

Stratocumulus and Stable Layers in the Solar System (boundaries optional)

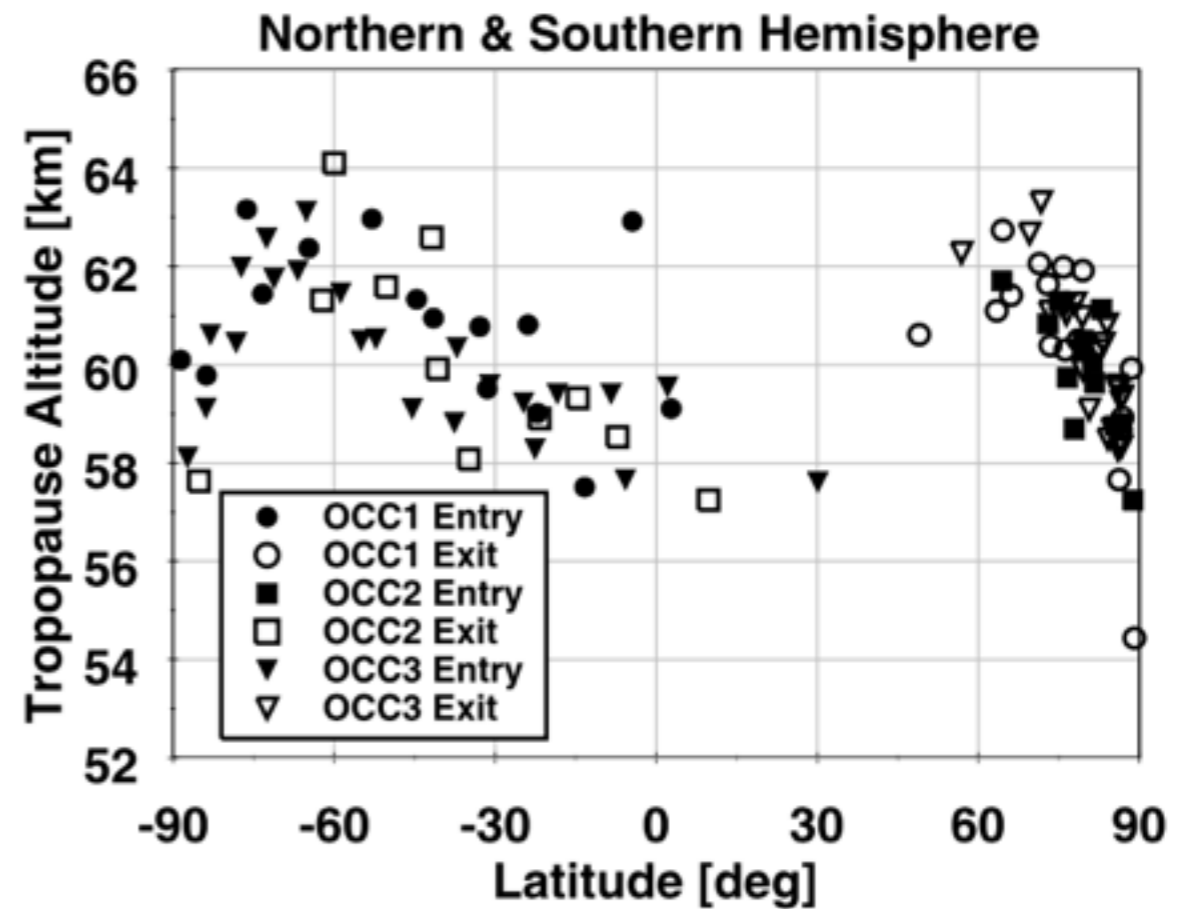
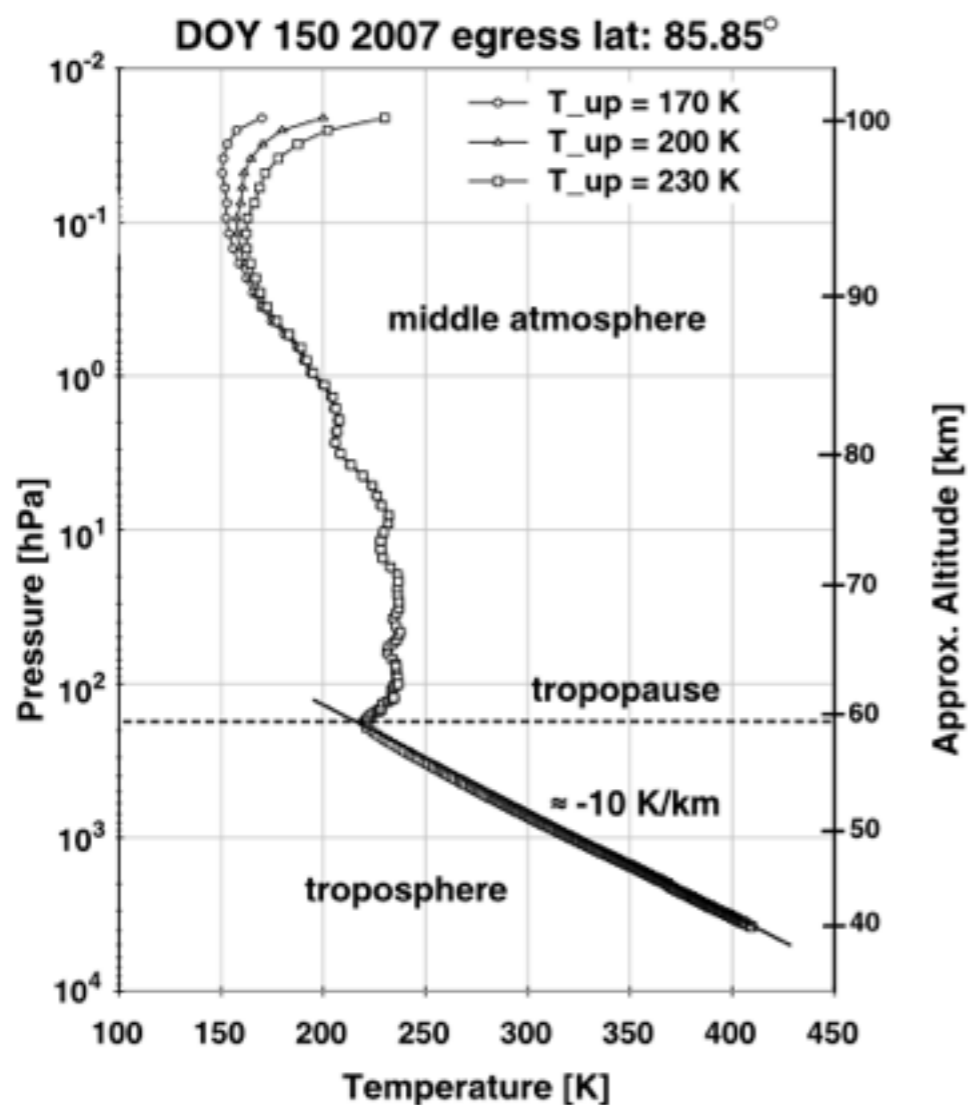
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[work in progress]

