

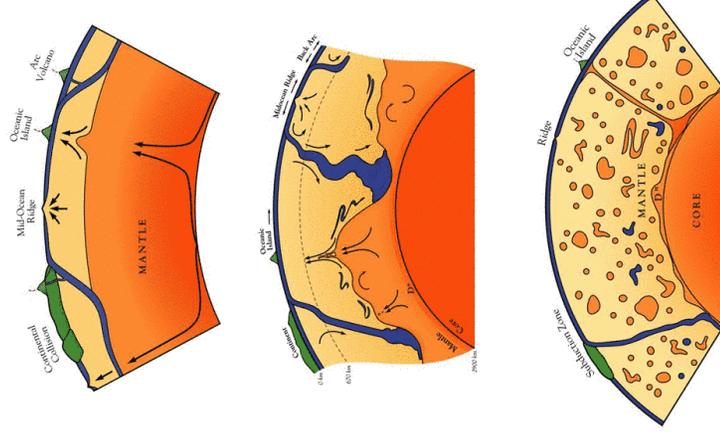
# Thermo-chemical convection Louise Kellogg, UC Davis

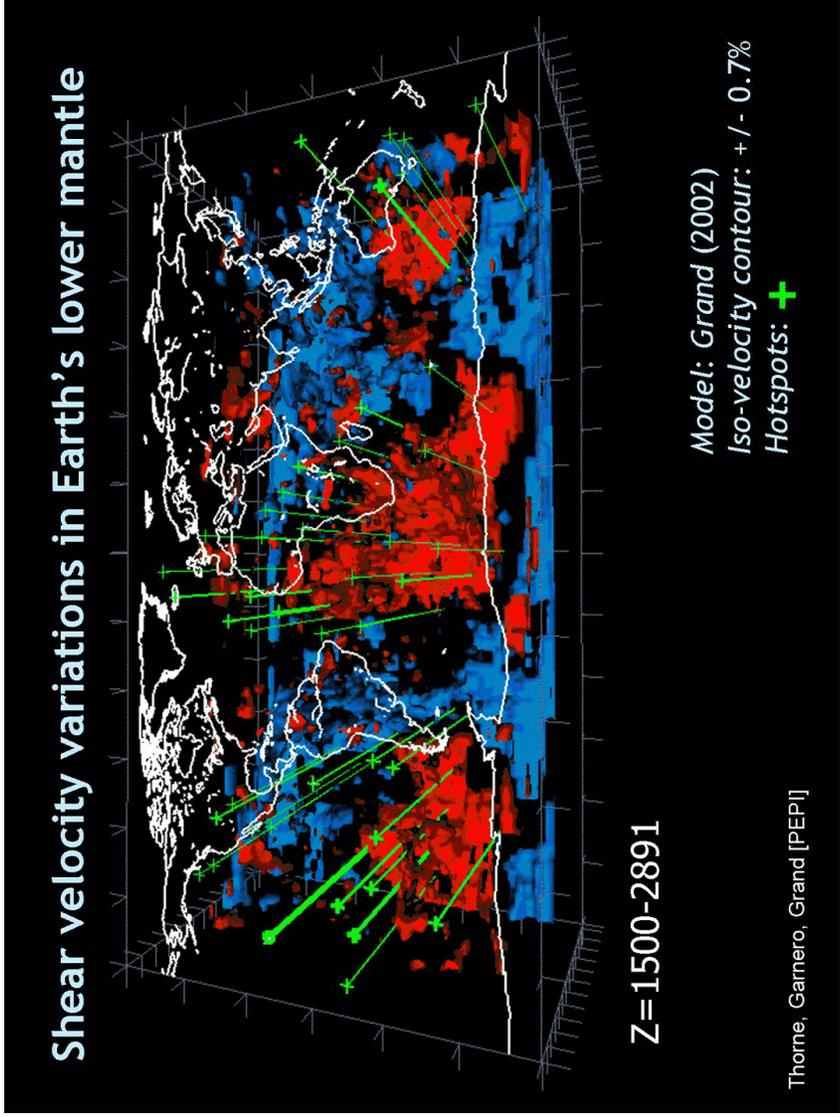
- Motivation: What is thermo-chemical convection, and how does it apply to the mantle?
- Methods: How do we study thermo-chemical convection?
- Examples of applications

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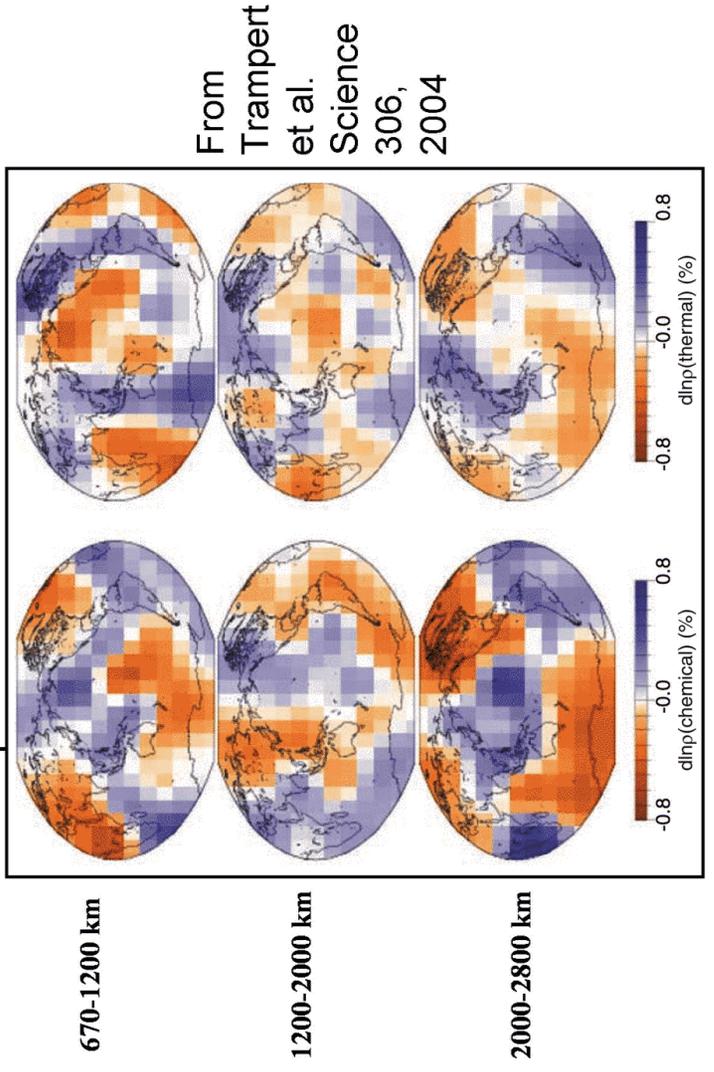
## Lines of argument for some compositional layering

