

Reduced Modeling of Strongly Stratified Shear Flows

Greg Chini[†] Guillaume Michel[‡]

Cesar Rocha[§] Keith Julien* Colm-Cille Caulfield[#]

[†]Program in Integrated Applied Mathematics & Department of Mechanical Engineering, University of New Hampshire

[‡]Laboratoire de Physique Statistique, Ecole Normale Supérieure, CNRS

[§]Scripps Institution of Oceanography, University of California San Diego

*Department of Applied Mathematics, University of Colorado, Boulder

[#]BP Institute, Dept. of Applied Mathematics & Theoretical Physics, Cambridge University

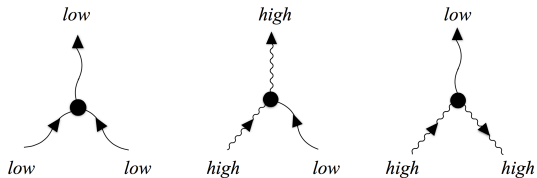
KITP18: Planetary Boundary Layers
June 14th, 2018

Themes

- 1 Stratified turbulence in absence of boundaries
- 2 Illustration of asymptotic derivation of QL system (temporal scale separation)
- 3 Extension of QL: slow spatial variation of “mean” field
⇒ Generalized Quasi-Linear (GQL) approx. [Marston, Chini & Tobias, *PRL* 2016]
- 4 Increasing efficiency of QL: full exploitation of temporal scale separation
⇒ New type of amplitude equation (which Guillaume will discuss. . .)

Quasi-Linear (QL) and Generalized Quasi-Linear (GQL) Formulations

- Decomposition into low and high modes in x and energy-conserving retention and removal of triadic interactions, w/cut-off Λ enabling homotopy b/w QL and DNS:

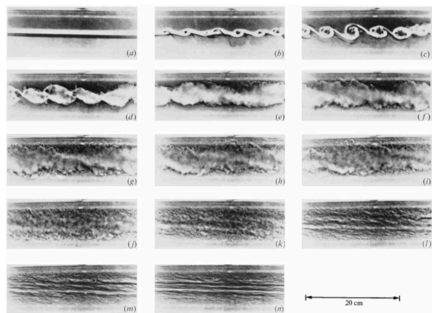


$$\text{Low modes: } \bar{u}(x, z, t) = \sum_{|k| \leq \Lambda} e^{ikx} u_k(z, t)$$

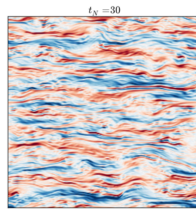
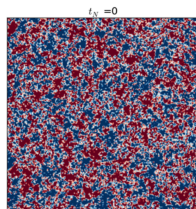
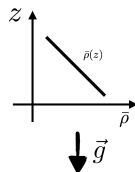
$$\text{High modes: } u'(x, z, t) = \sum_{|k| > \Lambda} e^{ikx} u_k(z, t)$$

- QL obtained for $\Lambda = 0$, while DNS (“NL”) obtained for $\Lambda \rightarrow \infty$
- Accuracy significantly improved for $0 < \Lambda \lesssim 5$
- Efficiency achieved by directly simulating *Statistical* Dynamical System (GCE2)?

Stratified Turbulence: Phenomenology & Significance



-Thorpe (1971)

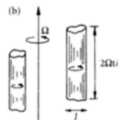
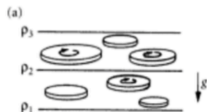
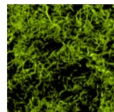


-1024x1024 DNS (C. Rocha)

- Anisotropic layers with $L \gg h$ coexist with small-scale isotropic features
- Mechanism for diabatic mixing (e.g. crucial for closing ocean circulation)
- SGS process in regional circulation and in computational climate models

Stratified Turbulence: Scales and Parameters

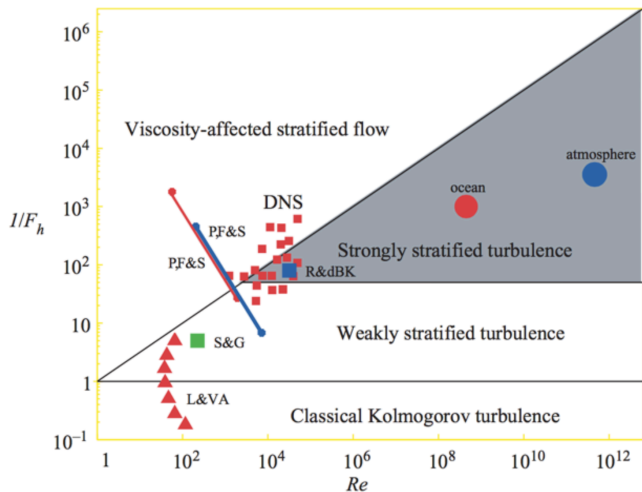
$$\text{Buoyancy frequency } N \equiv \sqrt{-\frac{g}{\rho_0} \frac{d\bar{\rho}}{dz}}$$