My original motivation

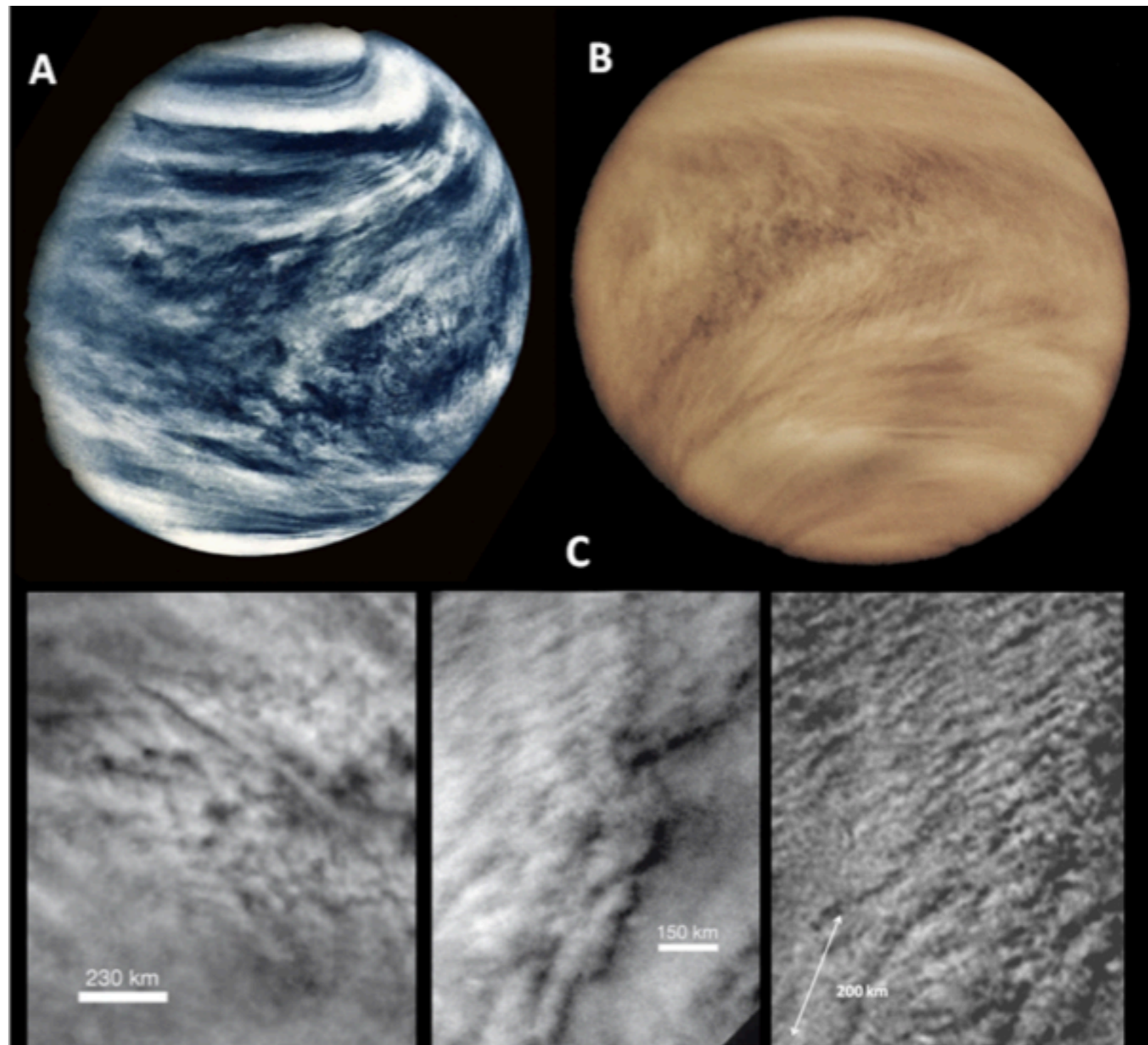
- Venus' sulfuric acid clouds deepen with lowering insolation, just like marine stratocumulus