- Geochemical (i.e. Ar, He, etc.)
- Heat production
- Seismic velocities in lowermost mantle
- Mineral physics (combined with seismic velocities)

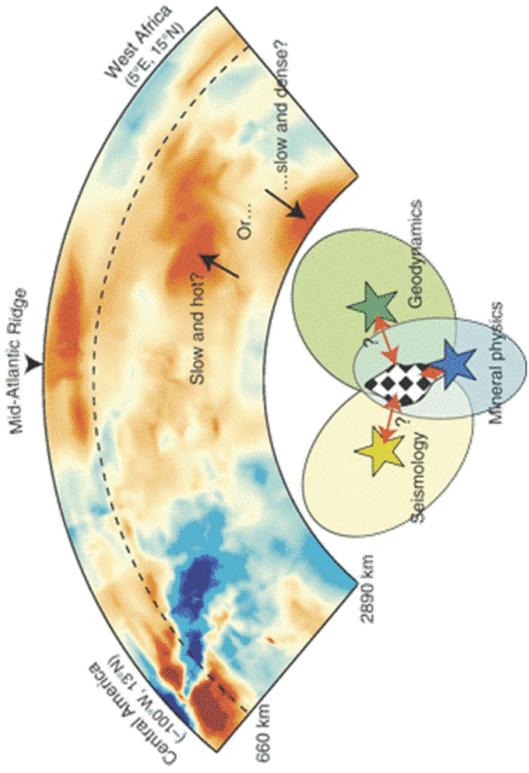




### Inferred compositional variations in the mantle



Do the seismic velocities show any evidence for a layer?



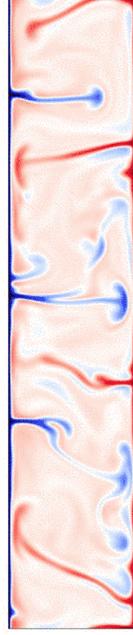
From R. van der Hilst

Mantle Convection: as seen last time

Conservation of mass:  $\nabla \cdot \vec{u} = 0$

Momentum:  $-\nabla P + \nabla \cdot \tau + Ra \hat{T}k$

Heat flow:  $\frac{\partial T}{\partial t} + \vec{u} \cdot \nabla T = \nabla^2 T + Ra_H / Ra$



$$Ra = \frac{\rho_m g \alpha \Delta T d^3}{\kappa \mu_0}$$

$$Ra_H = \frac{\rho_m^2 g \alpha H d^5}{\kappa \mu_0}$$

Mantle Convection with chemical buoyancy:

Conservation of mass:  $\nabla \cdot \vec{u} = 0$

Momentum:  $-\nabla P + \nabla \cdot \tau + RaT\hat{k} - BRaC\hat{k} = 0$

Heat flow:  $\frac{\partial T}{\partial t} + \vec{u} \cdot \nabla T = \nabla^2 T + Ra_H / Ra$

Advection/Diffusion:  $\frac{\partial C}{\partial t} + \vec{u} \cdot \nabla C = \frac{1}{Le} \nabla^2 C$

$Ra = \frac{\rho_m g \alpha \Delta T d^3}{\kappa \mu_0}$

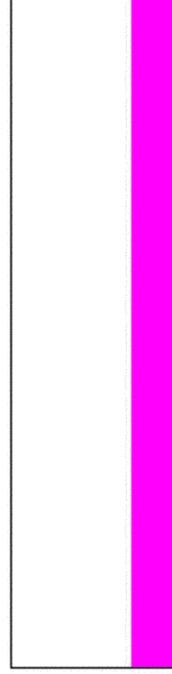
$Ra_H = \frac{\rho_m^2 g \alpha H d^5}{\kappa \mu_0}$

Lewis number:  $Le = \kappa / D$

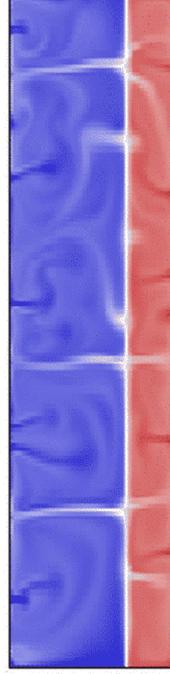
Compositional Buoyancy:

$B = \frac{\Delta \rho_c}{\rho_0 \alpha \Delta T}$

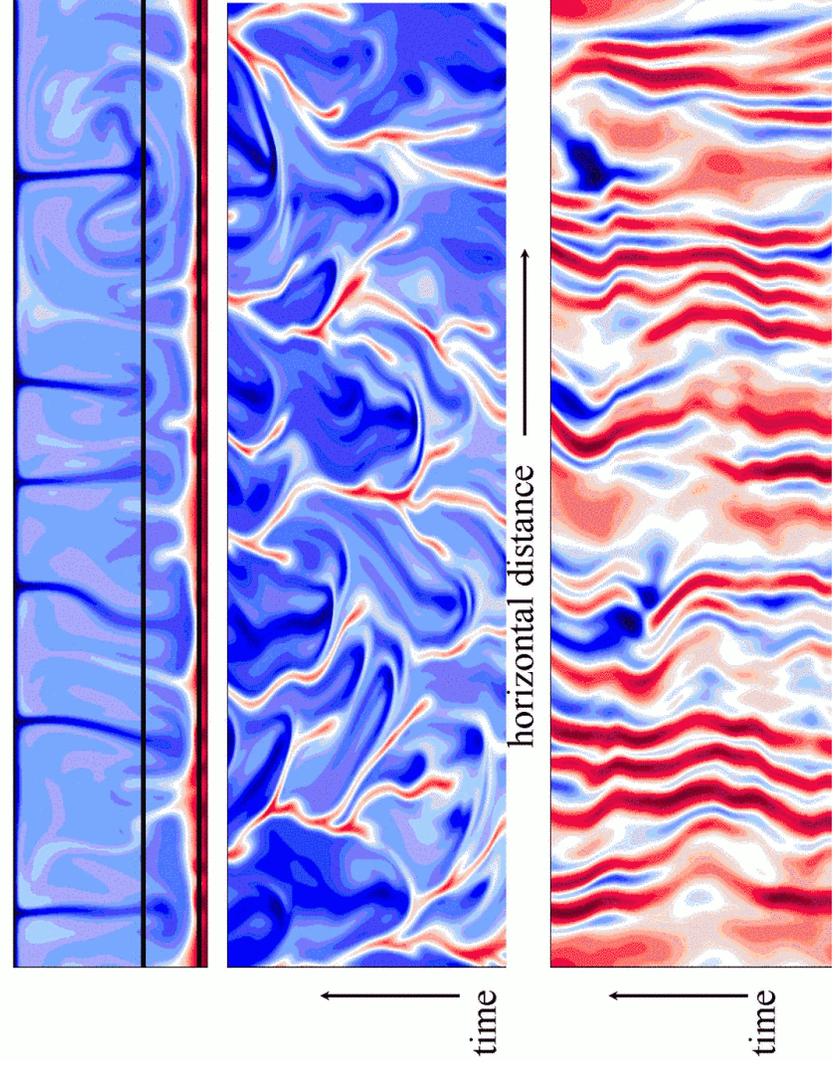
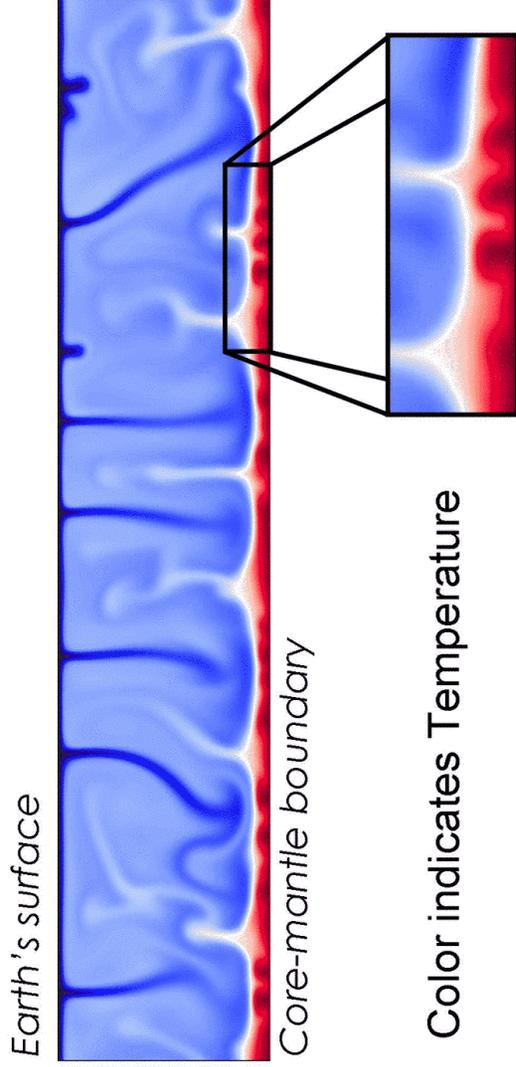
Initial Composition



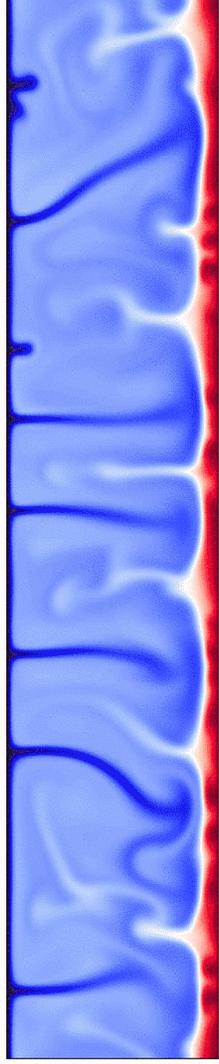
Initial temperature



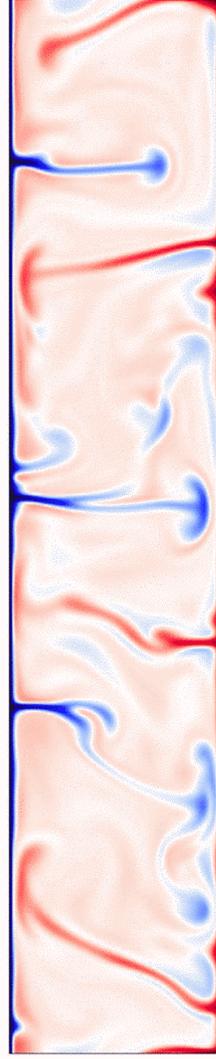
Double Diffusive Convection Model of D''  
 $B=1$   $Ra = 10^7$



