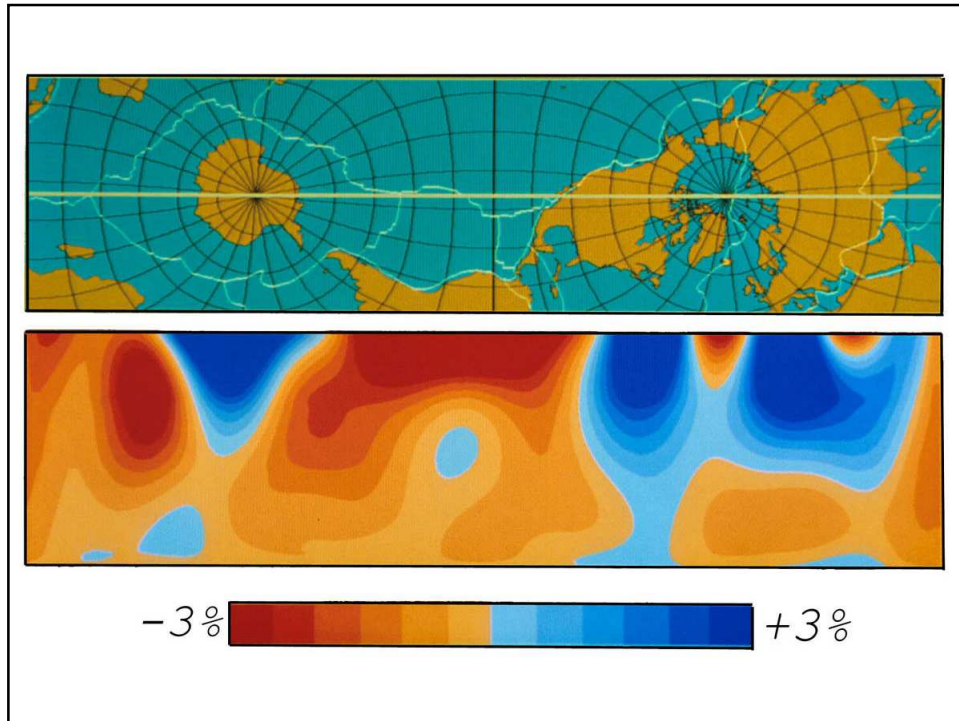
# ***Seismic Surface Waves***

***Adam M. Dziewonski***

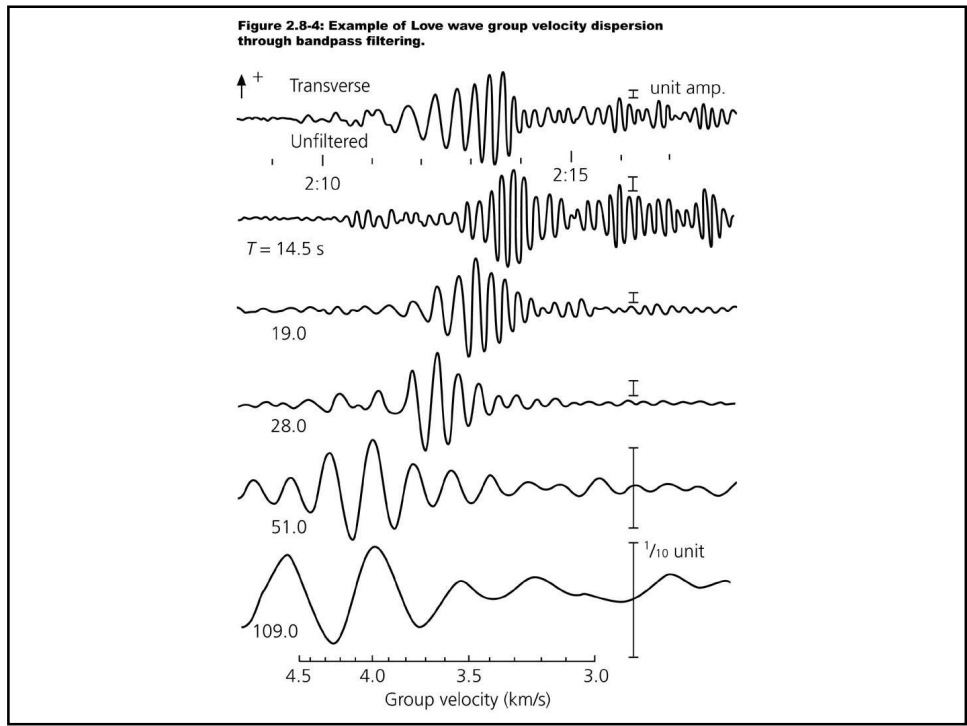
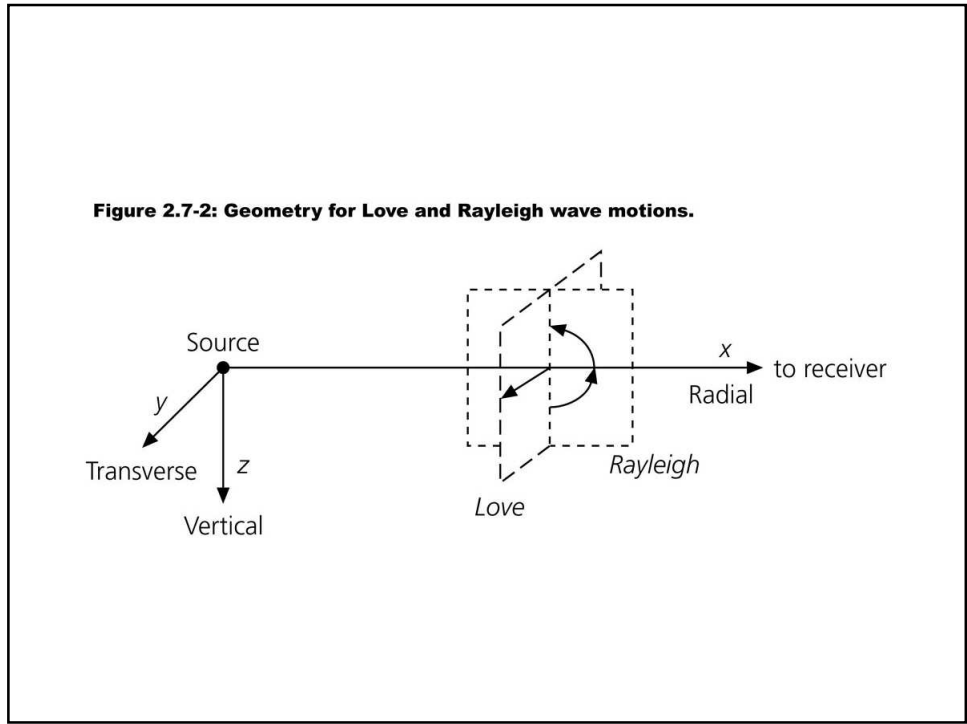
***July 17, 2004***

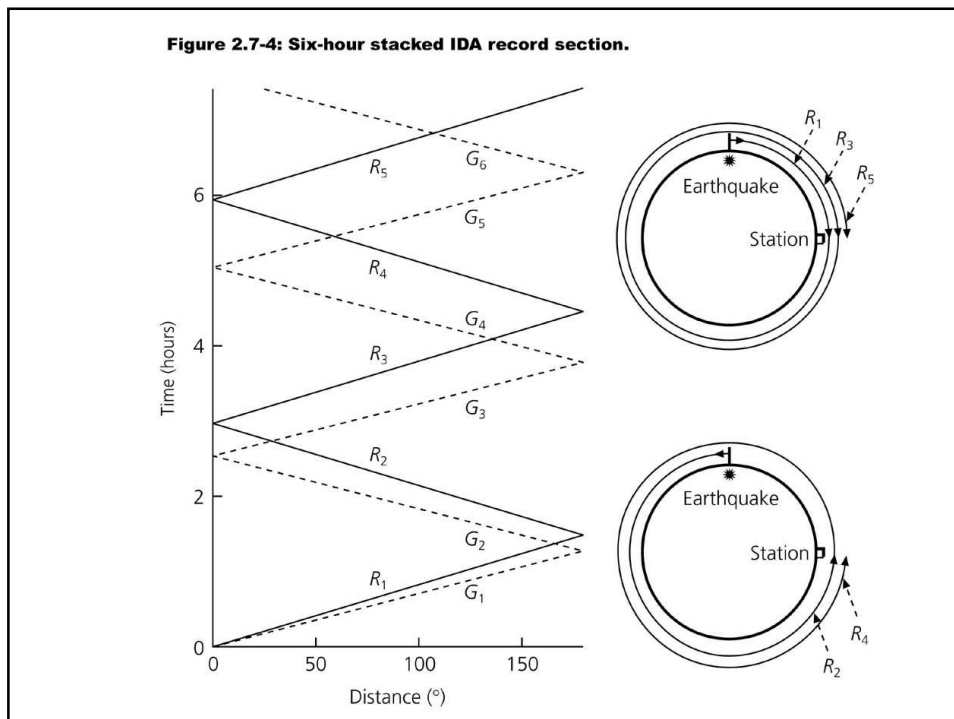
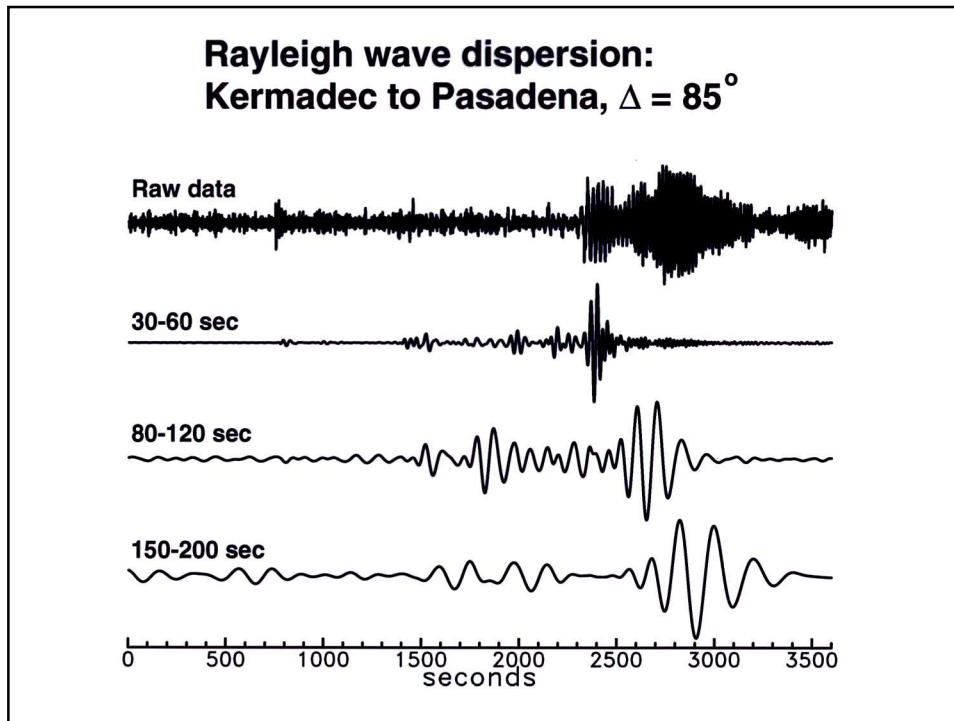


Seismology: Surface waves: excitation, measurements of dispersion and attenuation