Tellman et al. 2009

Imamura et al. 2014

**Venus' clouds are global, variable,
and account for a ~70% albedo**



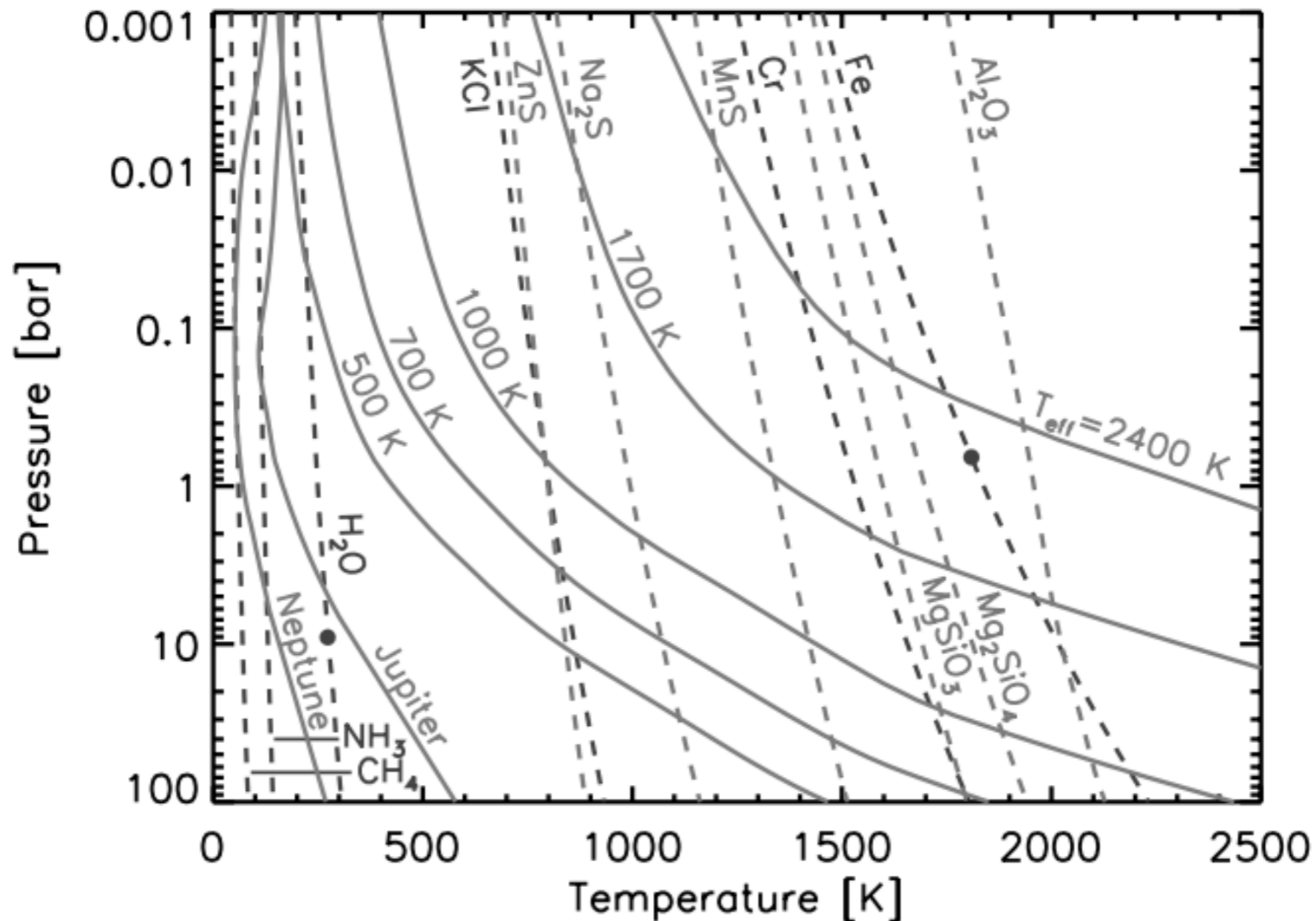
Venus in UV; Sanchez-Lavega et al. 2017

The big picture: What sets planetary albedo?

- Stratocumulus covers more area than any other cloud type on Earth (Wood 2012)
- Earth is an outlier, in the sense of having a large clear-sky component (Mars ~ doesn't have an atmosphere)
 - Venus — global cloud deck of sulfuric acid, albedo of ~70%
 - Titan — global stratospheric smog layer, albedo of ~55%
- Jupiter, Saturn, Uranus, Neptune have global and/or alternating zones of clouds

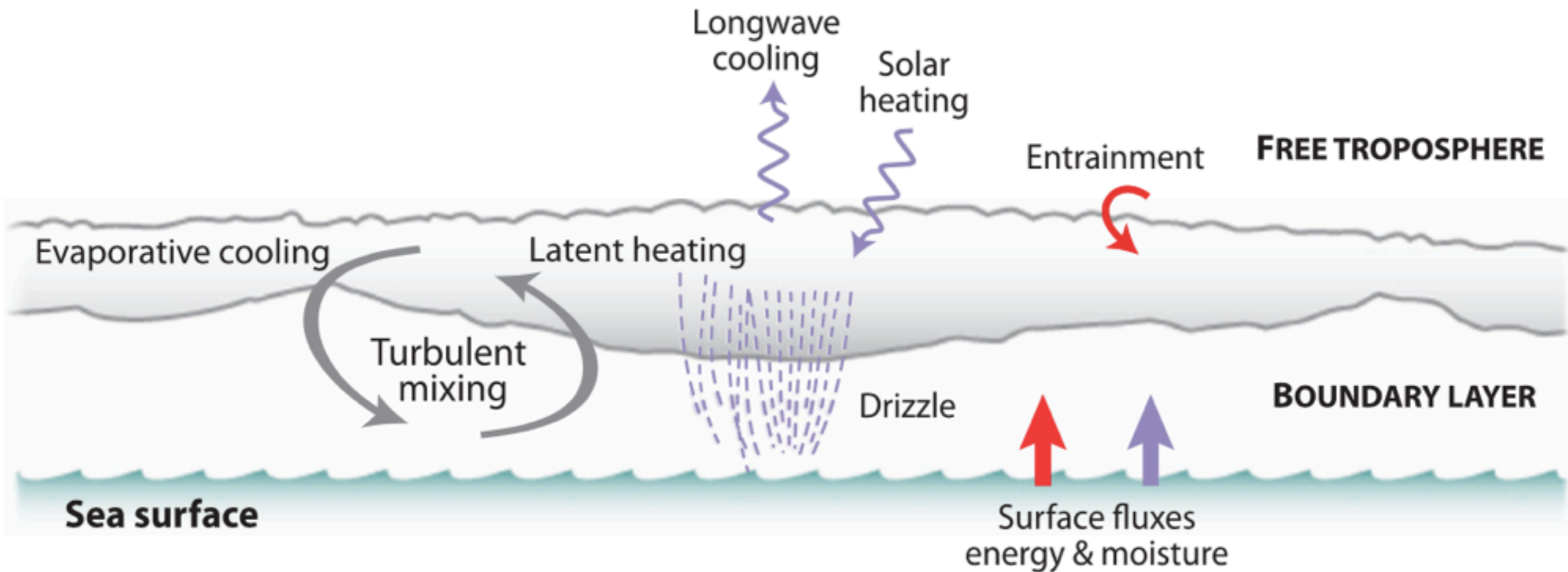
Exoplanet atmospheres have clouds of iron and rock...

If so, how do they form? Are they global? Diurnal?



Fortney, submitted

Stratocumulus forms in stably stratified layers by topside longwave cooling



A rough framework for cloud formation in stably stratified atmospheres

- Fundamentally, stratocumulus is driven by radiative cooling from the top.
- Clouds need to be optically thick to generate the driving buoyancy flux.
- A minimum flux of kinetic energy is required to prevent hydrometeors from settling out.
- In a bulk sense, radiative cooling balances latent and sensible heat fluxes and generates kinetic energy with some (given) efficiency, which must (at least in part) dissipate on hydrometeors.