Double Diffusive Convection Model of D"  
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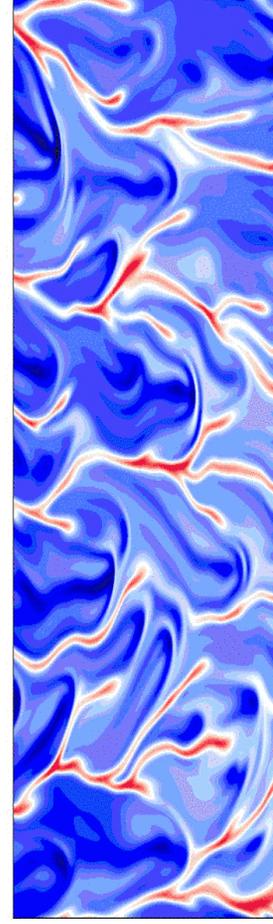


"Vanilla" Convection  
 $B=0$  (no composition)

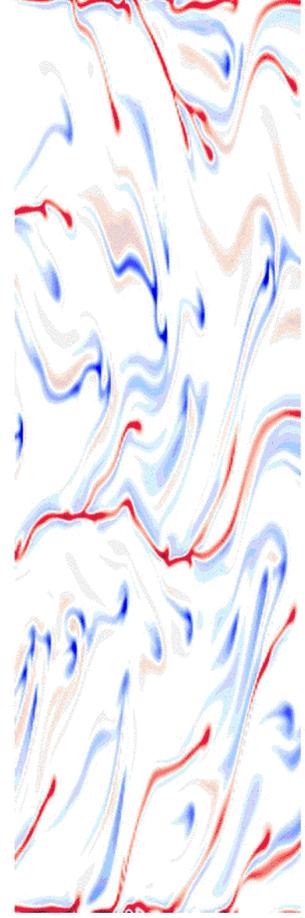


N. Montague and  
L. Kellogg, JGR,  
2000

A dense layer stabilizes the flow  
With a dense layer in D"



No dense layer

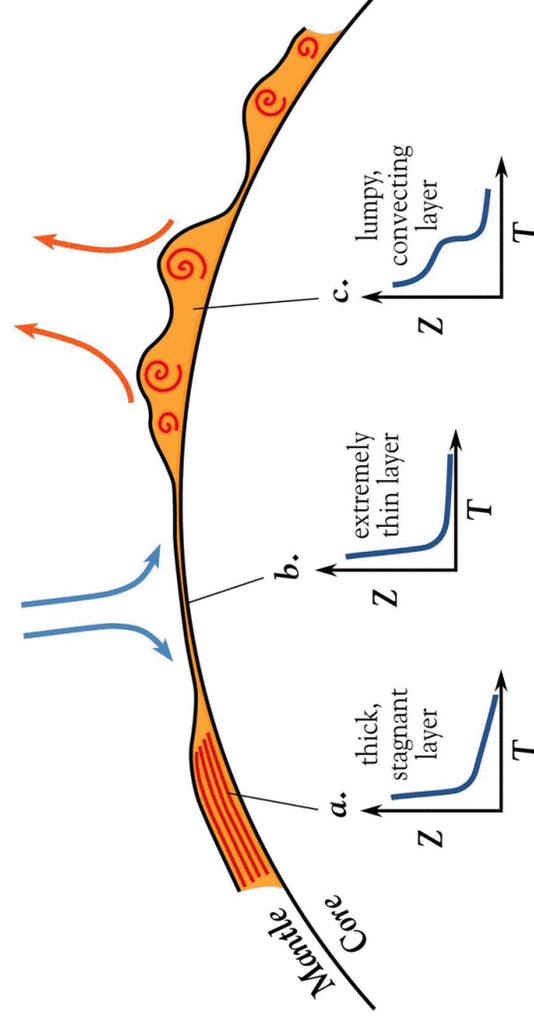


←  
time



Jellinek and Manga [Rev Geophys., 2005]

Possible processes in lowermost mantle



Mantle Convection with chemical buoyancy:

Conservation of mass:  $\nabla \cdot \vec{u} = 0$

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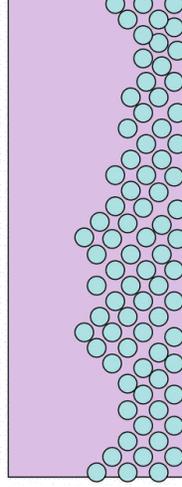
Lewis number:  $Le = \kappa / D$

Compositional Buoyancy:

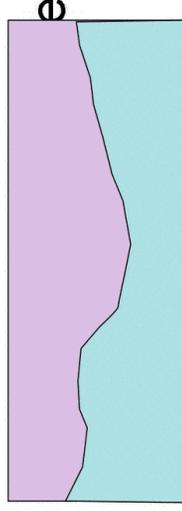
$B = \frac{\Delta \rho_c}{\rho_0 \alpha \Delta T}$