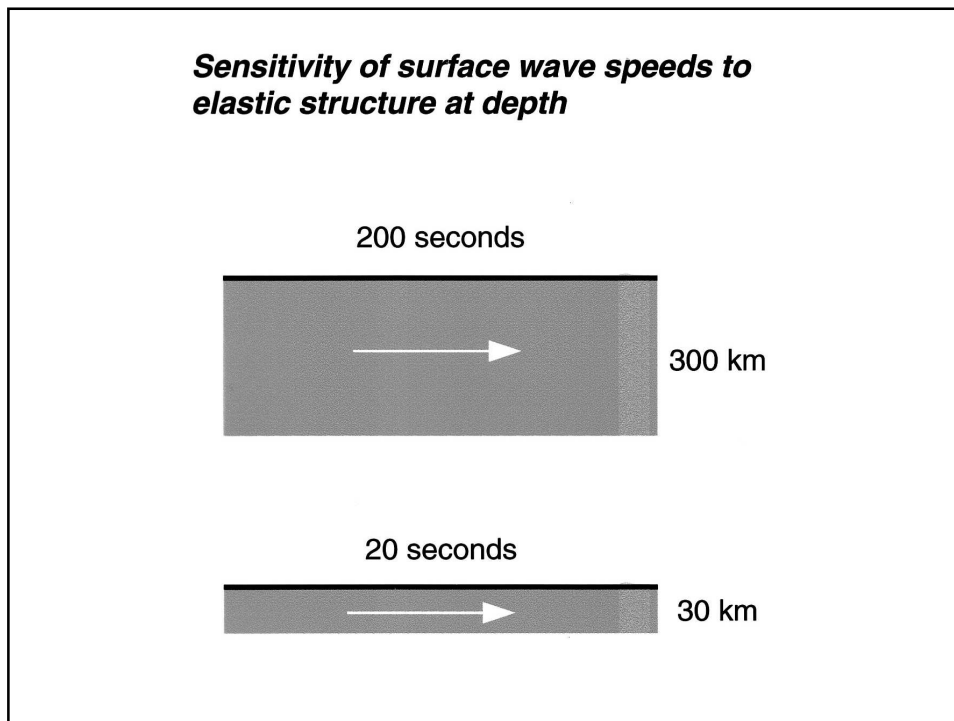
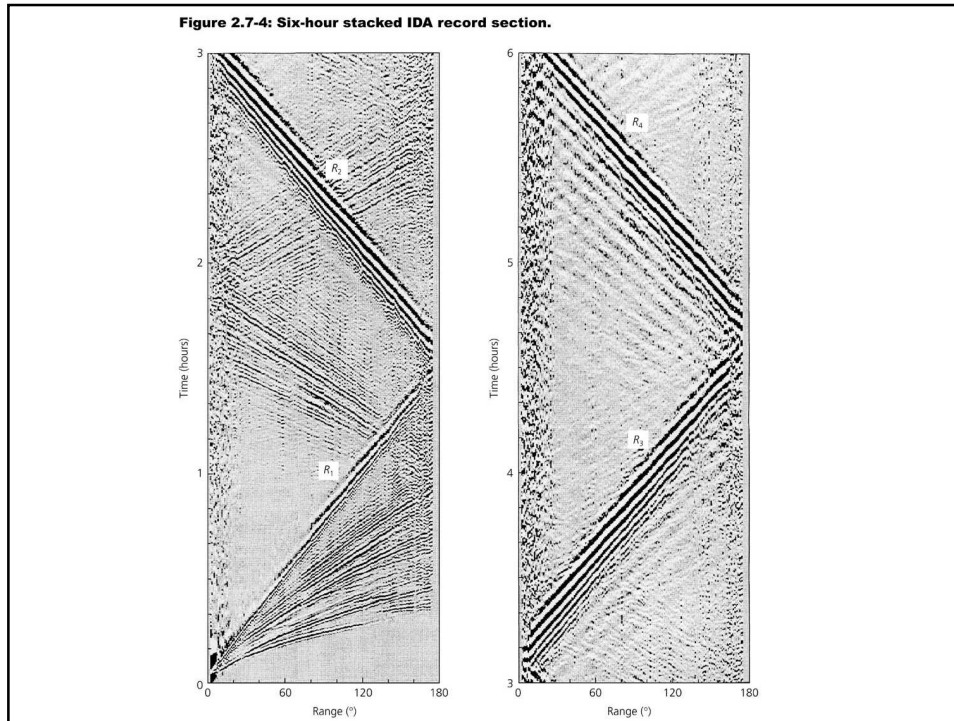


Seismology: Surface waves: excitation, measurements of dispersion and attenuation

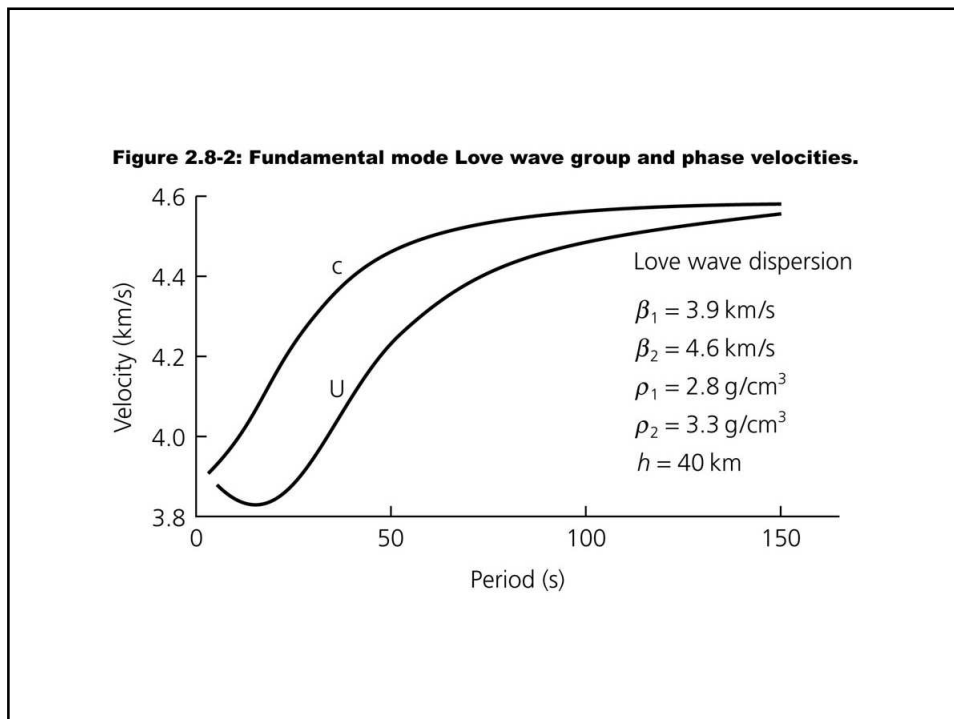
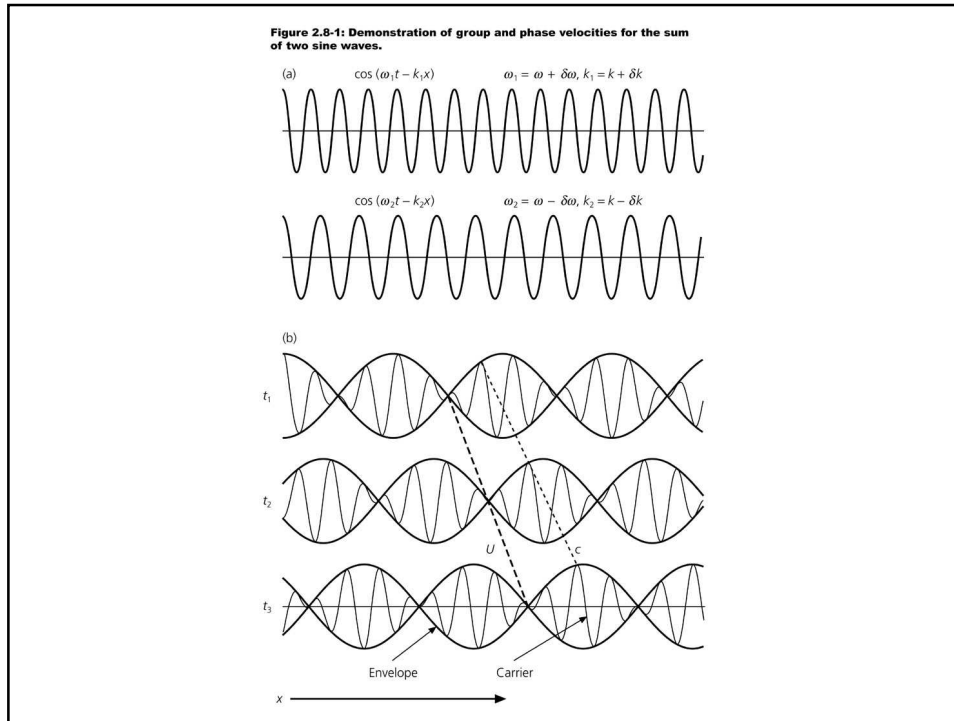




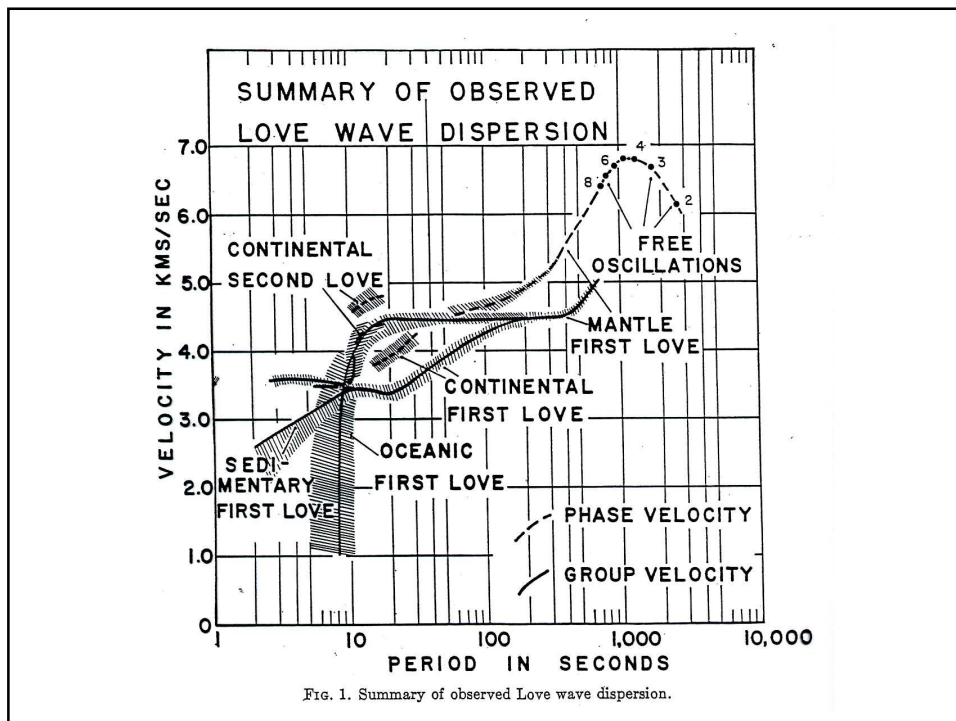
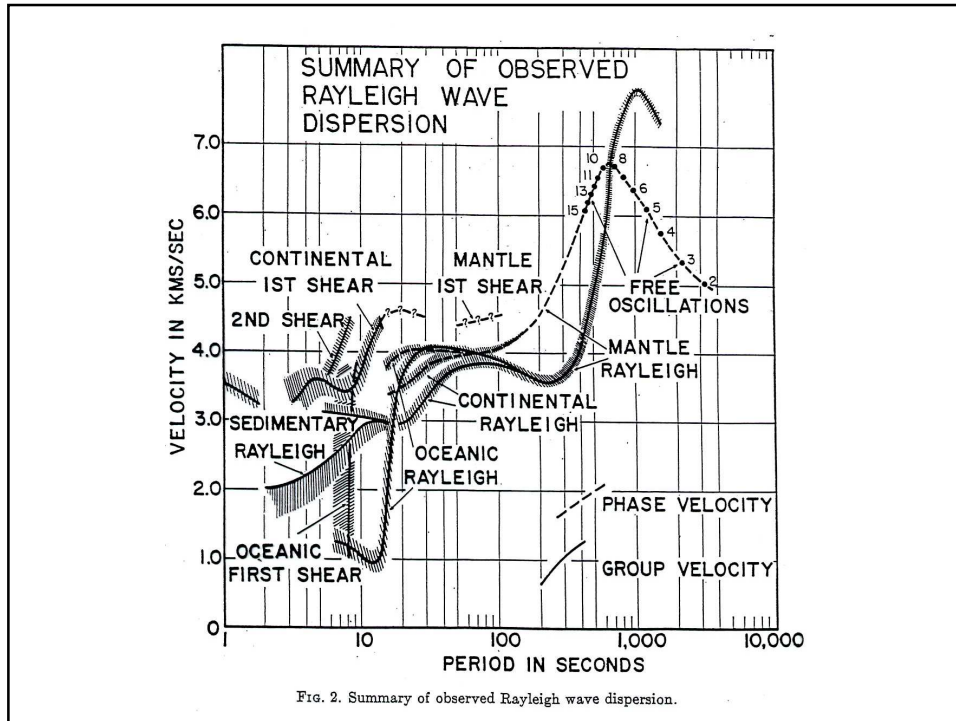
Seismology: Surface waves: excitation, measurements of dispersion and attenuation