Rotating
turbulence $1/L$ **Stratified turbulence** $1/L_O$ 3D
turbulence $1/L_K$

- (Horizontal) Froude number: $Fr \equiv \frac{U}{NL}$
- Reynolds number: $Re \equiv \frac{UL}{\nu}$

Simple scaling arguments give $L/L_O = O(Fr^{-3/2})$, where $L_O = (\epsilon/N^3)^{1/2}$

Stratified Turbulence: Computational Challenge

-Brethouwer *et al.* (2007)Note: $F_h \equiv Fr$

Stratified Turbulence: Fundamental Questions

- What sets vertical scale(s)?
- **Mixing properties?**
- Is horizontal spectrum of horizontal KE independent of Fr as $Fr \rightarrow 0$ for $Re \gtrsim 10/Fr^2$?

*Bartello & Tobias (2013) estimate that to demonstrate this independence even over one decade in Fr , namely, over the parameter range $0.01 \leq Fr \leq 0.1$, would require the ratio of maximum to minimum resolvable scale to be in the **millions** (i.e. in a single spatial direction), yielding a formidable computational challenge.*

Stratified Turbulence: Governing Equations

Anisotropic Scaling

$$\mathbf{x}_\perp : L \quad z : h \quad t : L/U \quad \mathbf{u}_\perp : U \quad w : Fr^2 UL/h \quad p : \rho_0 U^2 \quad b : U^2/h$$

Non-Rotating Boussinesq Equations

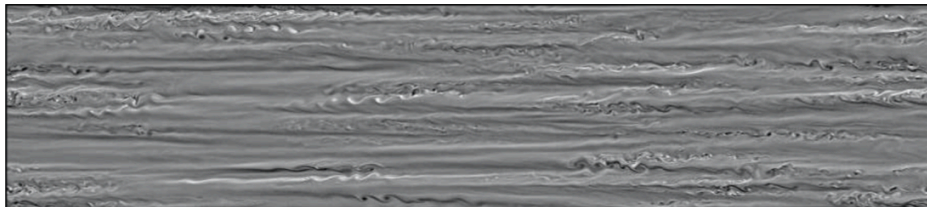
$$\begin{aligned} \partial_t \mathbf{u}_\perp + (\mathbf{u}_\perp \cdot \nabla_\perp) \mathbf{u}_\perp + \frac{Fr^2}{\alpha^2} W \partial_z \mathbf{u}_\perp &= -\nabla_\perp p + \mathcal{D}(\mathbf{u}_\perp) + \mathbf{f}_\perp \\ Fr^2 \left[\partial_t W + (\mathbf{u}_\perp \cdot \nabla_\perp) W + \frac{Fr^2}{\alpha^2} W \partial_z W \right] &= -\partial_z p + b + Fr^2 \mathcal{D}(W) \\ \nabla_\perp \cdot \mathbf{u}_\perp + \frac{Fr^2}{\alpha^2} \partial_z W &= 0 \\ \partial_t b + (\mathbf{u}_\perp \cdot \nabla_\perp) b + \frac{Fr^2}{\alpha^2} W \partial_z b &= -W + Pr^{-1} \mathcal{D}(b) \end{aligned}$$

where: $\mathcal{D} = \frac{1}{Re} \left[\nabla_\perp^2 + \frac{1}{\alpha^2} \partial_z^2 \right]$ and $\alpha \equiv h/L$, $Fr \equiv U/(NL)$

Limit Equations

- $Fr/\alpha \rightarrow 0$ as $Fr \rightarrow 0$ (Lilly 1983): Layerwise **2D** flow
- $Fr/\alpha = O(1)$ as $Fr \rightarrow 0$ (Billant & Chomaz 2001): Anisotropic **3D** flow