Some ways to represent thermo-chemical convection

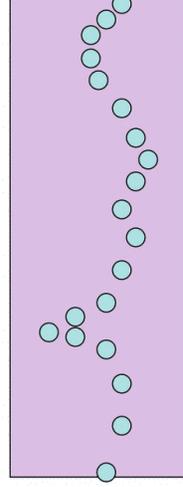
- Particle method



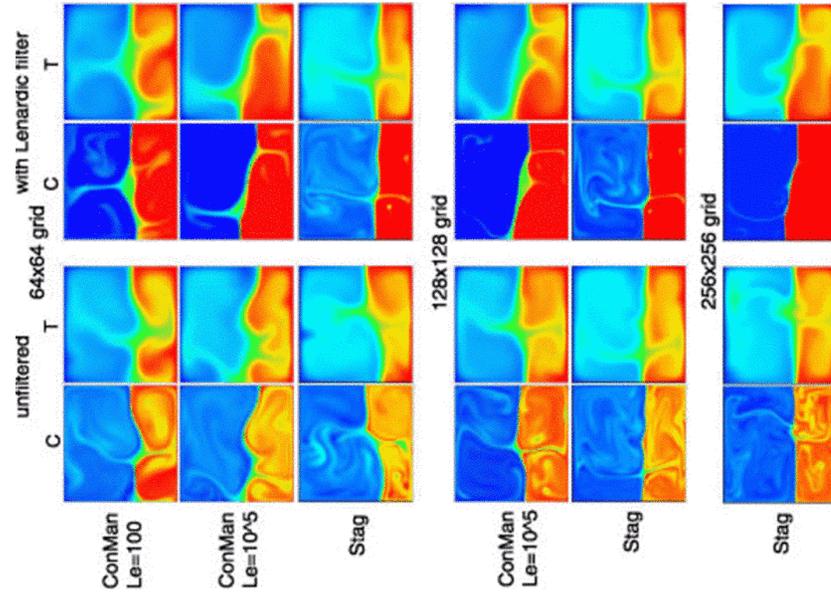
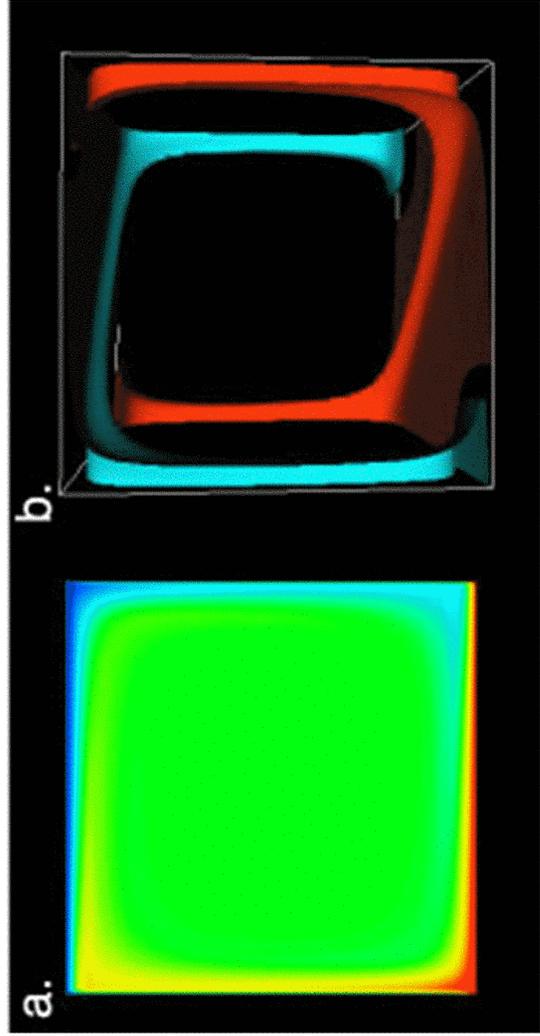
- Field method  
diffusive



- Marker chain



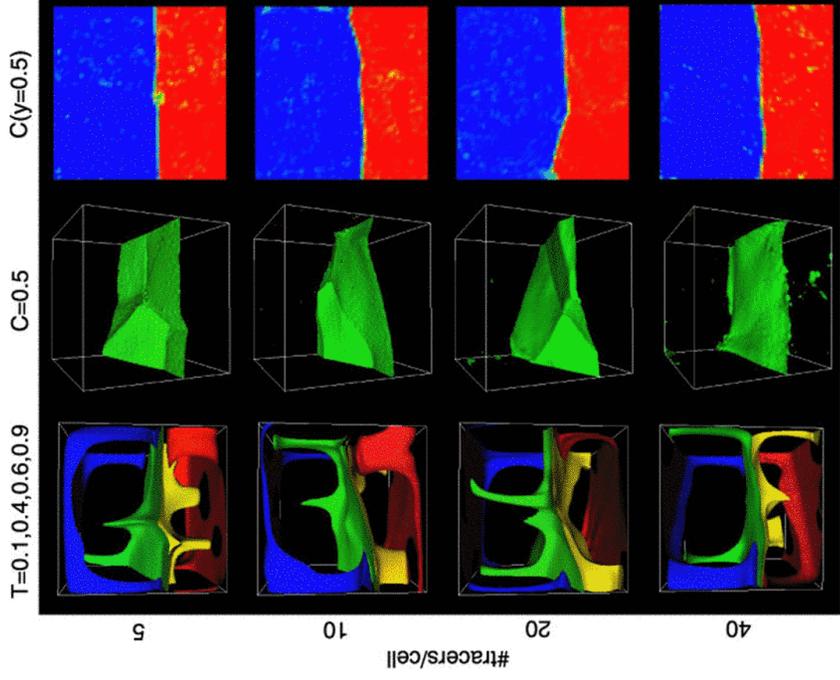
Testing methods for thermochemical convection  
(From Tackley & King, G-cubed, 2003)



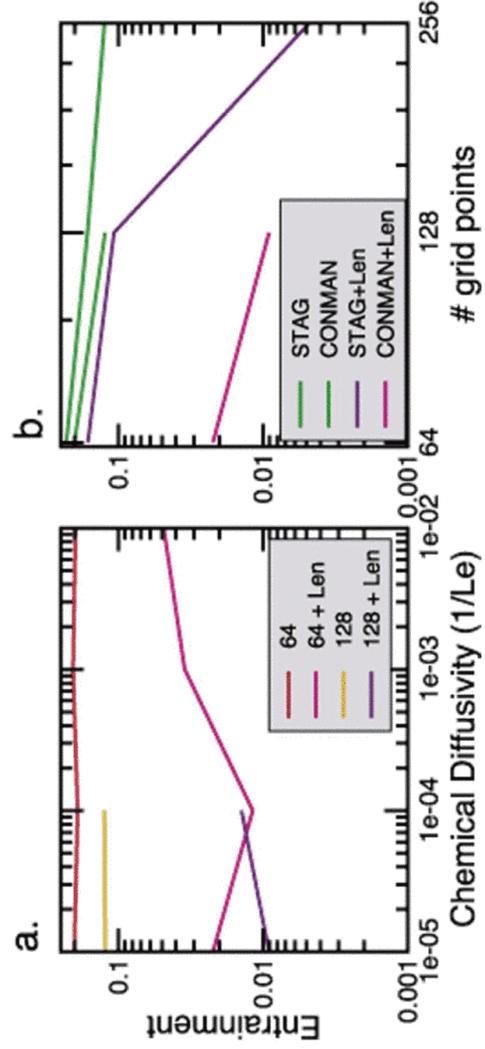
What's the best way  
to represent  
compositional  
heterogeneity in  
numerical models?