Preliminaries: BL heat-engine efficiency

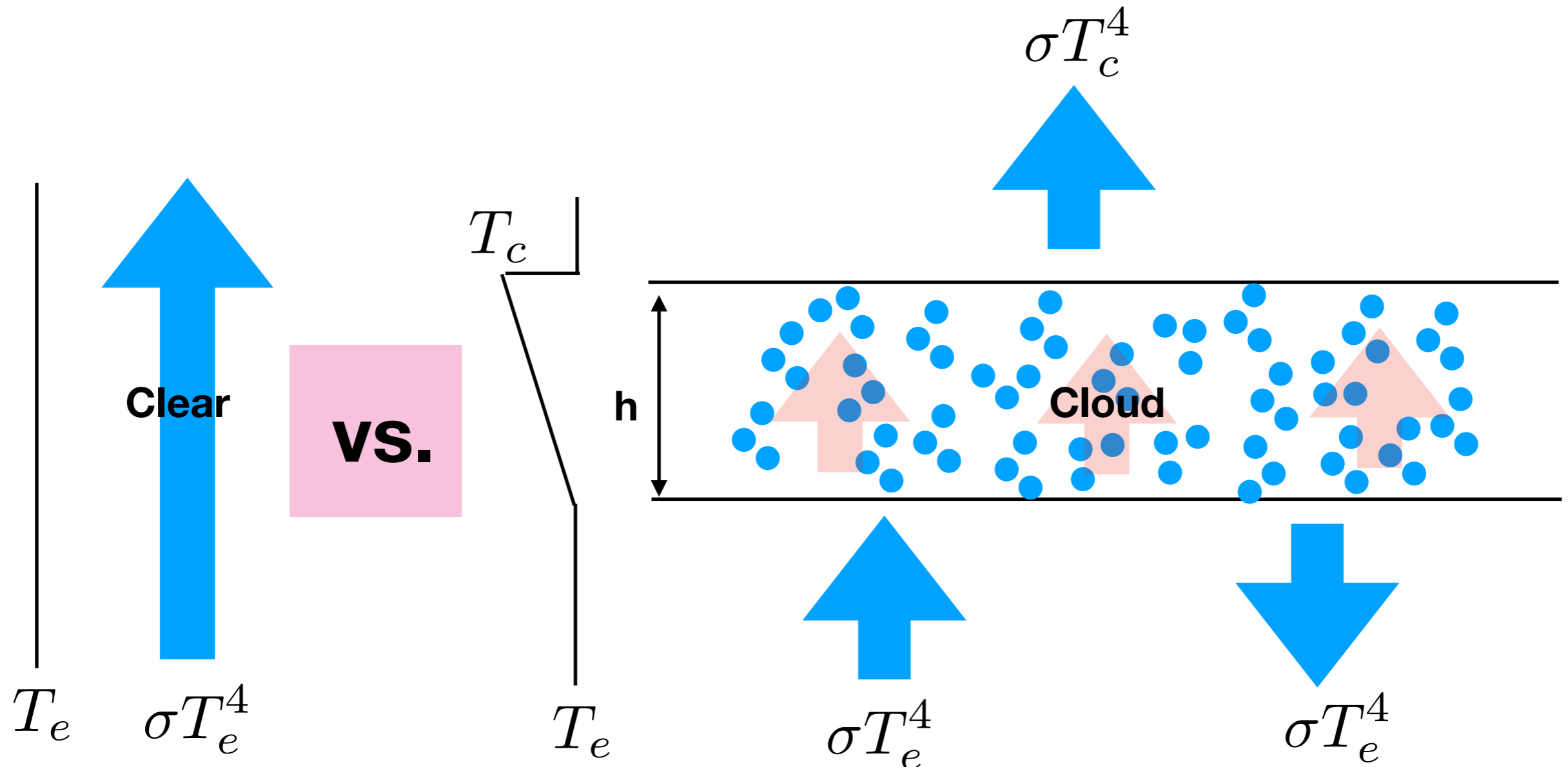
$$F_{\theta} \sim \rho w C_p d\theta \qquad b \equiv \left(g \frac{d\theta}{\theta} \right) \sim w^2 / h$$

↓ ↓

$$F_{\theta} \sim \rho w \left(g \frac{d\theta}{\theta} \right) \frac{C_p \theta}{g}$$
$$\sim \rho w^3 \frac{C_p \theta}{gh}$$
$$\sim \rho w^3 \frac{\theta}{\Delta T_{\text{ad}}}$$

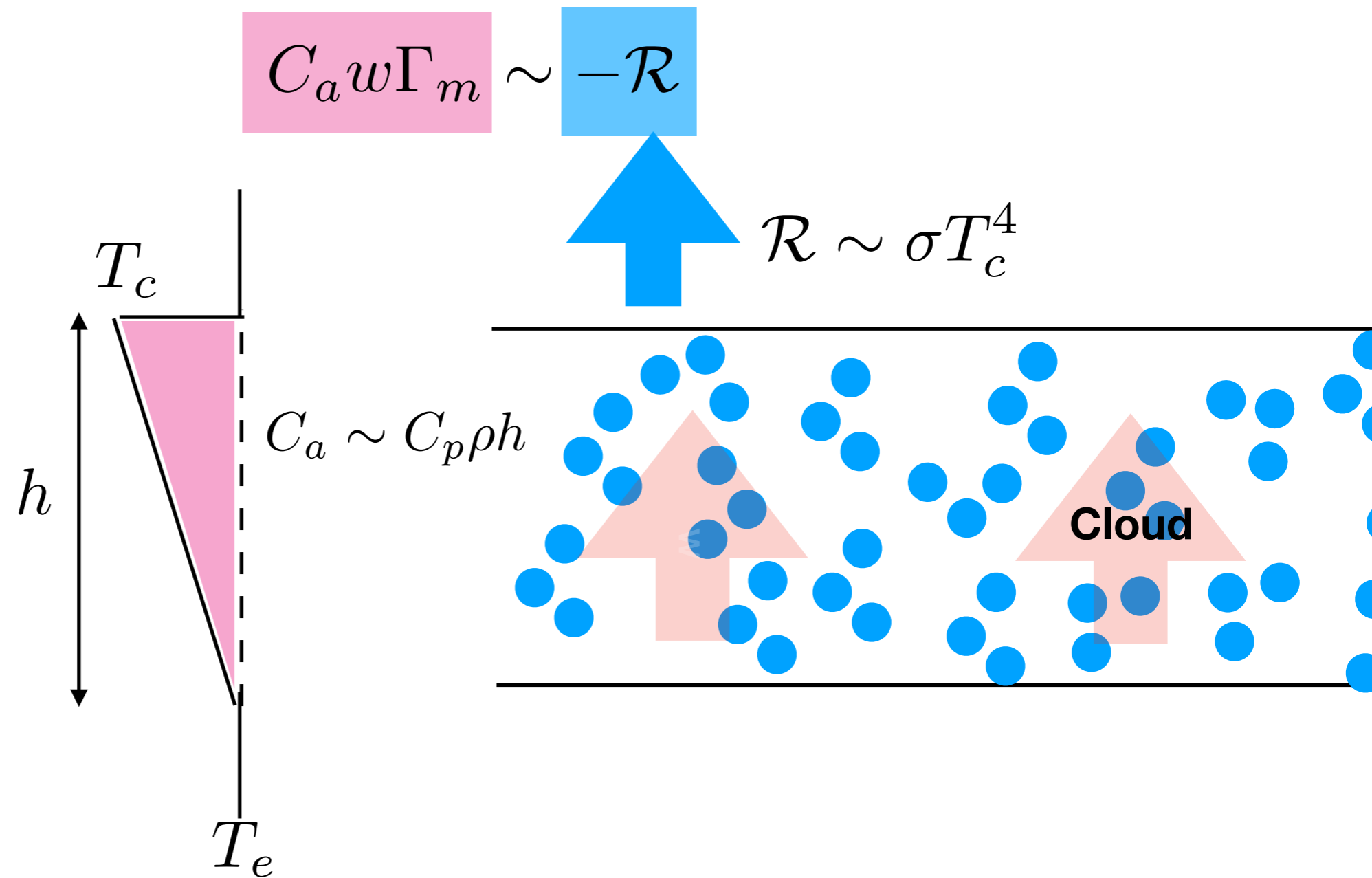
$$F_{\text{KE}} \sim \rho w^3 \sim \eta F_{\theta} \quad \text{with} \quad \eta \equiv \frac{\Delta T_{\text{ad}}}{\theta}$$