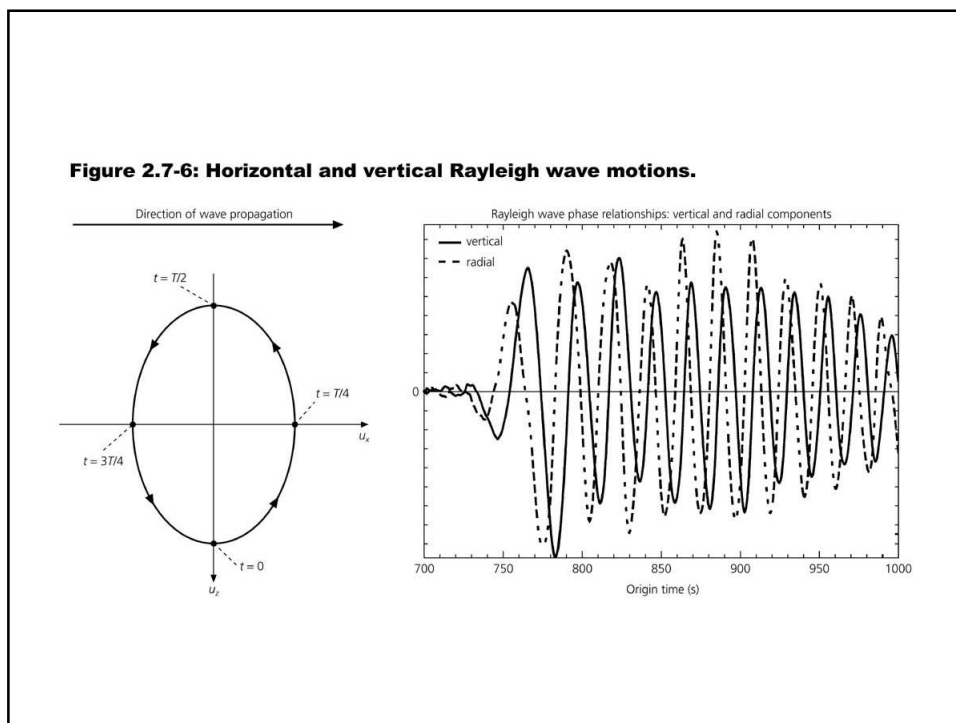
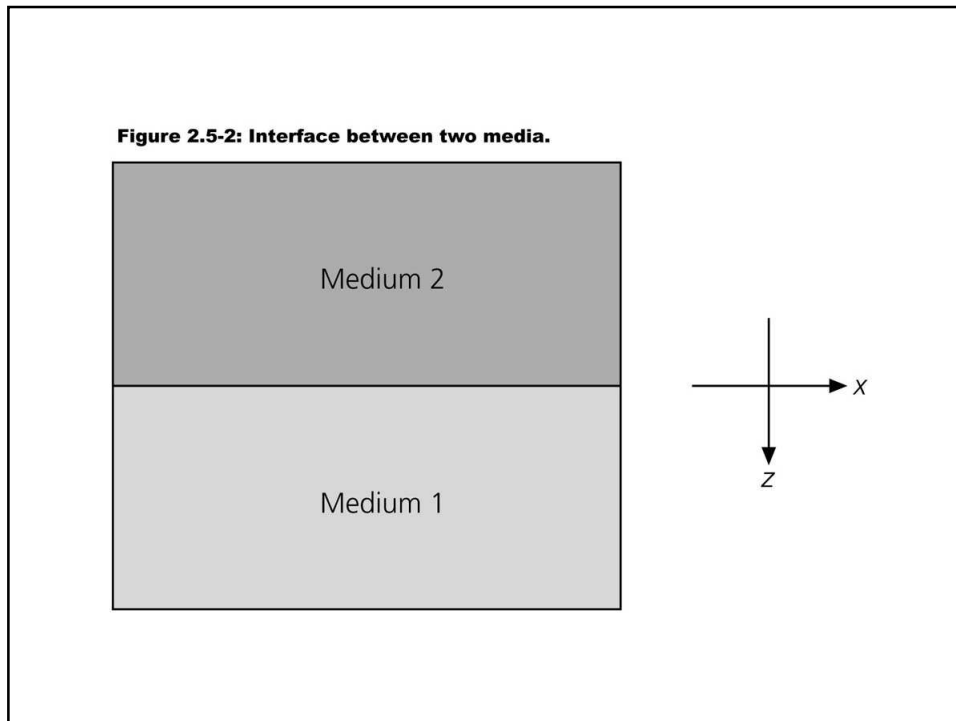


Seismology: Surface waves: excitation, measurements of dispersion and attenuation



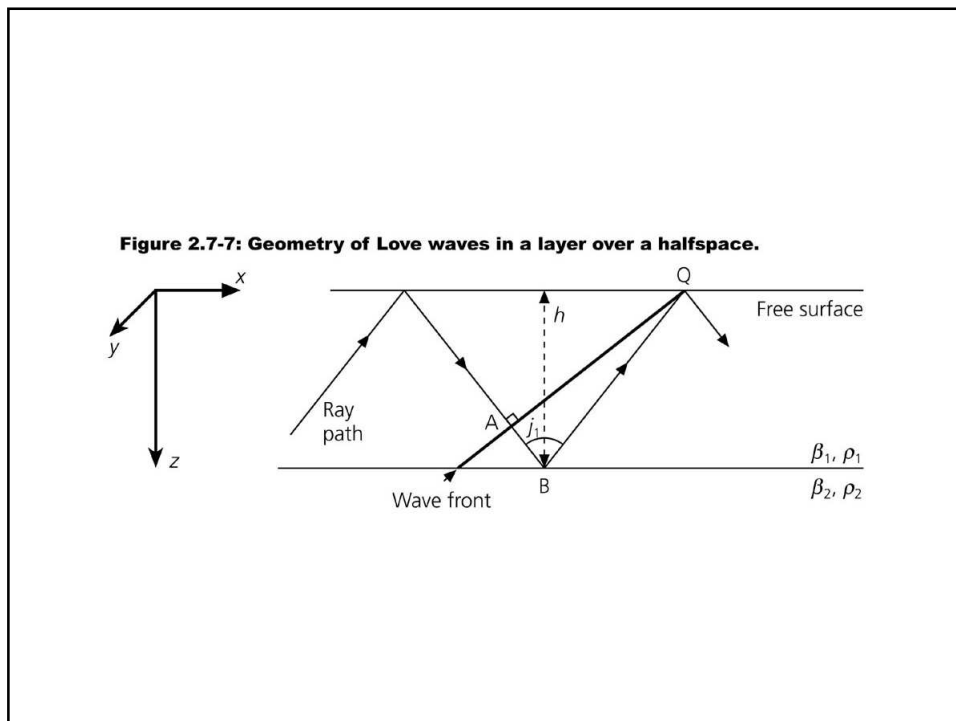
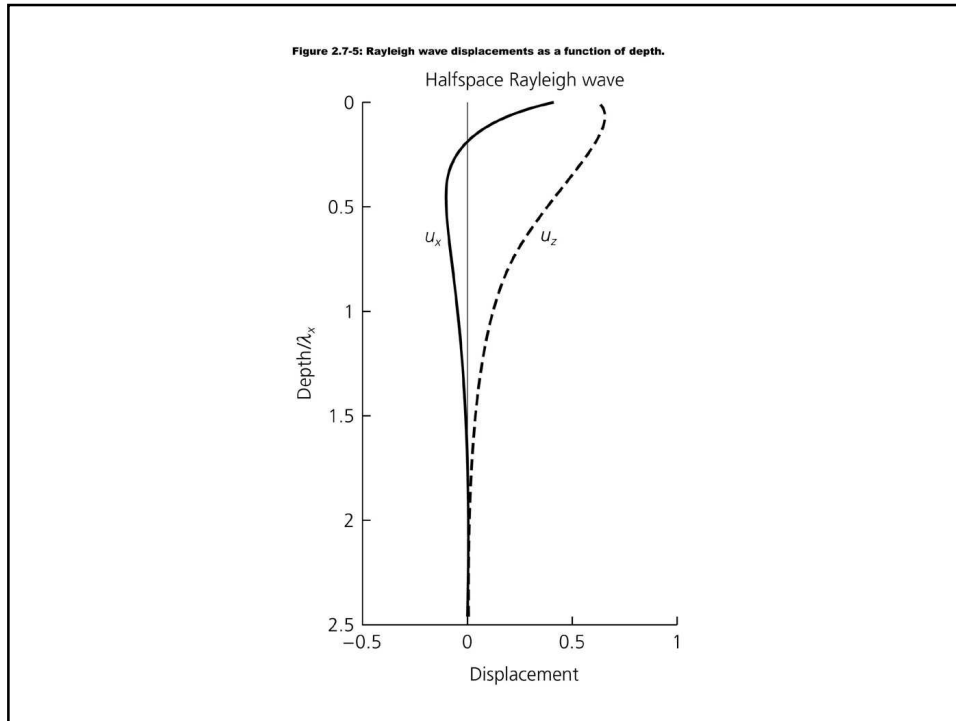
Seismology: Surface waves: excitation, measurements of dispersion and attenuation



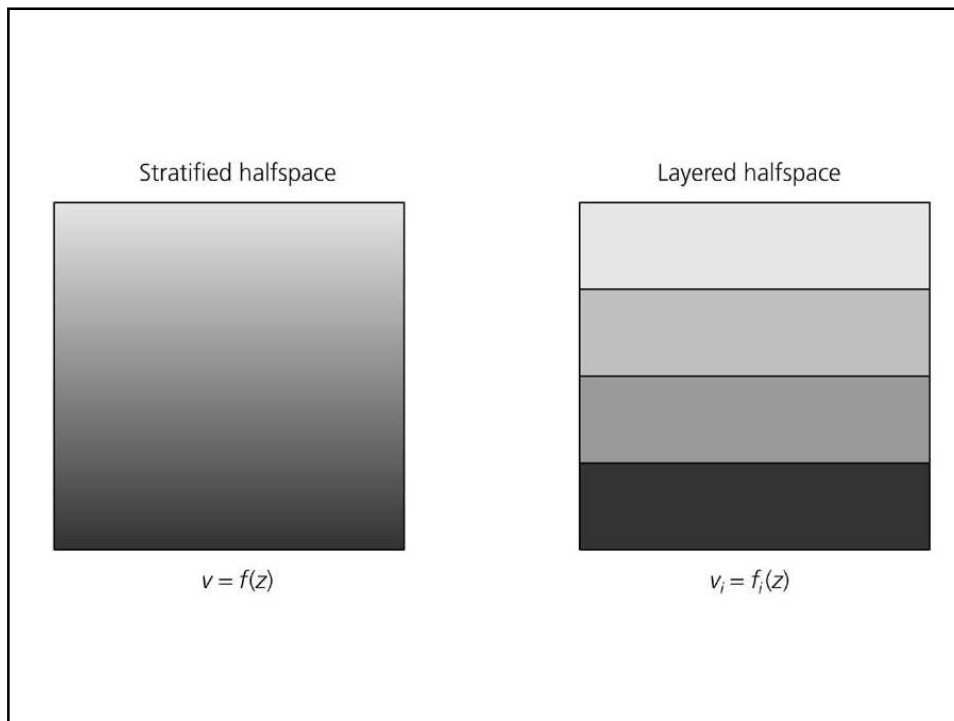
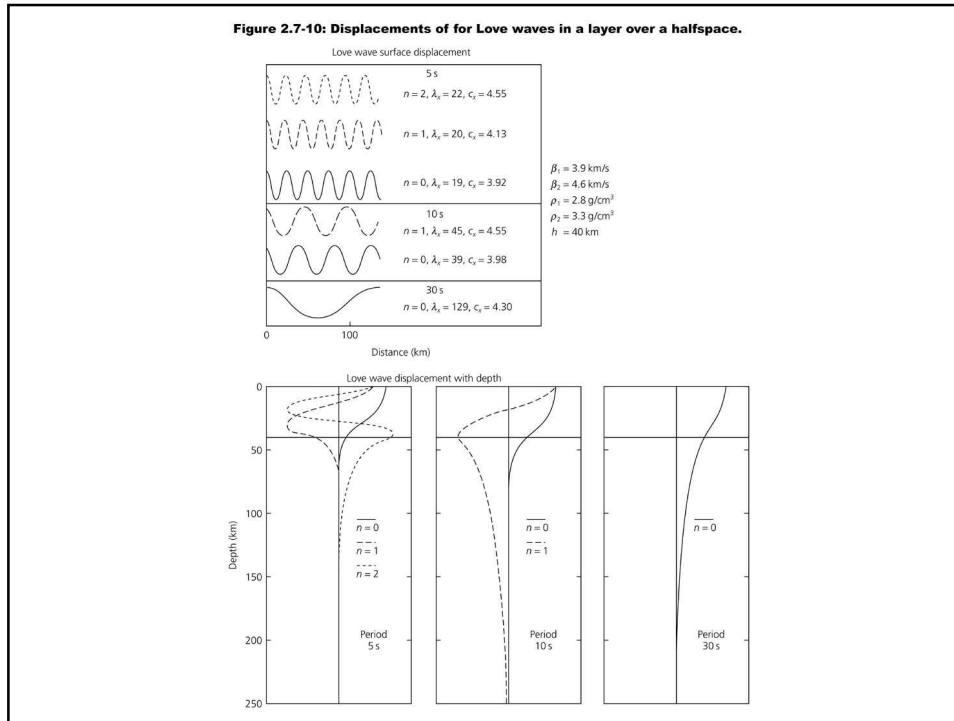