Emergence of Multiple Vertical and Horizontal Scales



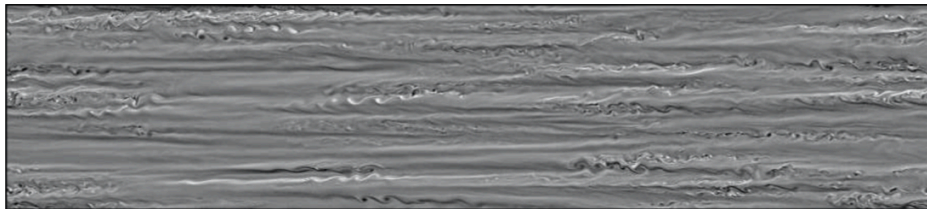
–Waite (2014)

- As $Fr \rightarrow 0$, vertical length scale self-adjusts so that $h = O(U/N) \Rightarrow \alpha = O(Fr)$
[Note: then $w : hU/L$]
- Stratified turbulence regime *defined* by $Fr \rightarrow 0$ with:

$\alpha = O(Fr)$ and **buoyancy Reynolds number** $\mathcal{R} \equiv ReFr^2 \gtrsim 10$
- Clear evidence of shear instabilities on horizontal scales $l \ll L$, and perhaps modulation on $O(L)$ vertical scales

... *Opportunity for asymptotically-reduced multiscale modeling...*

Multiple Scale Asymptotic Analysis



-Waite (2014)

- 1 Identify relevant **distinguished limit**: $\alpha = Fr$ and $\mathcal{R} \equiv ReFr^2 = O(1)$ as $Fr \rightarrow 0$
- 2 Introduce *fast* horizontal and temporal scales: $\chi_{\perp} = \mathbf{x}_{\perp}/Fr$ and $\tau = t/Fr$ so that

$$\nabla_{\perp} \rightarrow \nabla_{\mathbf{x}} + \frac{1}{Fr} \nabla_{\chi} \quad \partial_t \rightarrow \partial_t + \frac{1}{Fr} \partial_{\tau}$$

- 3 Introduce fast averaging operation and mean/fluctuation decomposition:

$$\phi(\mathbf{x}_{\perp}, z, t) \rightarrow \phi(\chi_{\perp}, \mathbf{x}_{\perp}, z, \tau, t) = \bar{\phi}(\mathbf{x}_{\perp}, z, t) + \phi'(\chi_{\perp}, \mathbf{x}_{\perp}, z, \tau, t), \text{ where } \bar{\phi}' \equiv 0$$

Expansion *Ansatz*

- Introduce $\epsilon \equiv \sqrt{Fr}$ and posit following asymptotic expansions for various fields:

$$\begin{aligned} [\mathbf{u}_\perp, b, p] &\sim [\mathbf{u}_{0\perp}, b_0, p_0] + \epsilon[\mathbf{u}_{1\perp}, b_1, p_1] + \epsilon^2[\mathbf{u}_{2\perp}, b_2, p_2] + \dots \\ W &\sim \epsilon^{-1}W_{-1} + W_0 + \epsilon W_1 + \dots \end{aligned}$$

- Key prescription is that vertical velocity (normalized by αU) is $O(\epsilon^{-1})$ on fine horizontal scales (i.e. when scaled by U , vertical velocity $w \sim \epsilon w_1 + \epsilon^2 w_2 + \dots$)
- This (re-)scaling ensures that feedback of fluctuations upon mean fields through vertical Reynolds stress divergence $\partial_z \left[\overline{W' u'_\perp} \right]$ arises at proper order s.t. there is a dominant balance with tendency $\partial_t \bar{\mathbf{u}}_\perp$ and vertical diffusion $\mathcal{R}^{-1} \partial_z^2 \bar{\mathbf{u}}_\perp$

Rescaling simultaneously ensures that fine-scale dynamics are **isotropic**

Compatible with DNS/scaling analysis showing $\langle w^2 \rangle = Fr U^2$ at **small scales** (Maffioli & Davidson, JFM 2016)

- Can then *deduce* that **fluctuating** horizontal velocity, buoyancy, and pressure fields arise at $O(\epsilon)$, a key simplification