Tackley and King, 2003

3D tests using a particle method

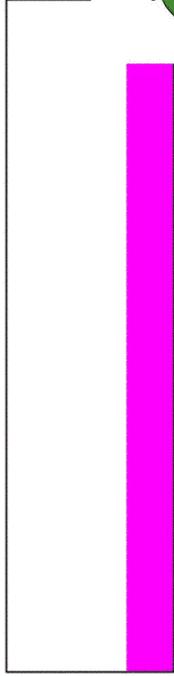


Comparisons of methods

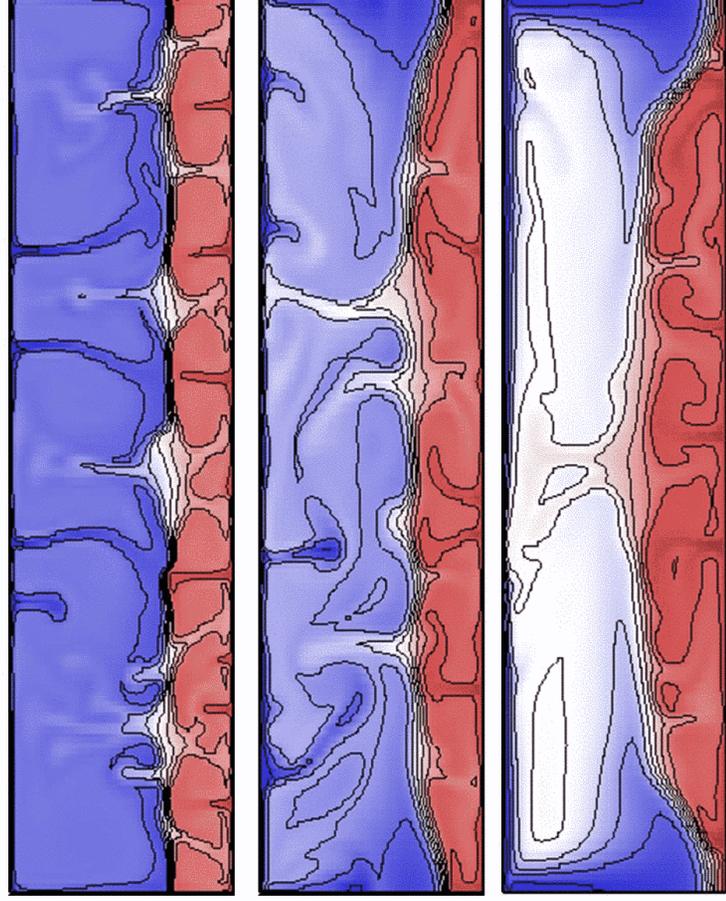
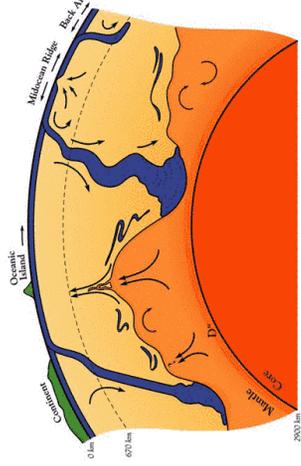
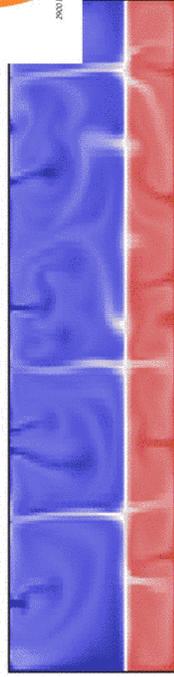