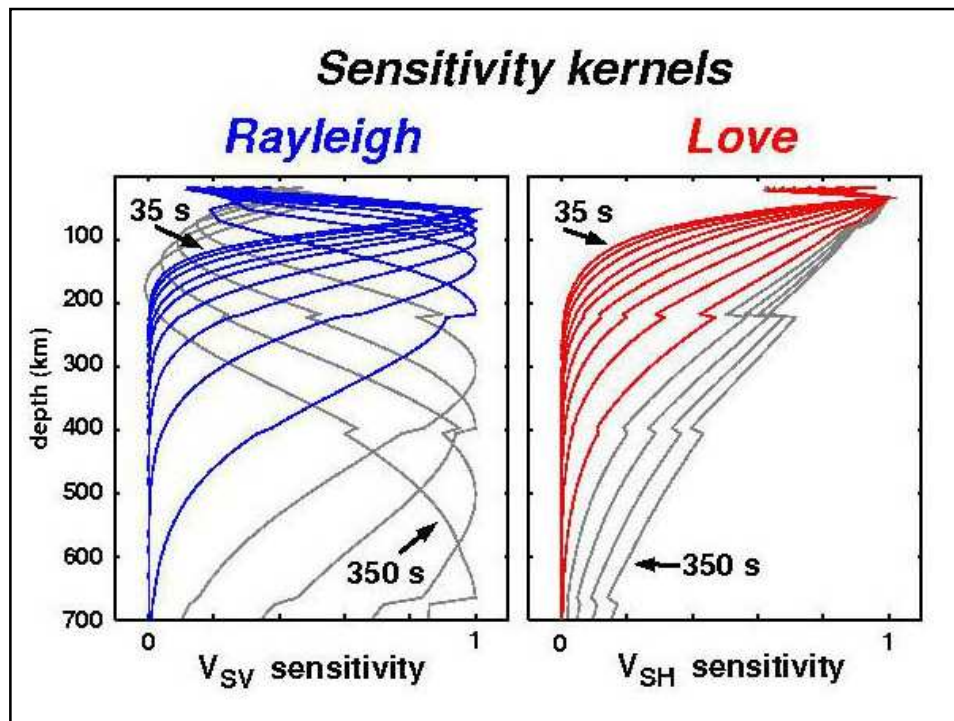


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Seismology: Surface waves: excitation, measurements of dispersion and attenuation





A surface wave seismogram  $u(\omega)$  can be written

$$u(\omega) = A(\omega) \exp[i\Phi(\omega)]. \quad (1)$$

The phase  $\Phi$  is the sum of two terms,

$$\Phi = \Phi_S + \Phi_P. \quad (2)$$

$\Phi_S$  is the source phase and  $\Phi_P$  is the propagation phase,

$$\Phi_P(\omega) = \int \frac{\omega}{c(\omega)} ds. \quad (3)$$

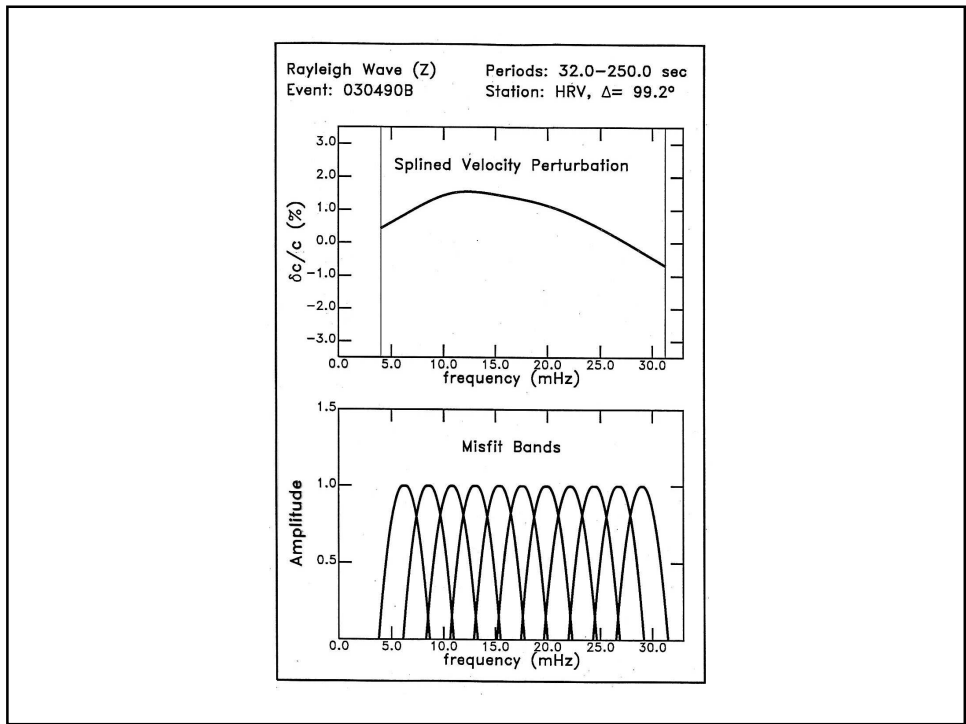
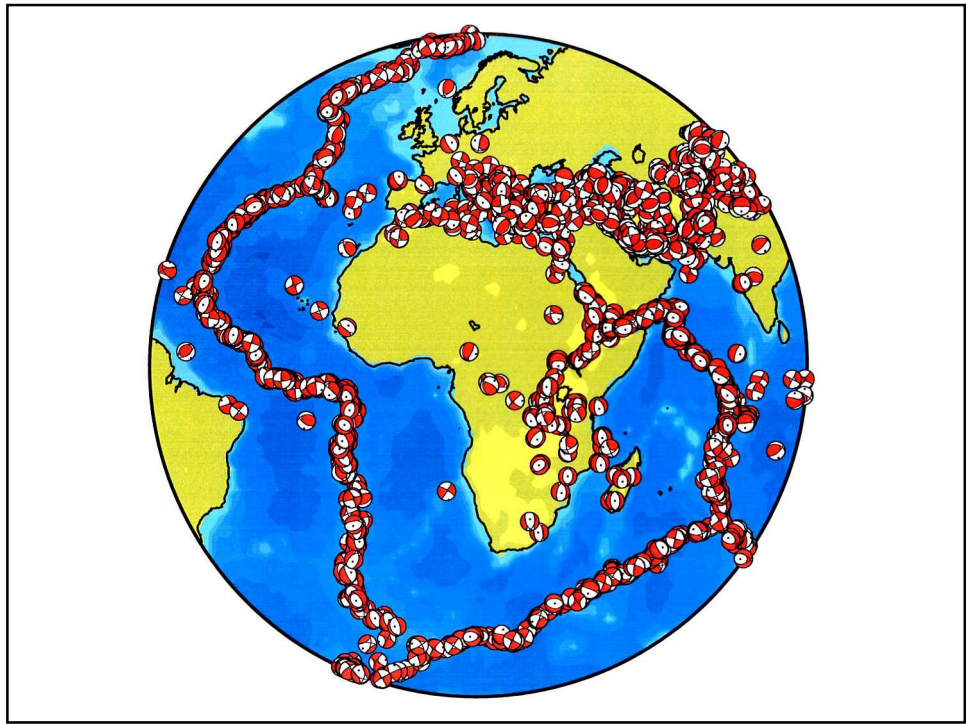
For a spherical reference Earth we write

$$\Phi_P^0(\omega) = \frac{\omega R \Delta}{c^0} = \frac{\omega X}{c^0}. \quad (4)$$

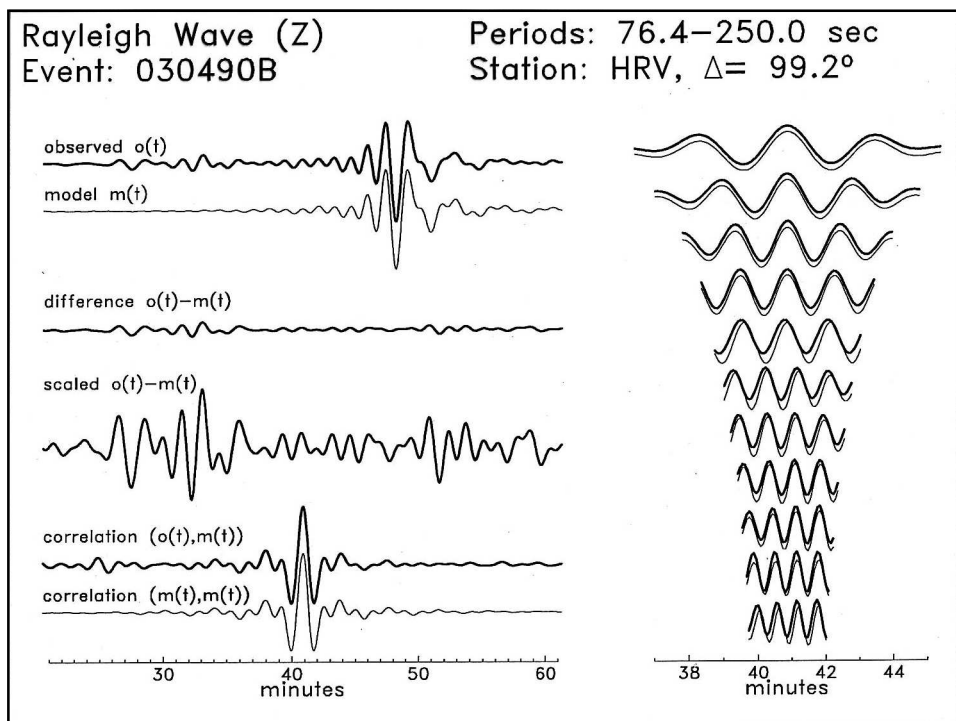
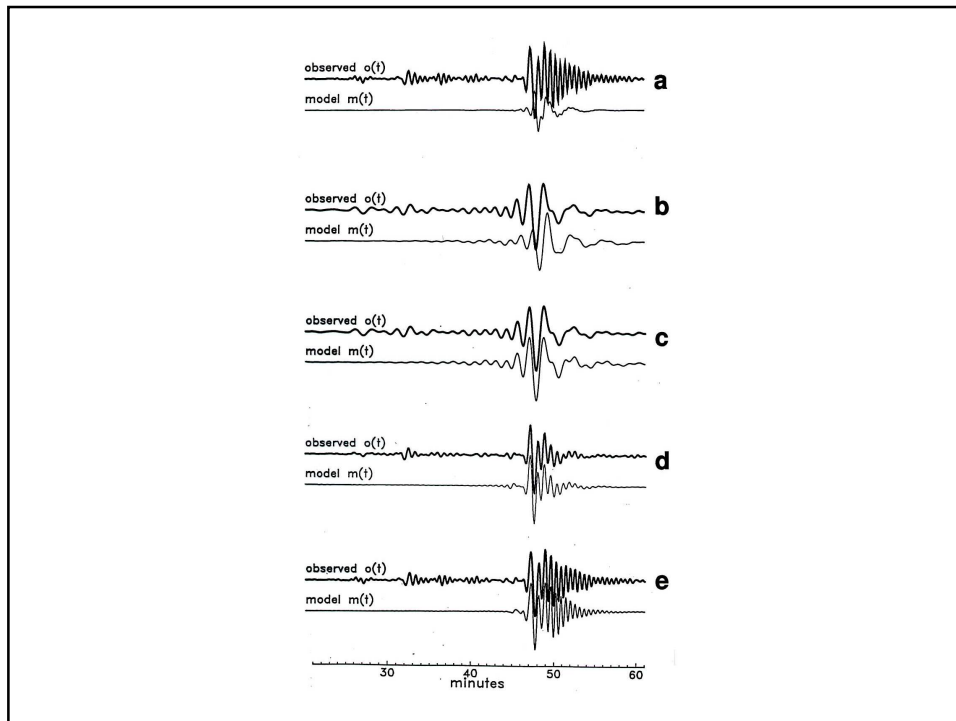
We approximate the observed phase as