Multiscale Reduced PDEs

Mean Equations

$$\begin{aligned}
 \left[\partial_t + (\bar{\mathbf{u}}_{0\perp} \cdot \nabla_{\mathbf{x}}) + \bar{W}_0 \partial_z \right] \bar{\mathbf{u}}_{0\perp} + \partial_z \left(\overline{W'_{-1} \mathbf{u}'_{1\perp}} \right) &= -\nabla_{\mathbf{x}} \bar{p}_0 + \frac{1}{\mathcal{R}} \partial_z^2 \bar{\mathbf{u}}_{0\perp} + \bar{\mathbf{f}}_{0\perp} \\
 0 &= -\partial_z \bar{p}_0 + \bar{b}_0 \\
 \nabla_{\mathbf{x}} \cdot \bar{\mathbf{u}}_{0\perp} + \partial_z \bar{W}_0 &= 0 \\
 \left[\partial_t + (\bar{\mathbf{u}}_{0\perp} \cdot \nabla_{\mathbf{x}}) + \bar{W}_0 \partial_z \right] \bar{b}_0 + \partial_z \left(\overline{W'_{-1} b'_1} \right) &= -\bar{W}_0 + \frac{1}{Pr\mathcal{R}} \partial_z^2 \bar{b}_0
 \end{aligned}$$

Multiscale Reduced PDEs

Mean Equations

$$\begin{aligned} \left[\partial_t + (\bar{\mathbf{u}}_{0\perp} \cdot \nabla_{\mathbf{x}}) + \bar{W}_0 \partial_z \right] \bar{\mathbf{u}}_{0\perp} + \partial_z \left(\overline{W'_{-1} \mathbf{u}'_{1\perp}} \right) &= -\nabla_{\mathbf{x}} \bar{p}_0 + \frac{1}{\mathcal{R}} \partial_z^2 \bar{\mathbf{u}}_{0\perp} + \bar{\mathbf{f}}_{0\perp} \\ 0 &= -\partial_z \bar{p}_0 + \bar{b}_0 \\ \nabla_{\mathbf{x}} \cdot \bar{\mathbf{u}}_{0\perp} + \partial_z \bar{W}_0 &= 0 \\ \left[\partial_t + (\bar{\mathbf{u}}_{0\perp} \cdot \nabla_{\mathbf{x}}) + \bar{W}_0 \partial_z \right] \bar{b}_0 + \partial_z \left(\overline{W'_{-1} b'_1} \right) &= -\bar{W}_0 + \frac{1}{Pr\mathcal{R}} \partial_z^2 \bar{b}_0 \end{aligned}$$

Fluctuation Equations

$$\begin{aligned} (\partial_\tau + \bar{\mathbf{u}}_{0\perp} \cdot \nabla_{\mathbf{x}}) \mathbf{u}'_{1\perp} + W'_{-1} \partial_z \bar{\mathbf{u}}_{0\perp} &= -\nabla_{\mathbf{x}} p'_1 + \frac{Fr}{\mathcal{R}} \left(\nabla_{\mathbf{x}}^2 + \partial_z^2 \right) \mathbf{u}'_{1\perp} \\ (\partial_\tau + \bar{\mathbf{u}}_{0\perp} \cdot \nabla_{\mathbf{x}}) W'_{-1} &= -\partial_z p'_1 + b'_1 + \frac{Fr}{\mathcal{R}} \left(\nabla_{\mathbf{x}}^2 + \partial_z^2 \right) b'_1 \\ \nabla_{\mathbf{x}} \cdot \mathbf{u}'_{1\perp} + \partial_z W'_{-1} &= 0 \\ (\partial_\tau + \bar{\mathbf{u}}_{0\perp} \cdot \nabla_{\mathbf{x}}) b'_1 + W'_{-1} \partial_z \bar{b}_0 &= -W'_{-1} + \frac{Fr}{Pr\mathcal{R}} \left(\nabla_{\mathbf{x}}^2 + \partial_z^2 \right) b'_1 \end{aligned}$$

Attributes of Multiscale System

- In absence of Reynolds stress (“eddy-flux”) divergences (RSDs), mean equations reduce to hydrostatic primitive equations (Billant & Chomaz 2001)
- Vertical RSDs provide crucial feedback of fluctuating fine-scale dynamics on evolution of mean fields, given here without need for *ad hoc* closure
- Fluctuation dynamics are **quasi-linear (QL)** about local mean fields:
 - Reverting to single set of (\mathbf{x}_\perp, t) scales yields [and interpreting $\overline{(\cdot)}$ as strict horizontal mean] yields a **QL reduction** of full Boussinesq equations
 - Suggests **2nd-order cumulant expansion (CE2)** approaches used by Marston & Tobias, Farrell & Ioannou, and Young & Srinivasan can be **formally justified** for stratified shear turbulence in the limit $Fr \rightarrow 0$
- By retaining multiple horizontal and temporal scales, can rationally extend QL/CE2 schemes \Rightarrow **Physical space variant of spectral GQL**