Conceptual cloud model



$$T_c = T_e - \Gamma_m h$$
$$T_c^4 \sim T_e^4 (1 - 4\Gamma_m h / T_e)$$

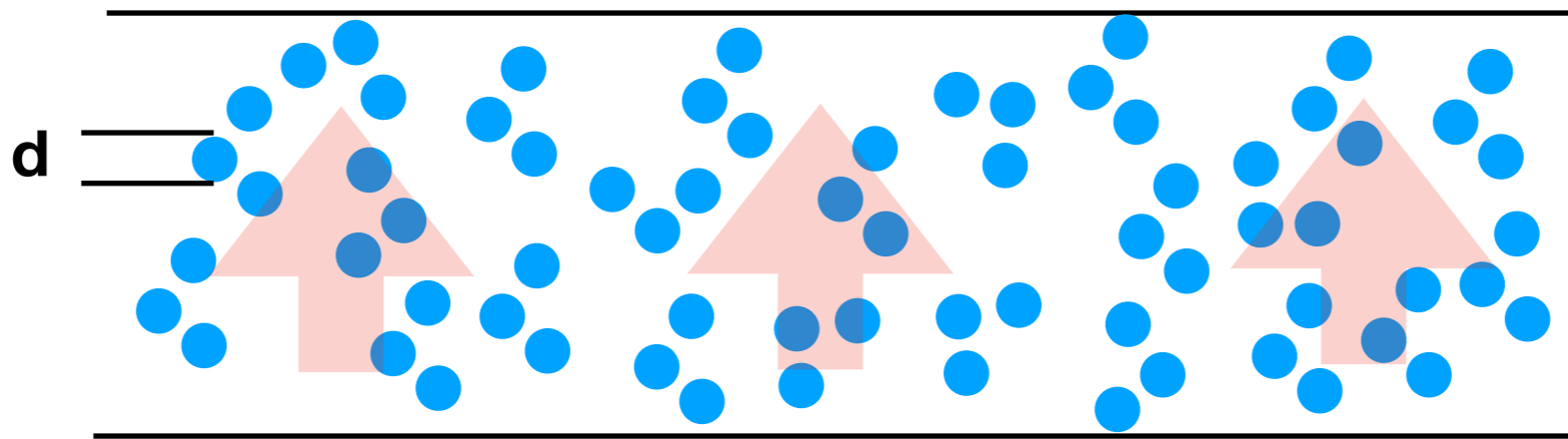
Adiabatic layer must cool



$$T_c^4 \sim T_e^4 (1 - 4\Gamma_m h / T_e)$$

Hydrometeors must not settle

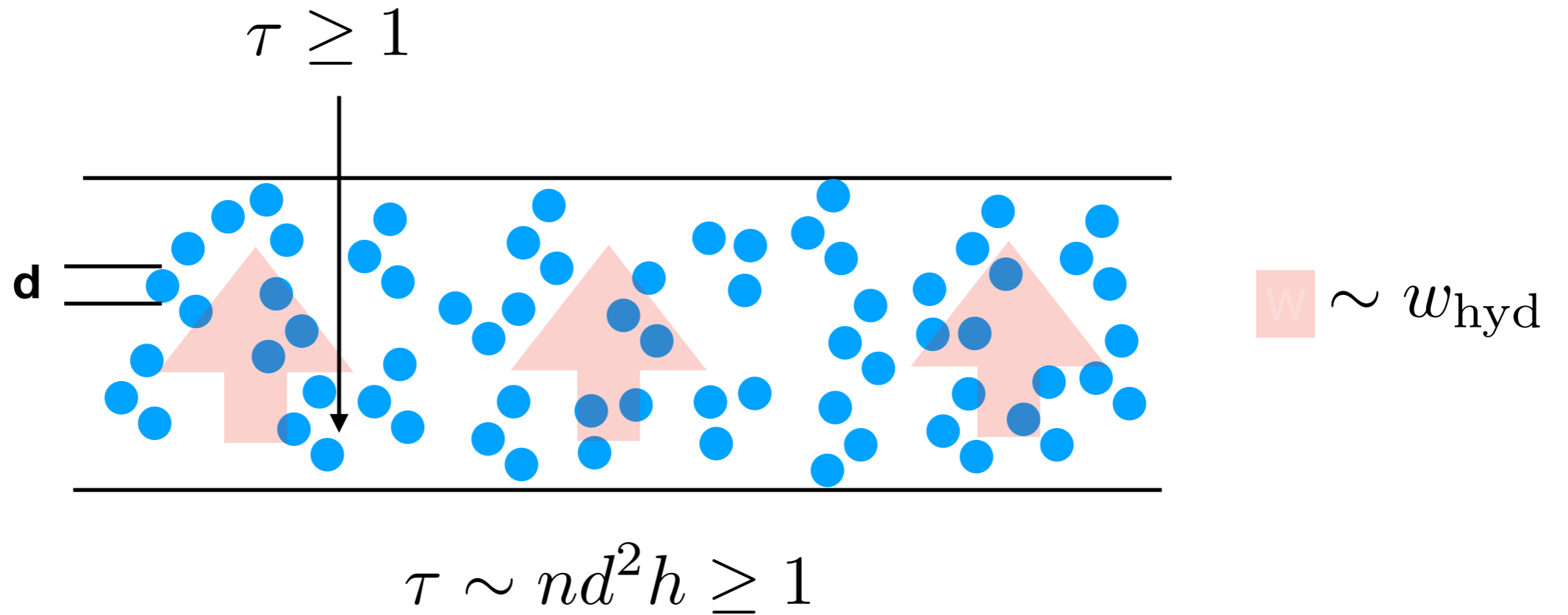
$$W \sim w_{\text{term}}$$



$$b_{\text{hyd}} \sim C_d w_{\text{term}}^2 / d$$
$$\sim C_d w^2 / d$$

$$C_a w \Gamma_m \sim -\mathcal{R}$$
$$T_c^4 \sim T_e^4 (1 - 4\Gamma_m h / T_e)$$

Cloud must be optically thick



Radius ~ 10 microns

Number density ~ 100 - 200 /cc

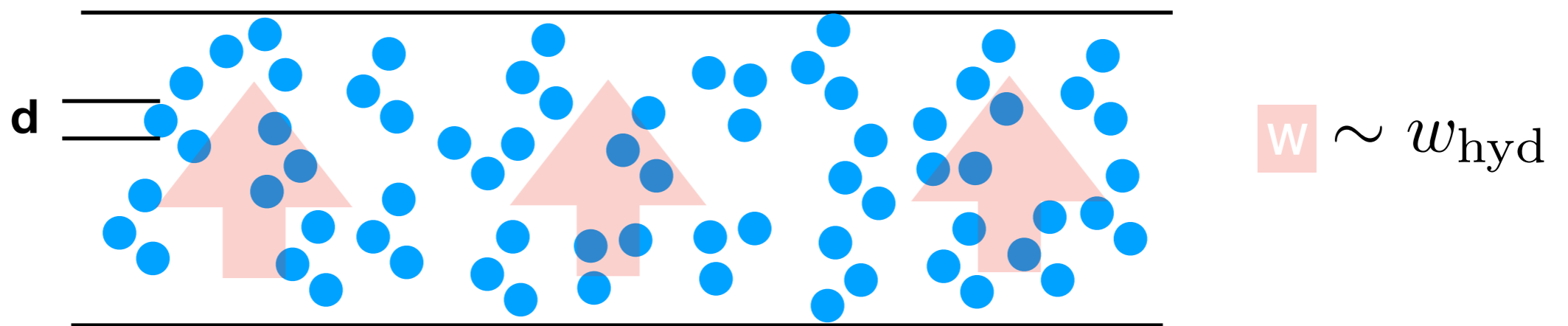
\rightarrow optical depth unity requires ~ 50 - 100 m layer