$$\Phi_P = \Phi_P^0 + \delta\Phi = \frac{\omega X}{c^0 + \delta c}. \quad (5)$$

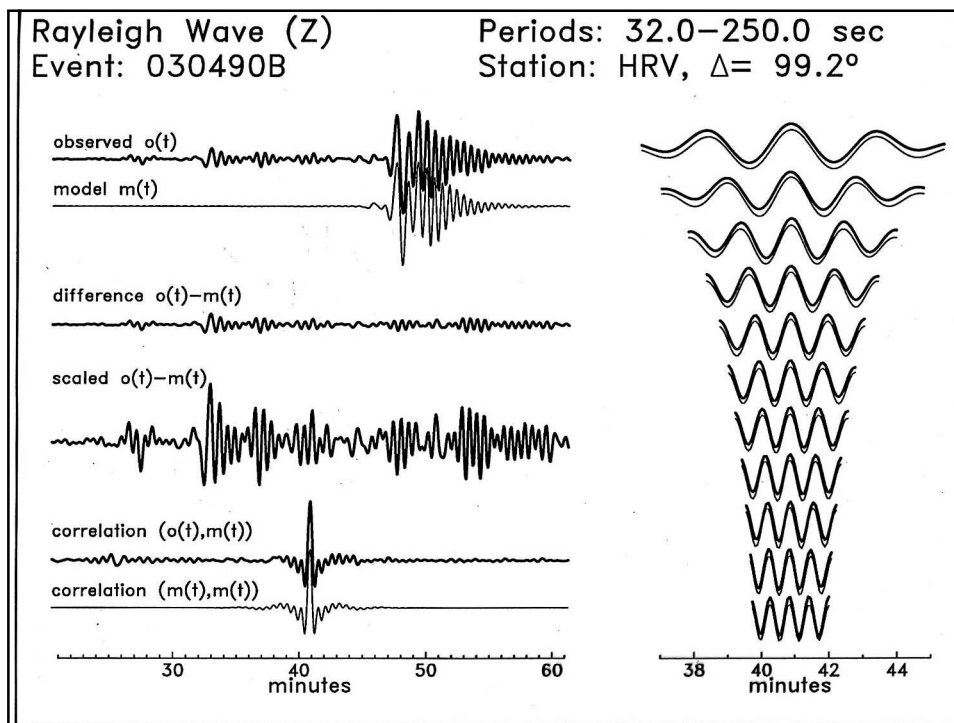
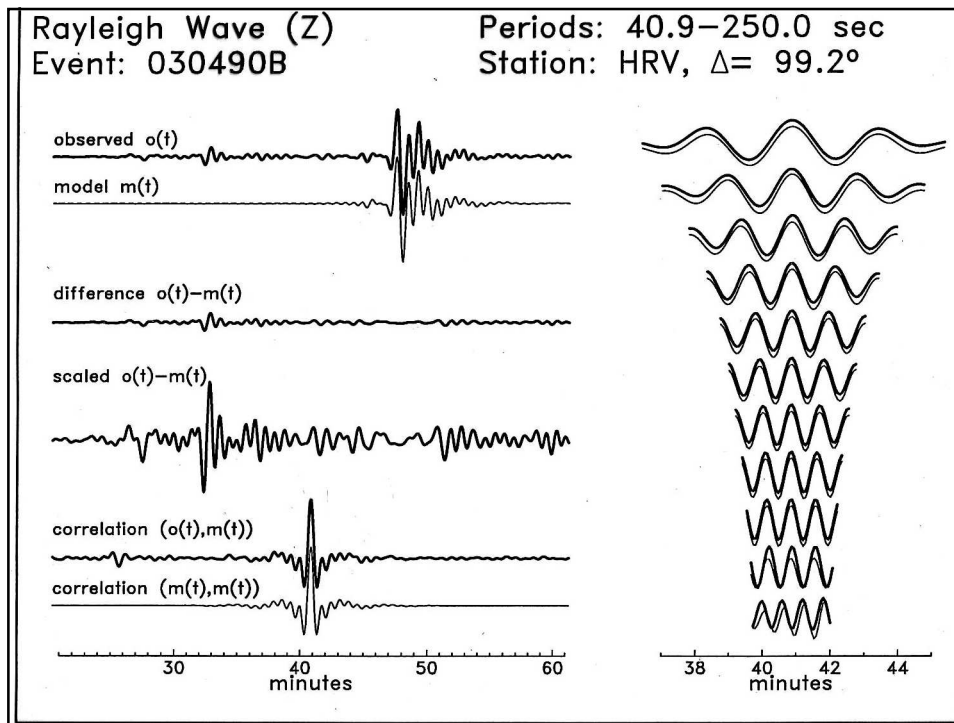
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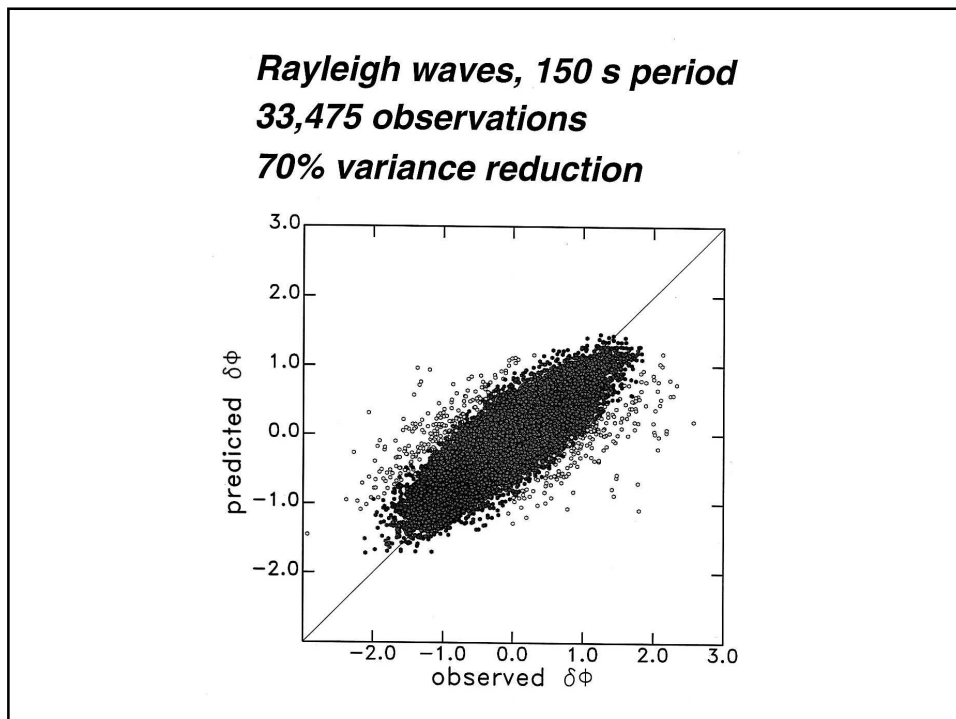
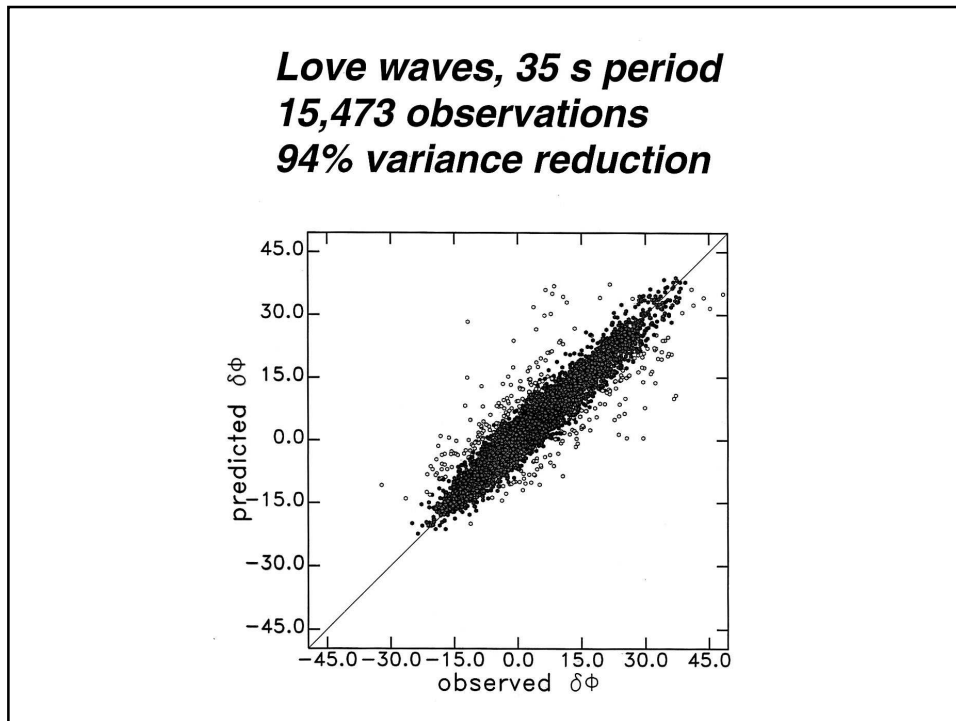


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Seismology: Surface waves: excitation, measurements of dispersion and attenuation





Seismology: Surface waves: excitation, measurements of dispersion and attenuation

The phase anomaly is written

$$\delta\Phi = -\frac{\omega}{c^0} \int_{(\theta_S, \phi_S)}^{(\theta_R, \phi_R)} \frac{\delta c(\theta, \phi)}{c^0} ds, \quad (1)$$

and the phase velocity perturbation is expressed in complex spherical harmonics

$$\frac{\delta c(\theta, \phi)}{c^0} = \sum_{l=0}^L \sum_{m=-l}^{m=l} C_{lm} Y_{lm}(\theta, \phi). \quad (2)$$

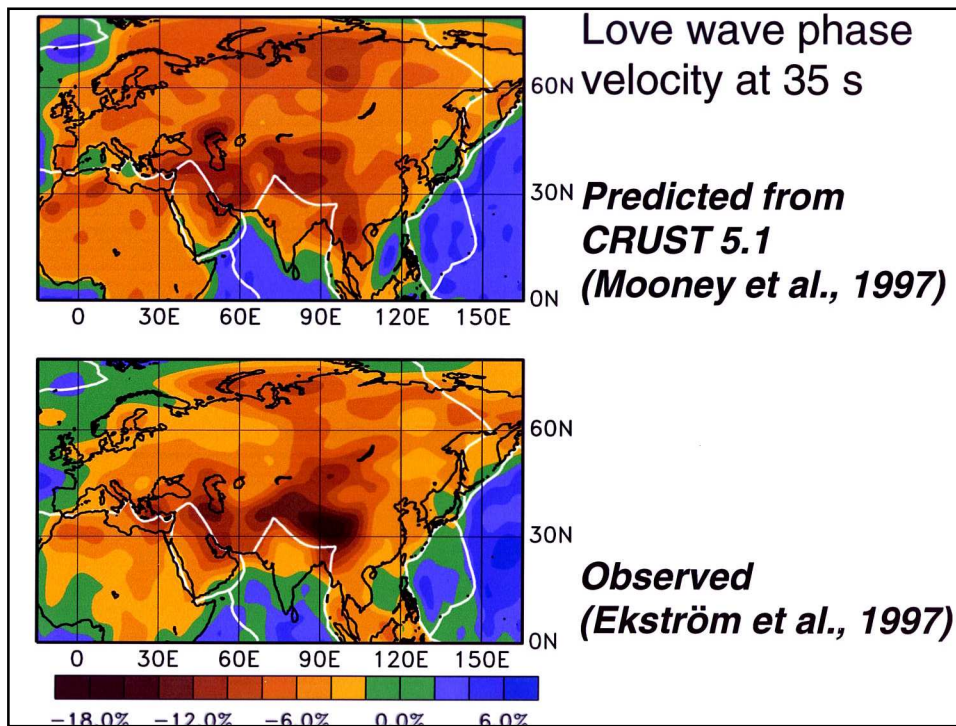
The rms gradient is

$$\mathcal{R} \propto \left[ \int_S (\nabla \frac{\delta c}{c^0}) \cdot (\nabla \frac{\delta c}{c^0}) d\Omega \right]^{1/2}. \quad (3)$$

We determine the model coefficients which solve the minimization problem

$$\min \left( \frac{\chi^2}{N} + \gamma \mathcal{R}^2 \right), \quad (4)$$

where  $\gamma$  is a damping parameter.