Strict QL Model

- Re-interpreting $\overline{(\cdot)}$ as strict horizontal mean
- Single horizontal scale χ and revert to single time scale τ via $\partial_t \rightarrow (1/Fr)\partial_\tau$

Mean Equations

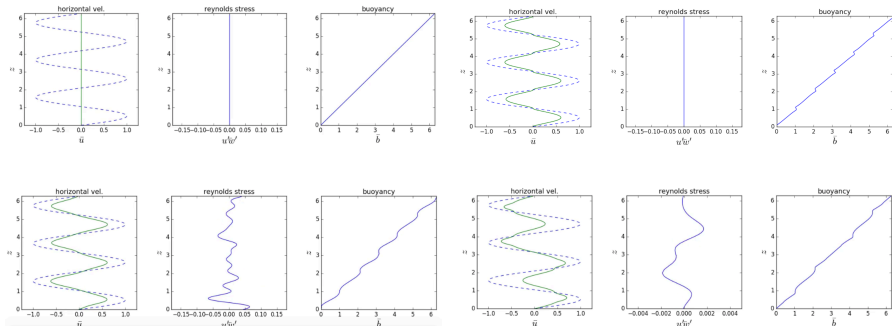
$$\begin{aligned} \frac{1}{Fr} \partial_\tau \bar{\mathbf{u}}_{0\perp} + \partial_z \left(\overline{W'_{-1} \mathbf{u}'_{1\perp}} \right) &= \frac{1}{\mathcal{R}} \partial_z^2 \bar{\mathbf{u}}_{0\perp} + \bar{\mathbf{f}}_{0\perp} \\ 0 &= -\partial_z \bar{p}_0 + \bar{b}_0 \\ \frac{1}{Fr} \partial_\tau \bar{b}_0 + \partial_z \left(\overline{W'_{-1} b'_1} \right) &= \frac{1}{Pr\mathcal{R}} \partial_z^2 \bar{b}_0 \end{aligned}$$

Fluctuation Equations

$$\begin{aligned} (\partial_\tau + \bar{\mathbf{u}}_{0\perp} \cdot \nabla_\chi) \mathbf{u}'_{1\perp} + W'_{-1} \partial_z \bar{\mathbf{u}}_{0\perp} &= -\nabla_\chi p'_1 + \frac{Fr}{\mathcal{R}} \left(\nabla_\chi^2 + \partial_z^2 \right) \mathbf{u}'_{1\perp} \\ (\partial_\tau + \bar{\mathbf{u}}_{0\perp} \cdot \nabla_\chi) W'_{-1} &= -\partial_z p'_1 + b'_1 + \frac{Fr}{\mathcal{R}} \left(\nabla_\chi^2 + \partial_z^2 \right) b'_1 \\ \nabla_\chi \cdot \mathbf{u}'_{1\perp} + \partial_z W'_{-1} &= 0 \\ (\partial_\tau + \bar{\mathbf{u}}_{0\perp} \cdot \nabla_\chi) b'_1 + W'_{-1} \partial_z \bar{b}_0 &= -W'_{-1} + \frac{Fr}{Pr\mathcal{R}} \left(\nabla_\chi^2 + \partial_z^2 \right) b'_1 \end{aligned}$$

Evolution of Mean Fields for (2D) Strongly Stratified Kolmogorov Flow

- Body forcing $\bar{f}_{0x} = (m^2/\mathcal{R}) \sin(mz)$ and imposed (initially) linear stratification
- Parameters: $m = 3$, $\mathcal{R} = 200$, $Fr = 0.02$



Extending QL Framework for Increased Accuracy and Efficiency

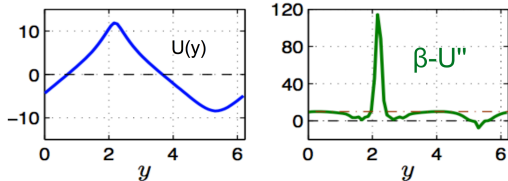
Some observations. . .