$$b_{\text{hyd}} \sim C_d w^2 / d$$

$$C_a w \Gamma_m \sim -\mathcal{R}$$

$$T_c^4 \sim T_e^4 (1 - 4\Gamma_m h / T_e)$$

Work done on hydrometeors requires a minimum KE flux



$$F_{\text{KE}} \sim N m_{\text{hyd}} b_{\text{hyd}} w_{\text{hyd}} / A$$

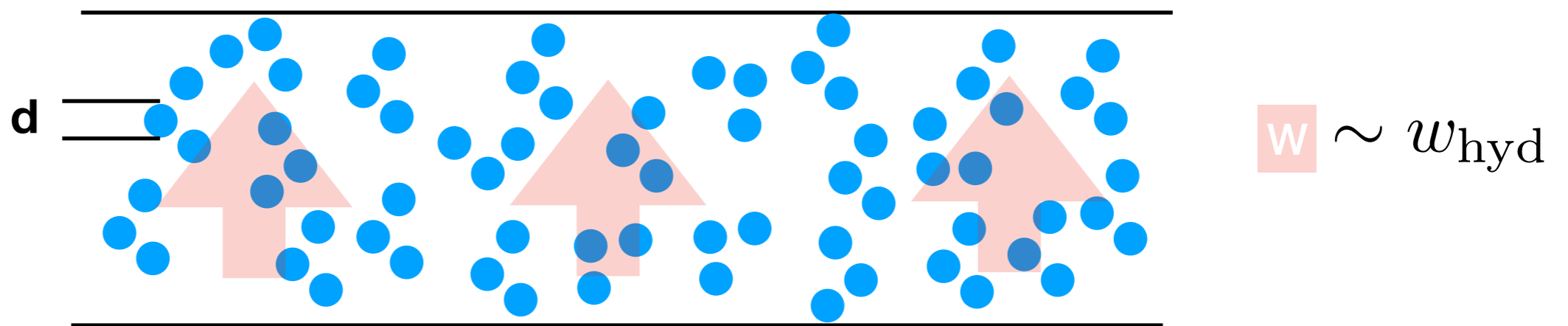
$$\tau \sim n d^2 h \geq 1$$

$$b_{\text{hyd}} \sim C_d w^2 / d$$

$$C_a w \Gamma_m \sim -\mathcal{R}$$

$$T_c^4 \sim T_e^4 (1 - 4 \Gamma_m h / T_e)$$

Work done on hydrometeors requires a minimum KE flux



$$F_{\text{KE}} \sim N m_{\text{hyd}} b_{\text{hyd}} w_{\text{hyd}} / A$$

$$\sim n d^2 h \rho_{\text{hyd}} C_d w^3$$

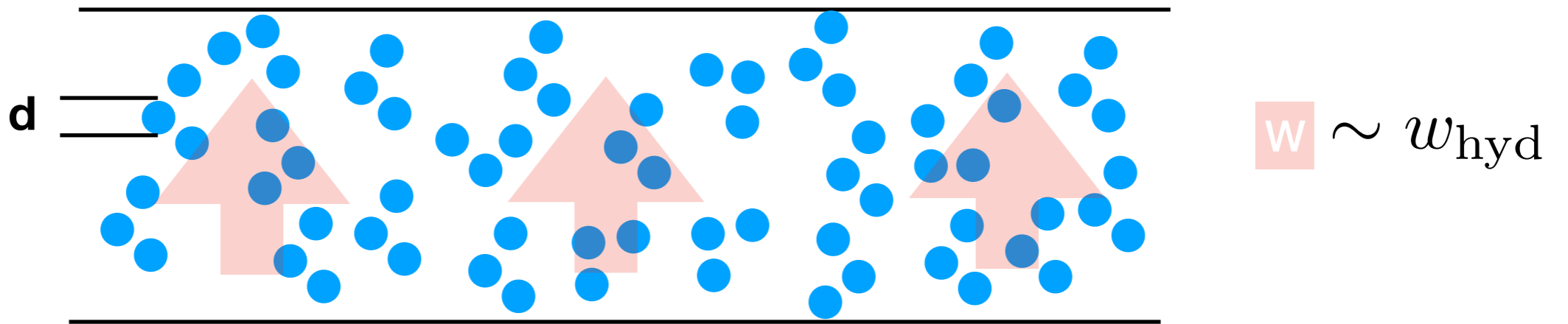
$$\tau \sim n d^2 h \geq 1$$

$$b_{\text{hyd}} \sim C_d w^2 / d$$

$$C_a w \Gamma_m \sim -\mathcal{R}$$

$$T_c^4 \sim T_e^4 (1 - 4 \Gamma_m h / T_e)$$

Putting it together



$$F_{\text{KE}} \sim N m_{\text{hyd}} b_{\text{hyd}} w_{\text{hyd}} / A$$

$$\sim nd^2 h \rho_{\text{hyd}} C_d w^3$$

$$\geq C_d \rho_{\text{hyd}} w^3$$

$$\tau \sim nd^2 h \geq 1$$

$$b_{\text{hyd}} \sim C_d w^2 / d$$

$$C_a w \Gamma_m \sim -\mathcal{R}$$

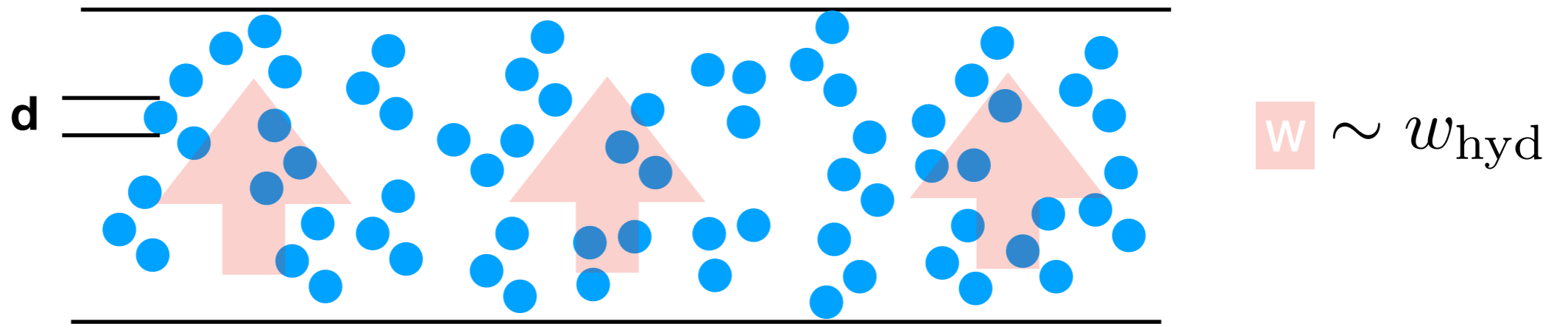
$$T_c^4 \sim T_e^4 (1 - 4\Gamma_m h / T_e)$$

Suspending hydrometeors and being optically thick requires...

$$\frac{1}{C_d} \geq \frac{\rho_{\text{hyd}}}{\rho}$$

Sorta amazing it doesn't depend on droplet size or number density.