Interpreting the observed phase anomalies in terms of velocity variations in the six tectonic regions of GTR1 (Jordan, 1981).

The observed phase anomaly is approximated as

$$\delta\Phi = \frac{\omega X}{c^0 + \delta c} - \frac{\omega X}{c^0} \approx -\frac{\omega X \delta c}{(c^0)^2}. \quad (1)$$

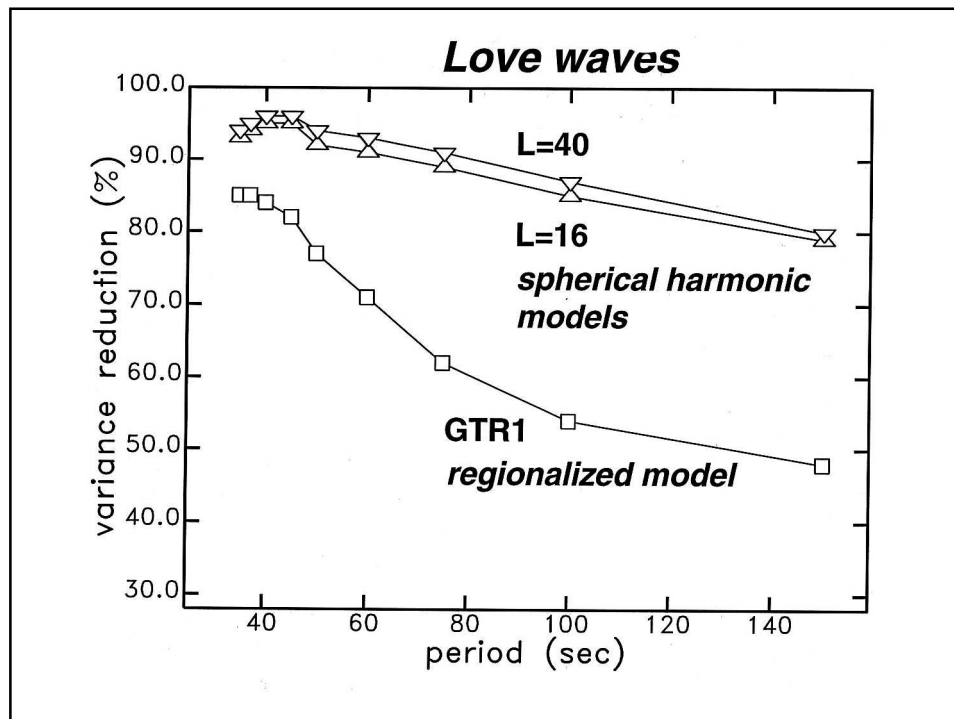
The phase anomaly  $\delta\Phi$  is modeled as

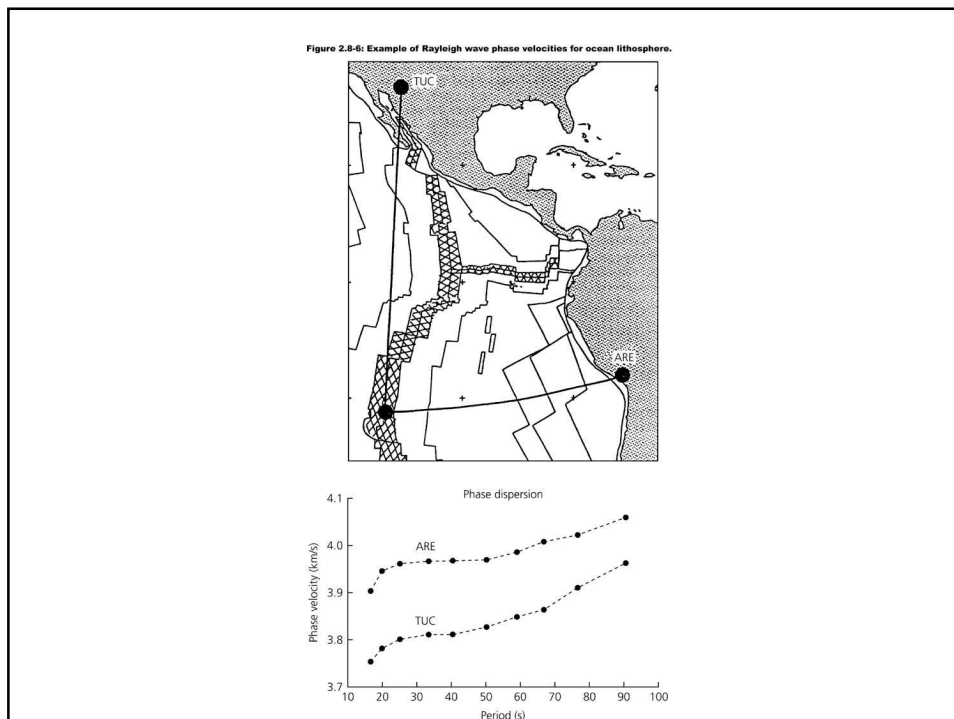
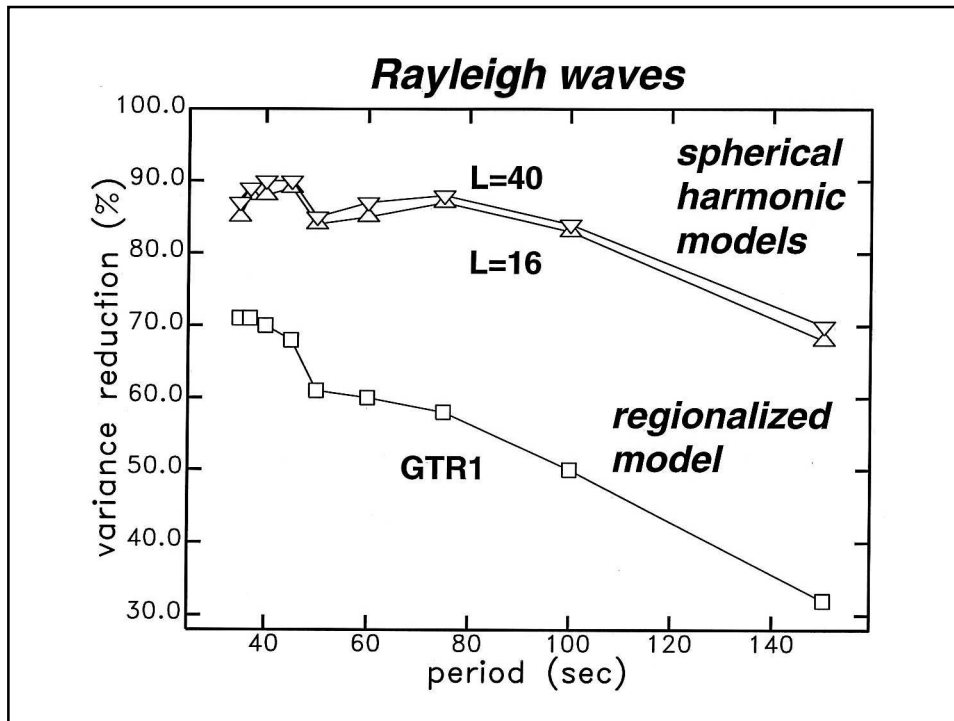
$$\delta\Phi = -\frac{\omega}{c^0} \sum_{i=1}^6 X_i \frac{\delta c^i}{c^0}. \quad (2)$$

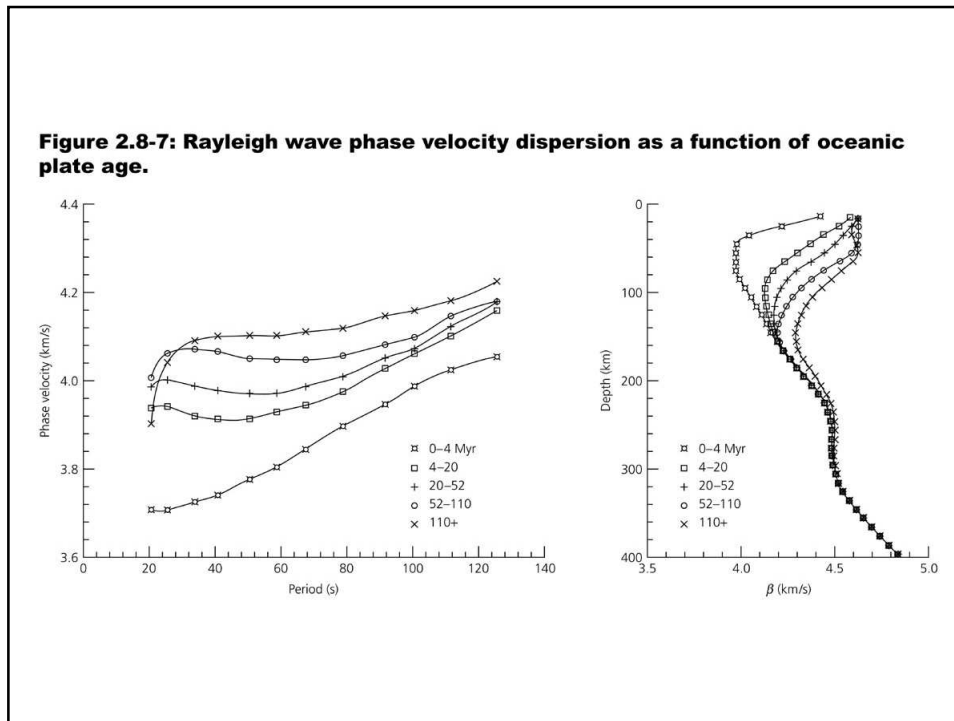
The misfit function  $\chi^2$  is formed as

$$\chi^2 = \sum_{j=1}^N \frac{1}{\sigma_j^2} \left( \delta\Phi_j - \frac{\omega}{c^0} \sum_{i=1}^6 X_j^i \frac{\delta c^i}{c^0} \right)^2, \quad (3)$$

where  $j$  is the index of the observation,  $N$  is the total number of observations, and  $\sigma_j$  is the associated observational uncertainty.