- QL simulations can be as expensive as DNS (!) owing to temporal stiffness and inclusion of many fluctuation (zonally-varying) modes, particularly in large domains
- Generalized QL (GQL) formulation has been shown to increase accuracy, but not necessarily efficiency
- Statistical formulations of (G)QL dynamics [(G)CE2] may offer certain efficiencies but increases spatial dimensionality. . .

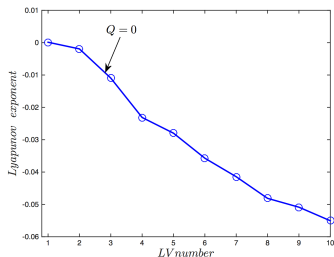
Alternative approach for systematically extending QL reduction and increasing computational efficiency relies on 3 ingredients:

- 1 Self-tuning of QL systems toward **“statistical marginal stability”**
- 2 Evolution of QL fluctuation energy spectrum toward **narrow-banded distribution**
- 3 Exploitation of **multiple scales** in t , x (and possibly z)

Statistical Marginal Stability of QL/CE2 Systems

DSS of Barotropic Jets

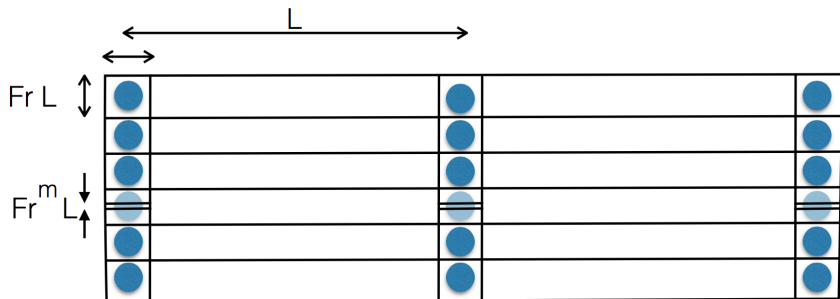
N. Constantinou. Formation of large-scale structures by turbulence in rotating planets. arXiv:1503.07644, 2015.

DSS of Plane Couette Flow

$$\partial_t \mathbf{U} = \mathbf{P}_L \left(-\mathbf{U} \cdot \nabla \mathbf{U} + \frac{1}{R} \Delta \mathbf{U} \right) + \mathcal{L}_k \mathbf{C}_k,$$

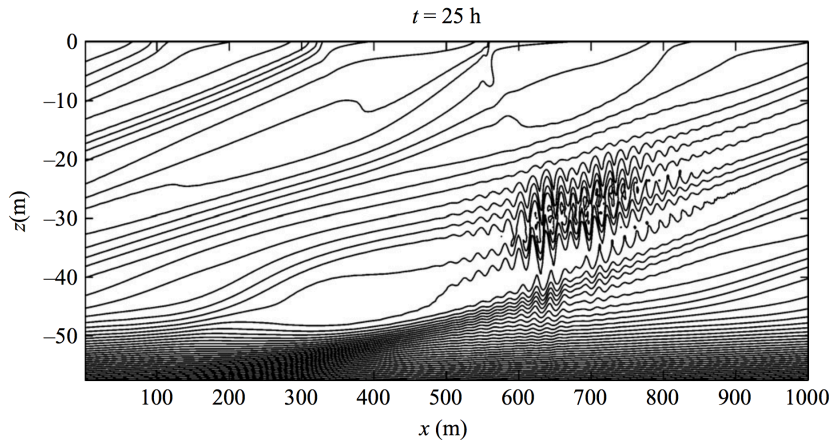
$$\partial_t \mathbf{C}_k = \mathbf{A}_k(\mathbf{U}) \mathbf{C}_k + \mathbf{C}_k (\mathbf{A}_k(\mathbf{U}))^\dagger + \mathbf{Q}_k.$$

B. Farrell & P. Ioannou. Structure and mechanism in a second-order statistical state dynamics model of self-sustaining turbulence in plane Couette flow. arXiv:1607.05020, 2016.

Beyond QL Dynamics: Multiple Scales in t and x and/or z 

- Desirable to extend model (i.e. beyond a QL representation) either to enhance accuracy or efficiency or both
- Develop multiscale numerical algorithm based on retention of multiple space and time scales

Other Applications: Secondary Shear Instability of Symmetrically-Unstable Fronts



J. Taylor & R. Ferrari. On the equilibration of a symmetrically unstable front via a secondary shear instability (JFM 2009).