Putting it all together



$$\begin{aligned}
 F_{\text{KE}} &\geq C_d \rho_{\text{hyd}} w^3 \\
 &\geq C_d \rho_{\text{hyd}} \left(\frac{-\mathcal{R}}{C_a \Gamma_m} \right)^3
 \end{aligned}$$

$$\begin{aligned}
 \tau &\sim nd^2 h \geq 1 \\
 b_{\text{hyd}} &\sim C_d w^2 / d \\
 C_a w \Gamma_m &\sim -\mathcal{R} \\
 T_c^4 &\sim T_e^4 (1 - 4\Gamma_m h / T_e)
 \end{aligned}$$

Putting it all together

Balance requires $F_{\text{KE}} = \eta F_{\theta} \sim -\eta \mathcal{R}$

$$-\eta \mathcal{R} \geq C_d \rho_{\text{hyd}} \left(\frac{-\mathcal{R}}{C_p \rho h \Gamma_m} \right)^3$$

—or—

$$-\eta (\sigma T_c^4)^2 \geq \frac{C_d \rho_{\text{hyd}}}{(C_p \rho h \Gamma_m)^3}$$

—or—

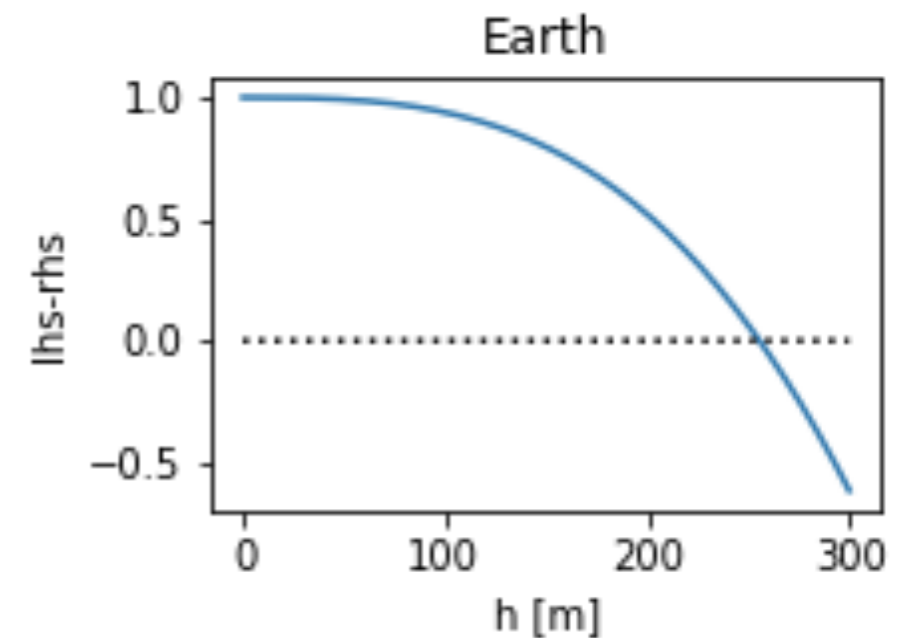
$$1 - 8\Gamma_m h / T_e \geq \eta \frac{(C_p \rho h \Gamma_m)^3}{C_d \rho_{\text{hyd}} \sigma^2 T_e^8}$$

[lhs]

[rhs]

Moist layers: Earth, Venus and Titan*

- Earth -> Few hundred meters (in PBL)
- Titan -> Few meters (in PBL)
- Venus -> Few meters (@40-60km)



$h \sim 250\text{m}$
 $W \sim 1 \text{ m/s}$

- So I'm under-estimating
 - Stratocumulus is a “topping”, not necessarily a “filling” mode
 - Lot's of order unity factors dropped