Radial Anisotropy =  
Transverse Isotropy  
with a Vertical Symmetry Axis

$$C_{ij} \rightarrow A, C, L, N, F$$

or

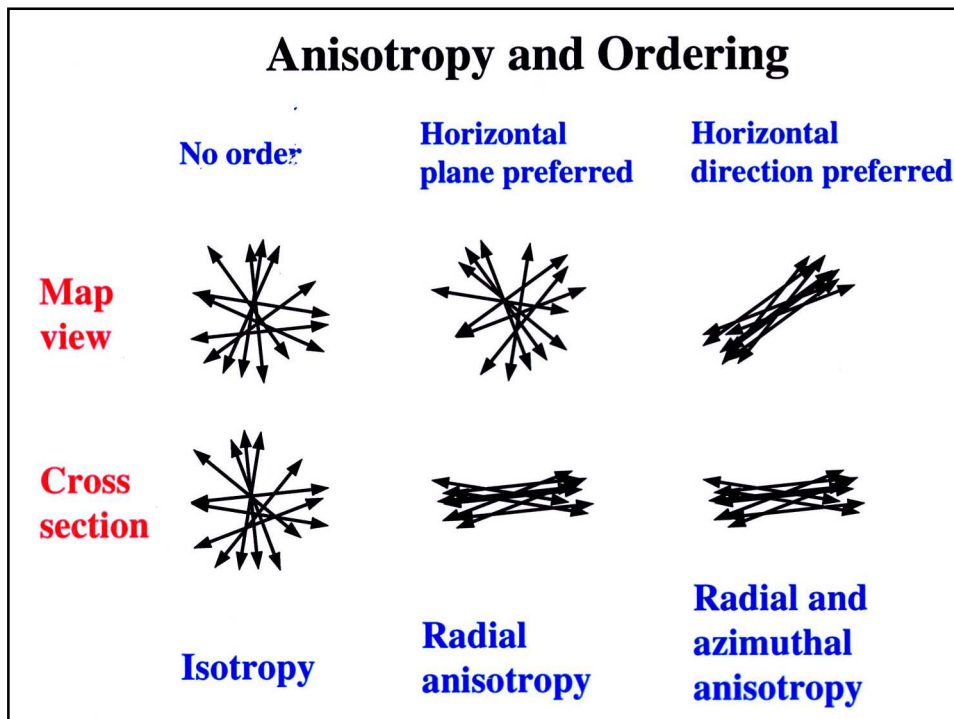
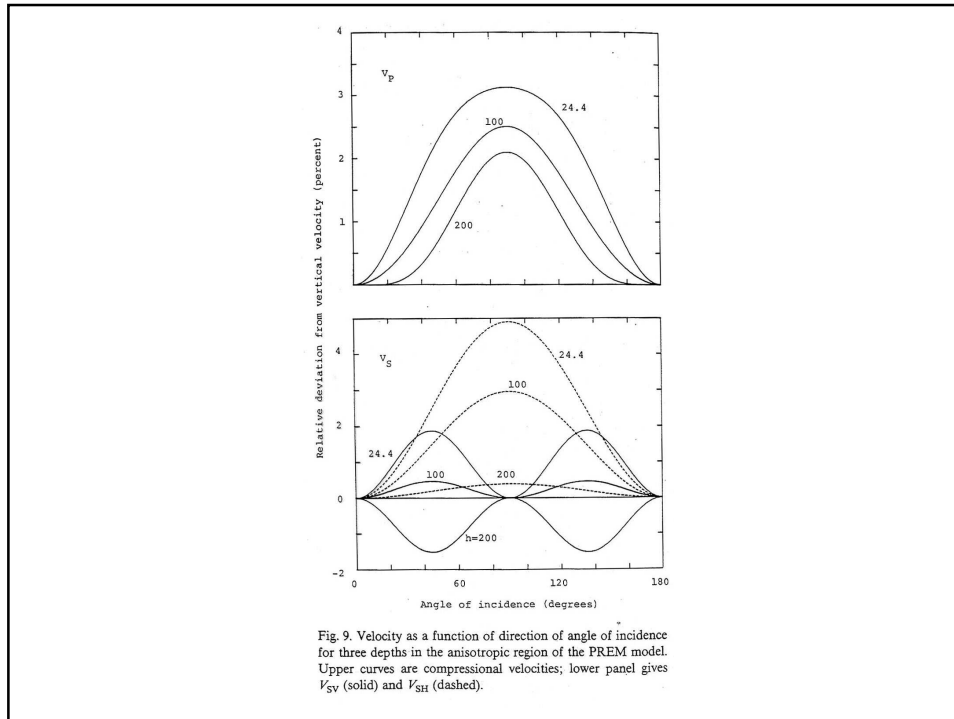
$$\rightarrow V_{SH}, V_{SV}, V_{PH}, V_{PV}, \eta$$

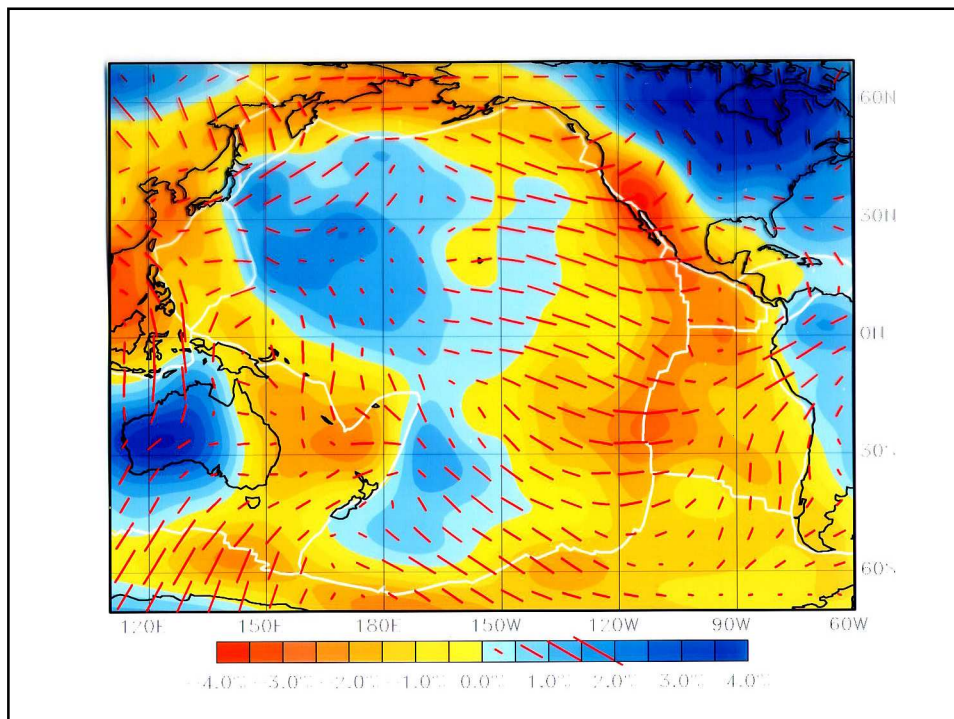
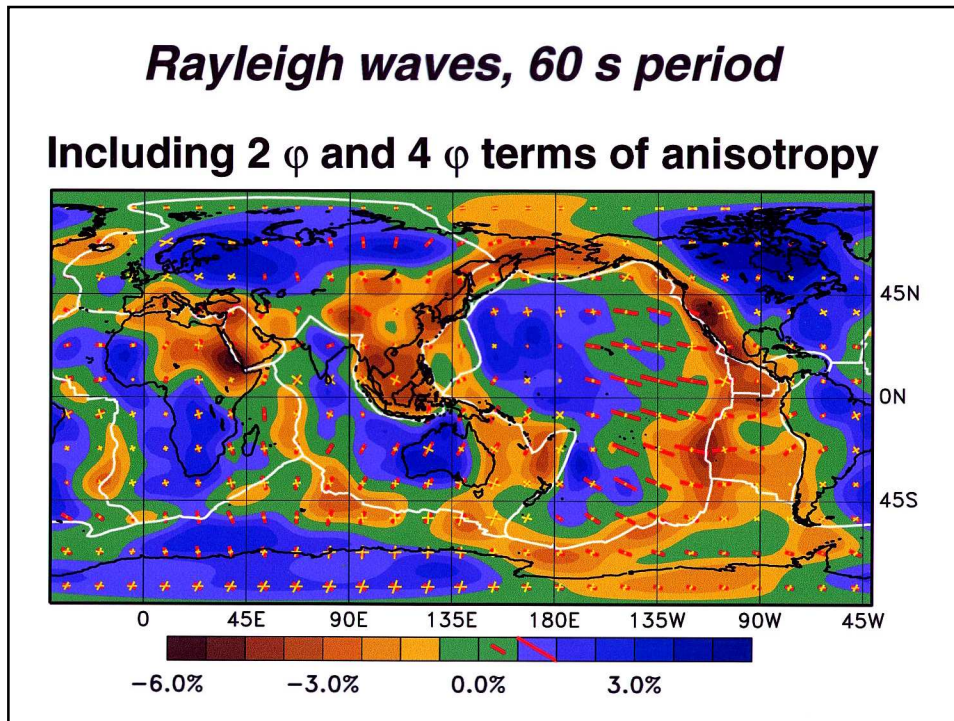
Common expectation:

$$\delta V_{PH} \sim \delta V_{SH}$$

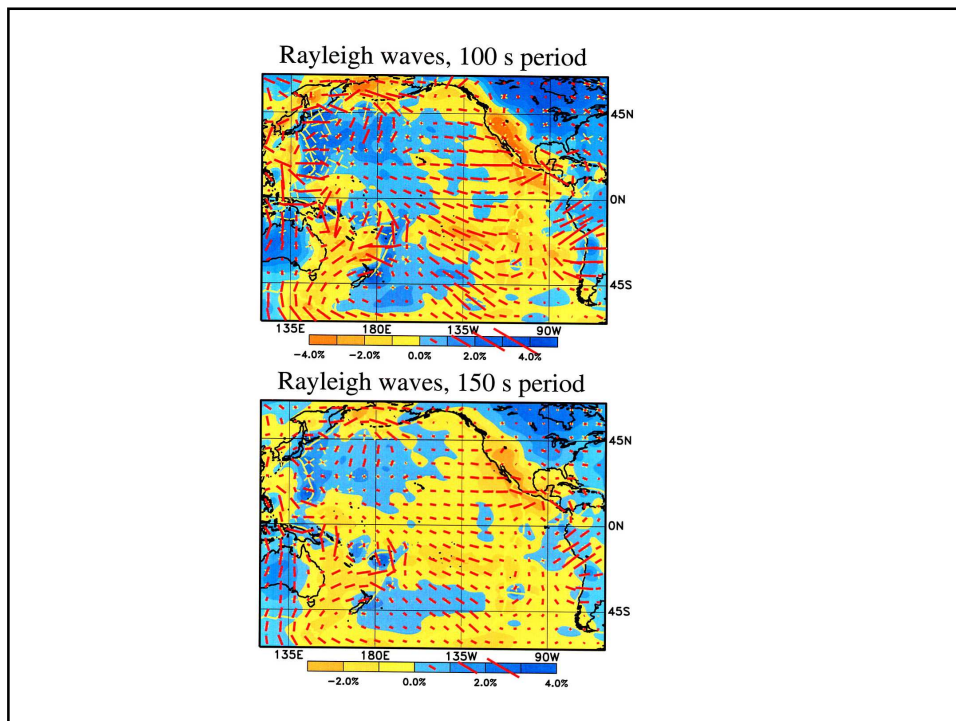
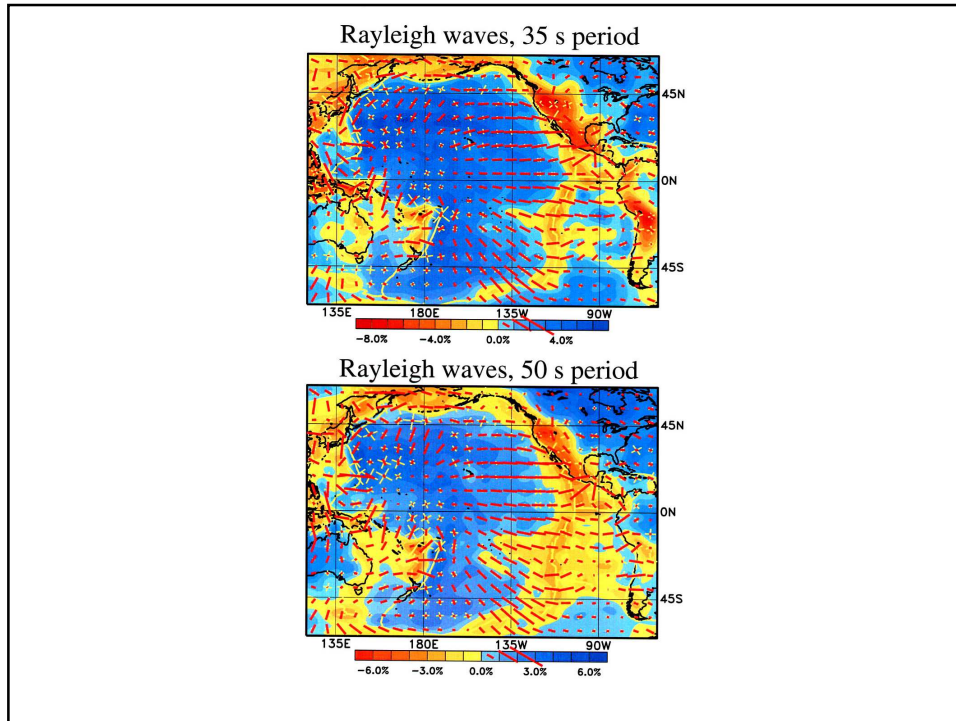
$$\delta V_{PV} \sim \delta V_{SV}$$