Other properties in conjunction with density: Effect of viscosity

Initial Composition

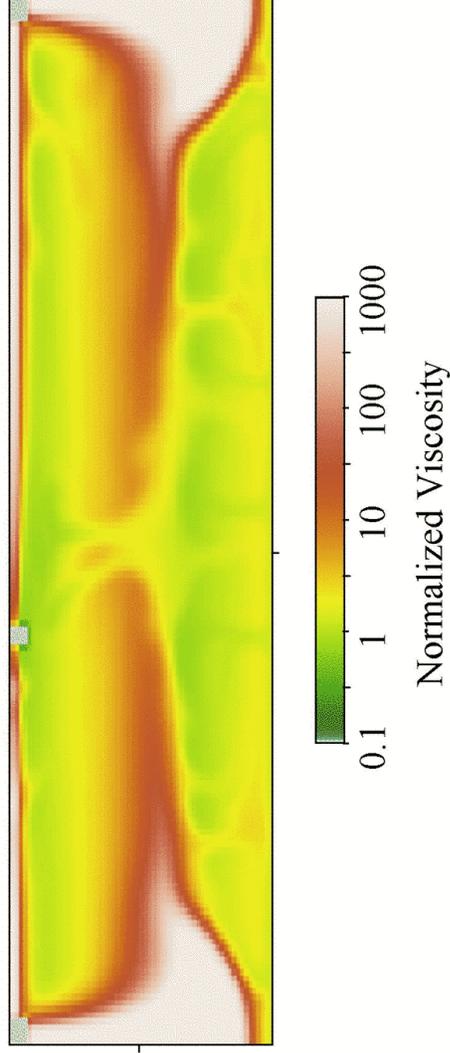


Initial temperature

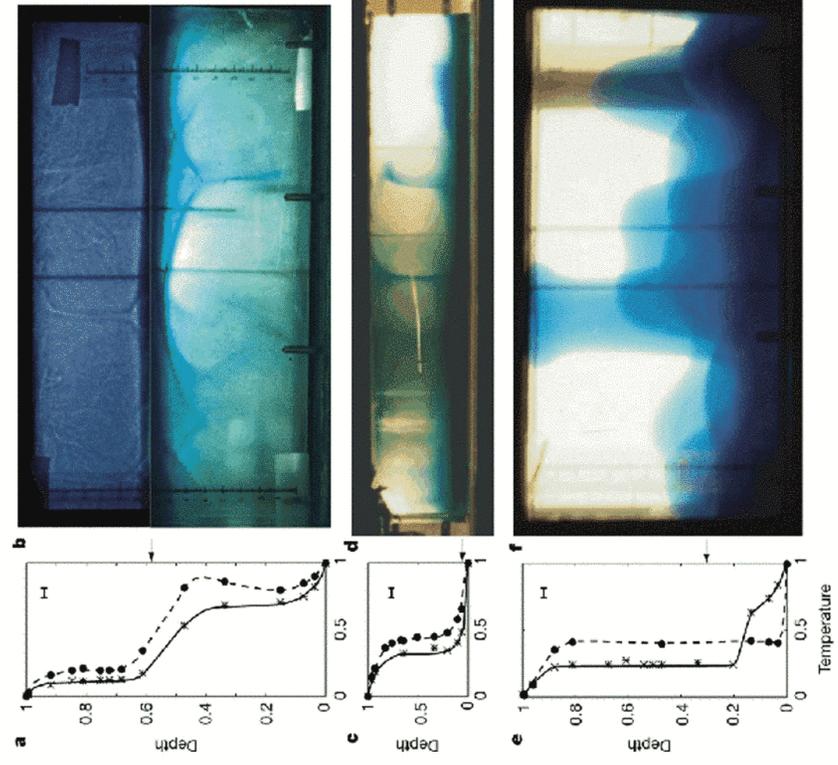


$B = 1$   →  
More temperature-dependent viscosity

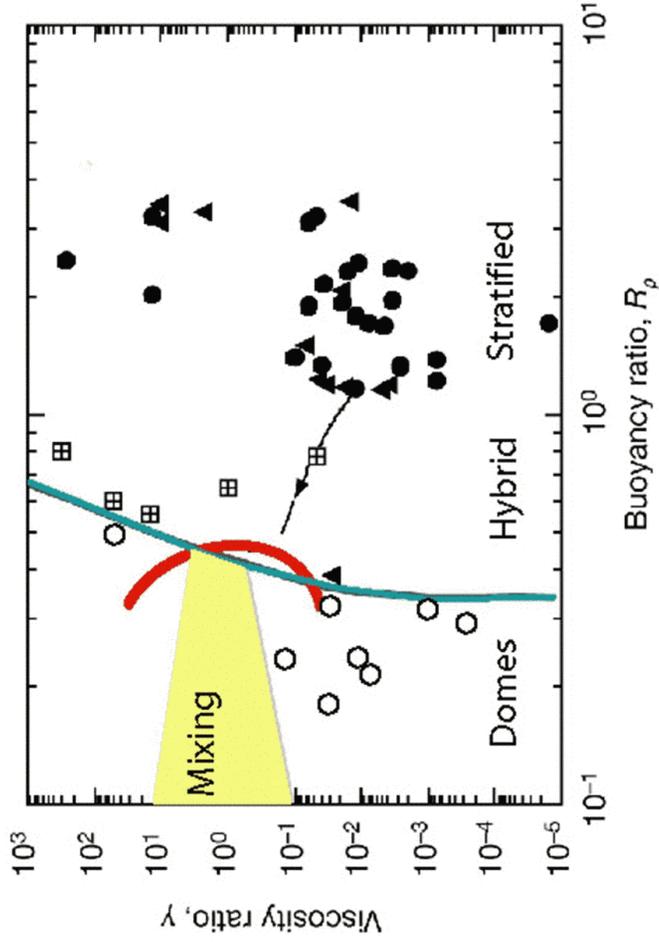
A viscosity inversion develops at the top of the layer



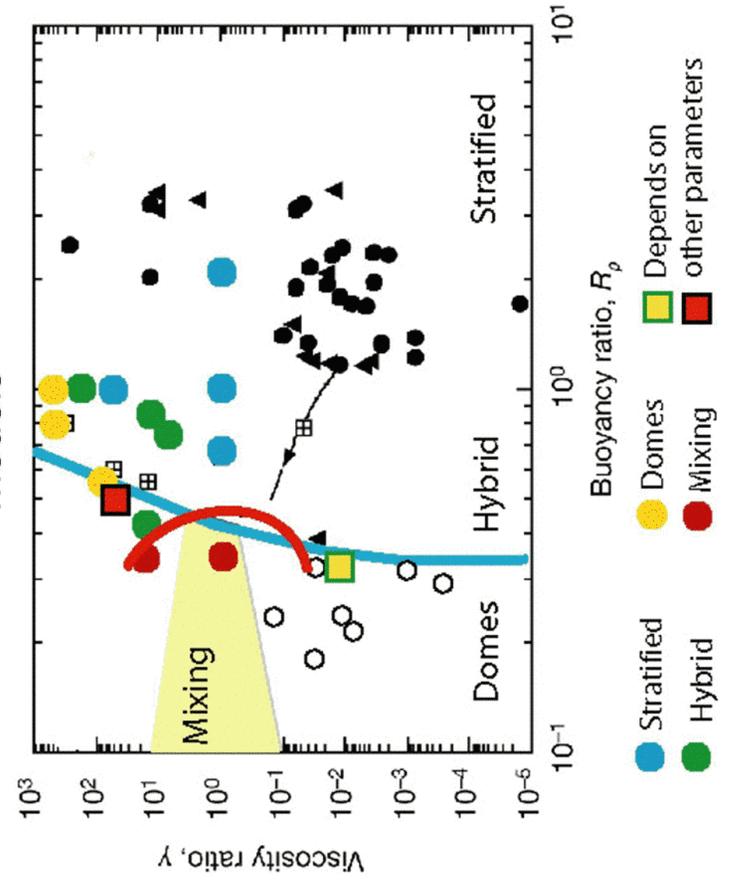
Layered convection experiments by Anne Davaille, (Nature 402, 756, Dec. 1999)