*Titan's atmosphere isn't optically thin

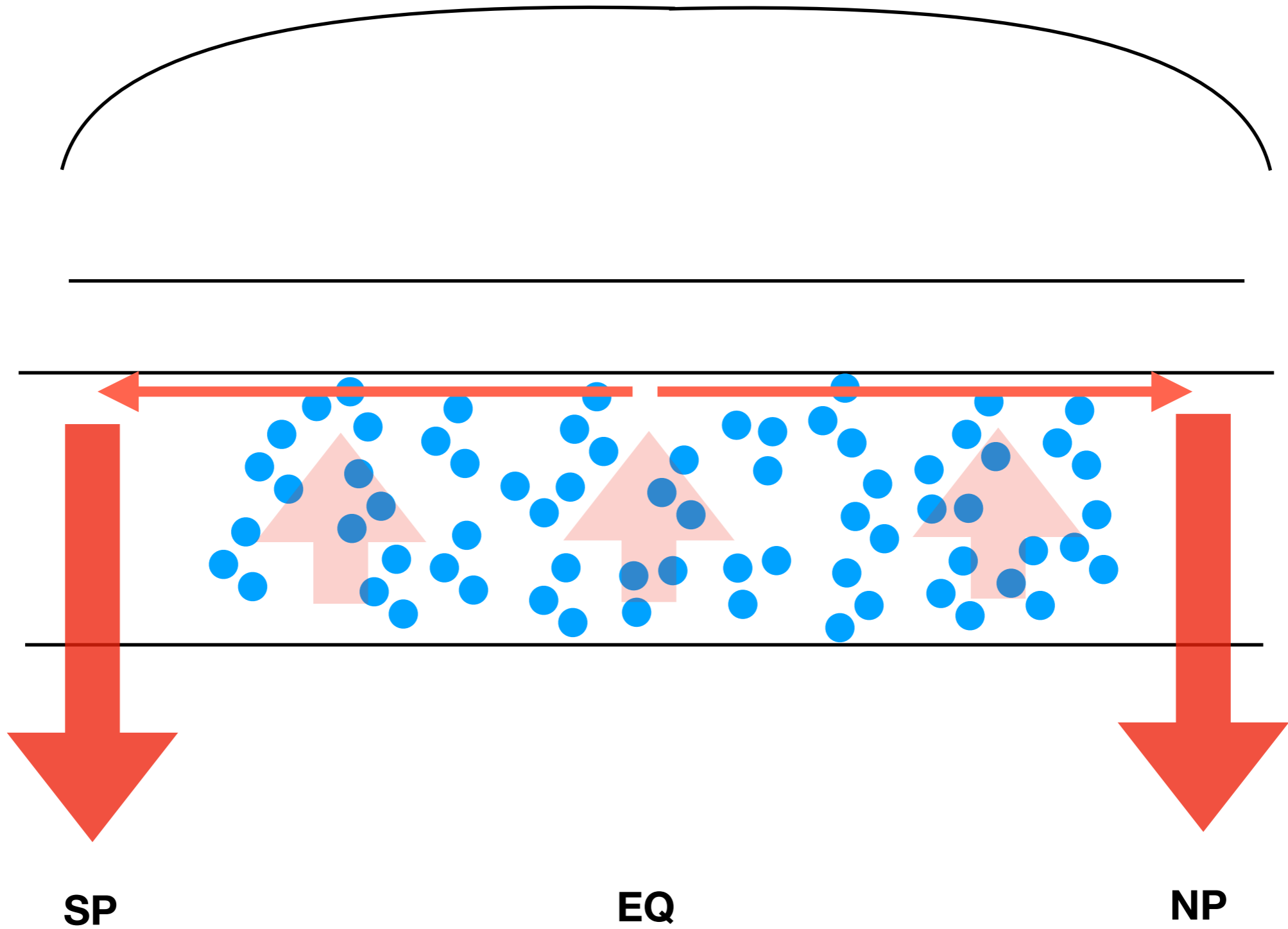
Venus and Jupiter as end-members

- Venus' cloud deck is ~global
 - Except at the poles where large buoyancy flux is removed, causing enhanced downwelling there? [akin to bottom-water formation]
 - Could polar downwelling balance upwelling over the rest of the globe?
 - How does this compare to the implied flux due to cloud granularity observed by Venus Express?
- Jupiter, being rotationally constrained, develops alternating up- and down-welling cells.
 - IR emission is concentrated in clear-sky downwelling regions, especially on the flanks of the equator
 - Polar region may host very deep convection, similarly to Venus. Perhaps we should be thinking about ocean analogies...

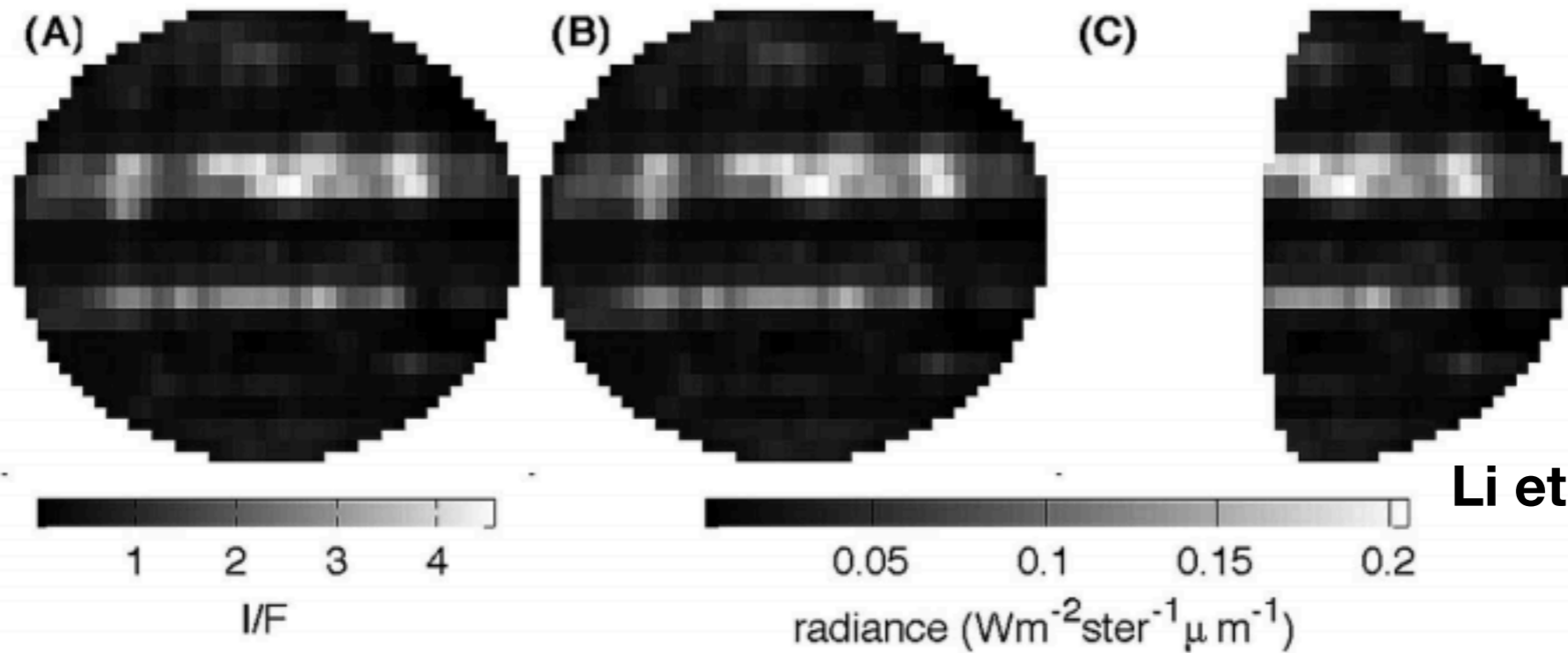
Venus model

Albedo

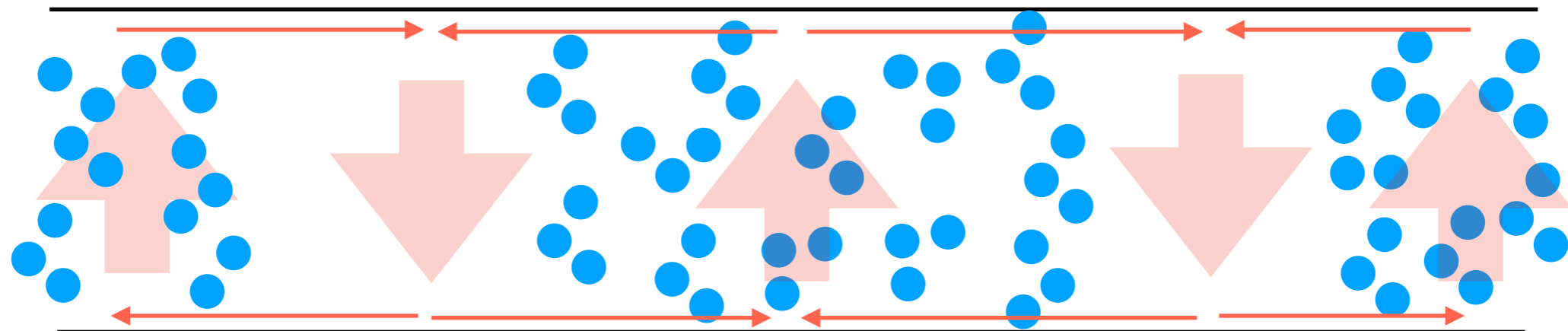
OLR



Jupiter model



Li et al. 2012



[here be monsters]

EQ

Questions

- How does the mass budget close in a (nearly) global mode of stratocumulus?
- What happens in a very hot, moist climate?
- Are there general, guiding principles of planetary albedo?