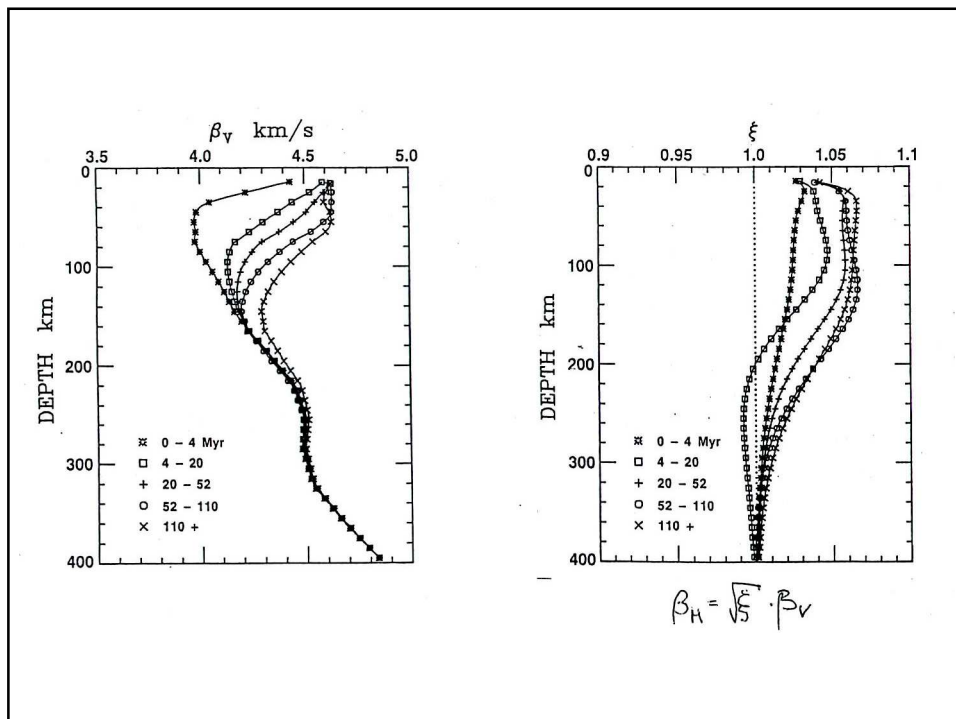
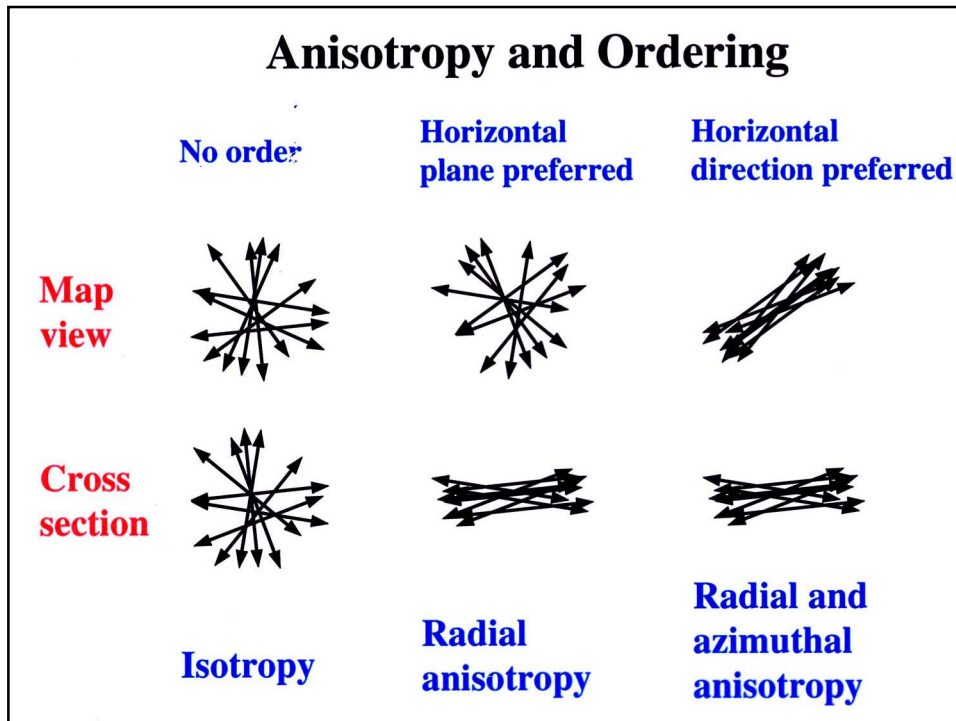
$$\eta \sim 1$$



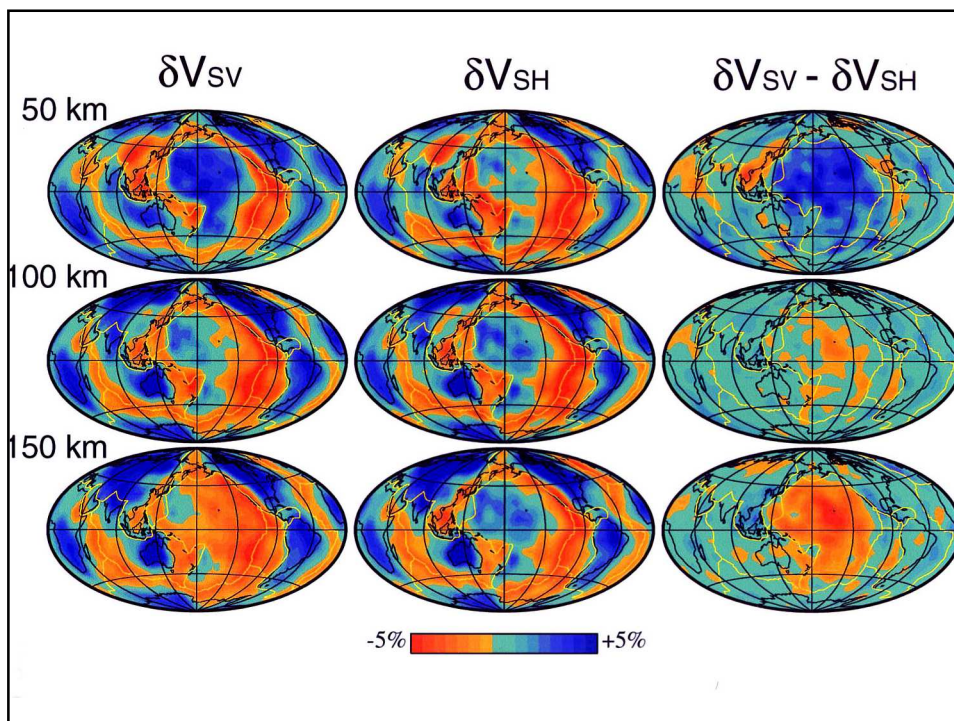
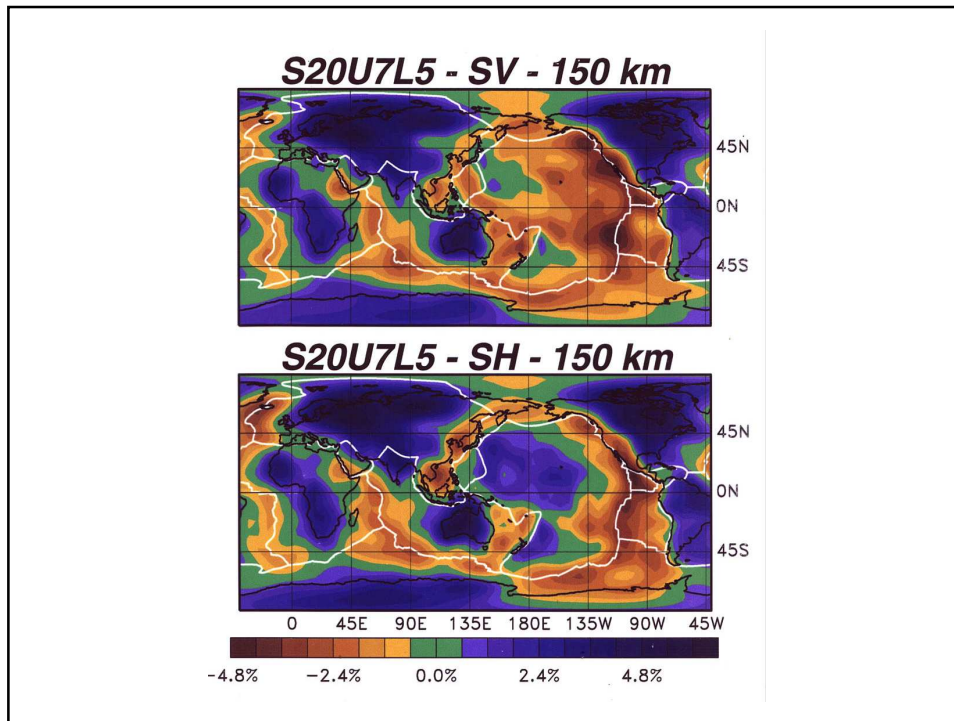


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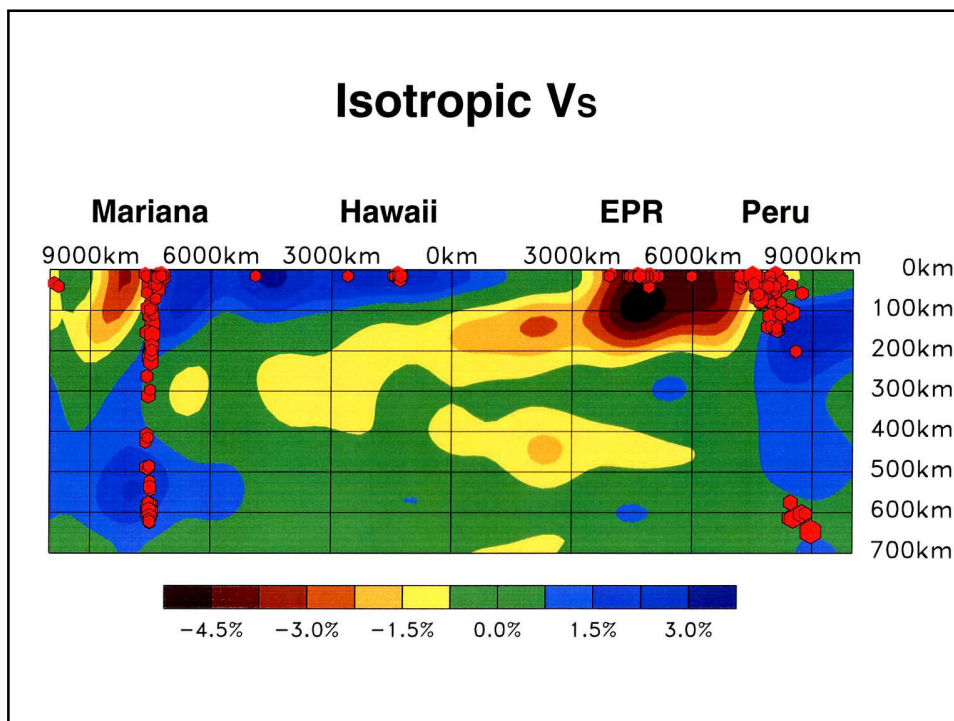
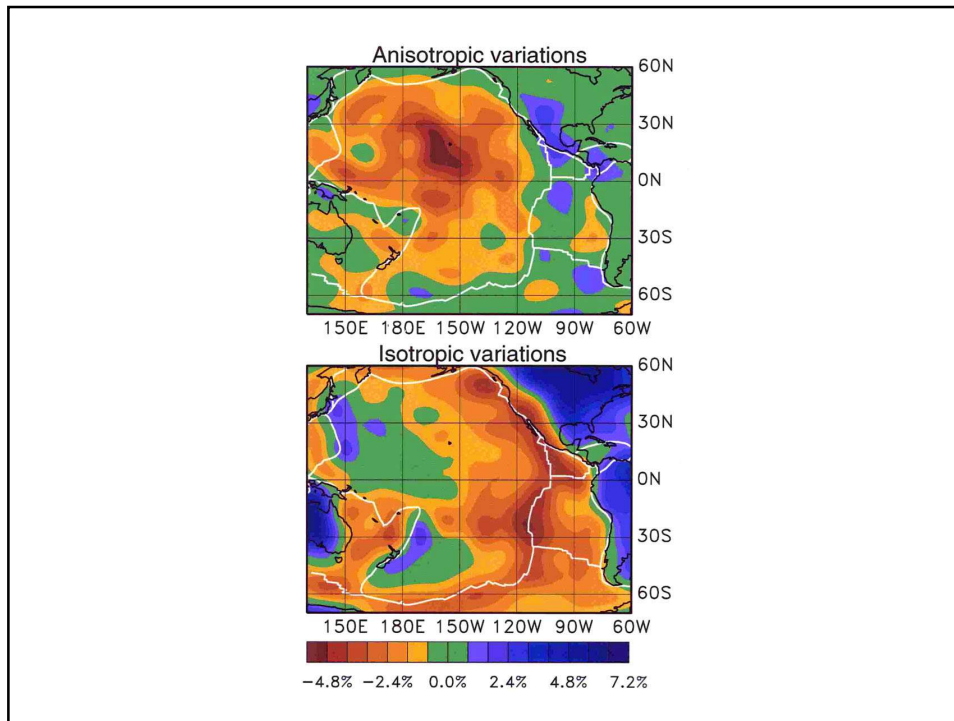


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Seismology: Surface waves: excitation, measurements of dispersion and attenuation