Davaille experiments

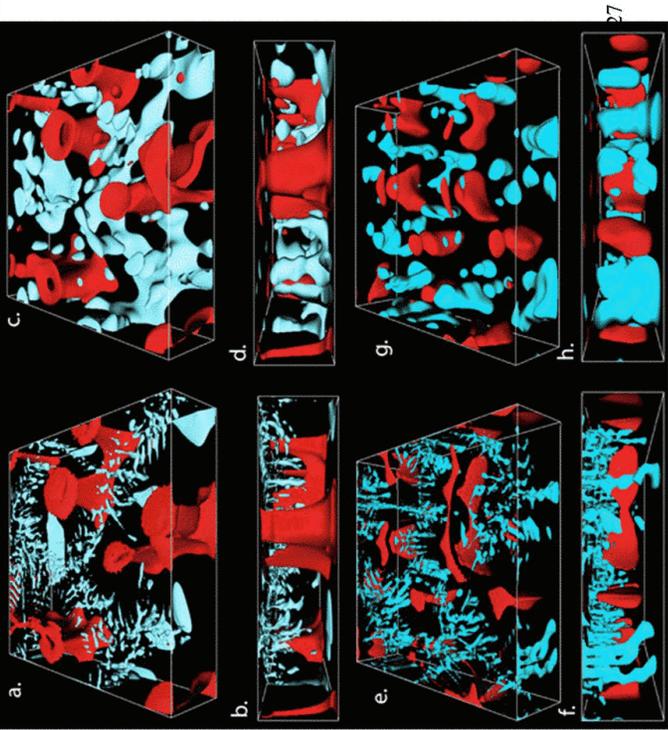


Davaille experiments + several numerical models



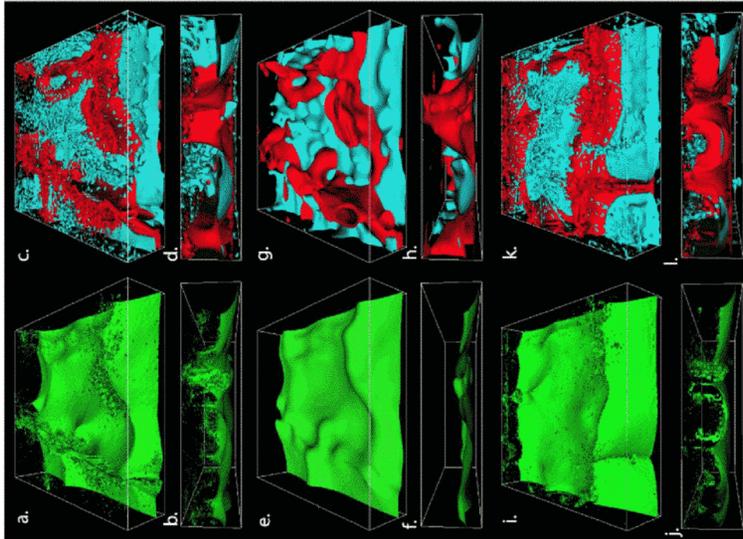
Fate of a dense layer (from Tackley, G-cubed, 2002)

Fixed T CMB



0 flux CMB

Next add a dense layer (30% of mantle)



Fixed T CMB  
B=0.3

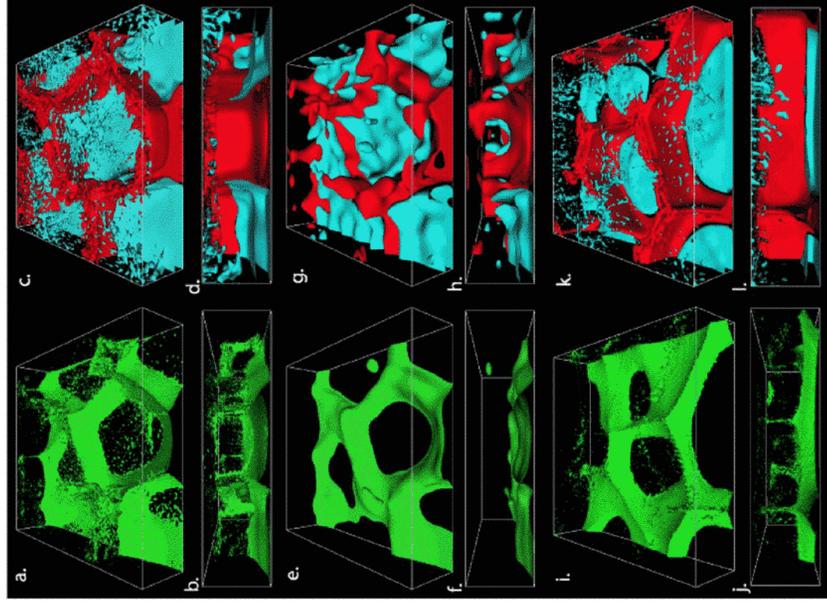
0 flux CMB

Tackley, G-cubed, 2002

A thinner dense layer can be discontinuous (10% of mantle)

Fixed T CMB  
 $B = 0.3$

0 flux CMB  
 $B = 0.25$



Tackley, G-cubed, 2002

Parameter space is many-dimensional!

